

The MgTiO₃-FeTiO₃ join at high pressure and temperature

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ABSTRACT

The phase relations at high pressure and high temperature for the FeTiO₃-MgTiO₃ join were determined using several different experimental methods. Through a series of multi-anvil experiments, a phase boundary with a negative slope was observed between MgTiO₃ I (ilmenite structure) and a high pressure phase with the MgTiO₃ II (lithium niobate structure) after quenching. The enthalpy of transformation of MgTiO₃ I to MgTiO₃ II was determined through transposed-temperature-drop calorimetry to be 28.78 ± 1.45 kJ/mol. The enthalpy of transformation from ilmenite to lithium niobate structure was also determined for three intermediate compositions on the FeTiO₃-MgTiO₃ join, Fe_{0.2}Mg_{0.8}TiO₃, Fe_{0.5}Mg_{0.5}TiO₃ and Fe_{0.8}Mg_{0.2}TiO₃, and confirmed for FeTiO₃, and was found to be a linear function of composition. These experiments represent one of the first successful calorimetric measurements on small samples (1 to 3 mg) synthesized at high pressures (15 to 21 GPa). X-ray analysis during compression of Fe_{0.5}Mg_{0.5}TiO₃ II in a diamond cell confirmed a room temperature transition at 28 GPa to Fe_{0.5}Mg_{0.5}TiO₃ III (a GdFeO₃-type perovskite structure), similar to the transitions previously observed in FeTiO₃ and MnTiO₃. The Fe_{0.5}Mg_{0.5}TiO₃ sample was heated to 802 °C at 21 GPa, and it was observed that the stable high temperature, high pressure phase is perovskite, Fe_{0.5}Mg_{0.5}TiO₃ III. The above data combined confirm the stability of a continuous perovskite solid solution at high pressure and temperature for the FeTiO₃-MgTiO₃ join.

INTRODUCTION

Under ambient conditions, both FeTiO₃ and MgTiO₃ are stable in the ilmenite structure ($R\bar{3}$), and complete solid solution occurs between them (e.g., Wood et al. 1991). Syono et al. (1980) described a phase transition at high pressures in FeTiO₃ from the ilmenite structure (FeTiO₃ I) to a hexagonal structure seen in multi-anvil quench experiments and later determined to be the lithium niobate structure (FeTiO₃ II) (Leinenweber et al. 1995; Ko and Prewitt 1988). At 16 GPa FeTiO₃ in this lithium niobate structure transforms to an orthorhombic perovskite polymorph (FeTiO₃ III) (Leinenweber et al. 1991). During quenching this perovskite (FeTiO₃ III) reverted to the lithium niobate structure (Leinenweber et al. 1991). Thermodynamic calculations, based on the observed P - T diagram and on the measured enthalpy of II-to-I transition, indicated the stable FeTiO₃ phase at high pressure has the perovskite structure while the lithium niobate polymorph is a metastable quench phase (Mehta et al. 1994).

A new high-pressure MgTiO₃ phase with the lithium niobate structure ($R3c$) (MgTiO₃ II) was recovered after synthesis at

21 GPa and 1200 °C, and phases with the lithium niobate structure were also synthesized for intermediate compositions between FeTiO₃ and MgTiO₃ (Linton et al. 1997). These magnesium-iron titanates with lithium niobate structures may be metastable quench phases from perovskites. This study shows through calorimetry and X-ray diffraction during compression of MgTiO₃-FeTiO₃ with the lithium niobate structure in a diamond anvil cell (DAC) that these compounds are stable in perovskite-type structures at high pressure. This is the first example of a complete solid solution between iron and magnesium in the perovskite structure.

EXPERIMENTAL METHODS

Synthesis

The geikielite (MgTiO₃ I) starting material was synthesized from stoichiometric quantities of MgO and TiO₂ ground together under ethanol for one hour. Both MgO and TiO₂ were dried prior to weighing in Pt crucibles at 1400 and 1150 °C, respectively, for ~24 hours. The mixture was first heated at 1200 °C for 48 hours and then reground and heated for 24 hours at 1450 °C. The temperature was reduced to 800 °C and the sample removed and cooled in a desiccator. X-ray diffraction patterns of the product showed only geikielite peaks.

MgTiO₃-FeTiO₃ ilmenite solid solutions were synthesized from mixtures of predried Fe₂O₃ and TiO₂ plus MgTiO₃. The mixture was ground in an agate mortar for 1 hour, then heated

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in a platinum crucible at 1200 °C. The oxygen fugacity was controlled at 10^{-10.5} atm by a flow of CO and CO₂. To limit iron loss to the platinum, two lots for each composition were sequentially reacted in the same crucible and the first batch was thrown away. The purity of the products was confirmed by XRD, microprobe and Mössbauer analyses.

Samples of MgTiO₃, Fe_{0.2}Mg_{0.8}TiO₃, Fe_{0.5}Mg_{0.5}TiO₃, Fe_{0.8}Mg_{0.2}TiO₃, and FeTiO₃ in the lithium niobate structure were synthesized from ilmenites of the same composition in a multi-anvil press at the Geophysical Laboratory at 21 GPa and 1300 °C, 19 GPa and 1300 °C, 20 GPa and 1300 °C, 15 GPa and 1300 °C, and 15 GPa and 1300 °C, respectively. Experiment durations ranged from 3 to 24 hours. Each entire sample was removed from its capsule onto a quartz low-background plate and X-rayed. A small quantity of sample was reserved for microprobe analysis (no cation inhomogeneities were found in fully reacted material).

Calorimetry

Transposed temperature drop calorimetry was performed using a miniaturized “ultra-sensitive” Calvet-type calorimeter optimized for use with very small samples (Topor and Navrotsky 1992; Navrotsky 1997). The 2 to 2.5 mg pressed pellets of MgTiO₃ I, FeTiO₃ I, MgTiO₃ II, and FeTiO₃ II were dropped from room temperature (22 °C) to a calorimeter temperature of 801 °C. Fe bearing samples were dropped under a flow of argon, which had been purified by a commercial oxygen getter down to 10 ppb O₂. Phase II samples were dropped, calorimetric signal was measured for forty minutes, and the samples were rapidly withdrawn from the calorimeter to room temperature. This was done to avoid any slow further transformation of the samples which might be undetected in the calorimetric signal. The samples were then repelletized, weighed, and dropped again. Occasionally, when the sample was broken into too many small pieces on the first drop, it was impossible to collect enough sample (>0.5 mg) to form a pellet for a second drop.

High-pressure phase equilibrium

The high pressure and temperature experiments were conducted in a Walker-type multi-anvil press at the Geophysical Laboratory of the Carnegie Institution of Washington. Two types of assemblies were used; for 5 to 9 GPa experiments (synthesis of FeTiO₃ II), we used a 18 mm edge length cast MgO octahedron with a graphite heater and gold or iron capsules. For 10–15 GPa experiments, a MgO octahedral pressure medium with 10 mm edge lengths was used, with 0.5–1 mm in diameter by 1–2 mm in length gold capsules separated by a MgO sleeve from a Re foil heater with a thermal gradient of less than 50 °C over the length of the capsule. Temperature was measured with W5%Re-W26%Re (Type C) thermocouples. Eight tungsten carbide cubes with a corner truncation of 11 mm for lower pressures or 5 mm for higher pressures applied a compressive force on the octahedra. Samples were fully pressurized at a rate of 2–5 GPa/hour then heated at 20 to 80 °C/min. At the end of the run, samples were thermally quenched by cutting off power to the furnace, then slowly depressurized. The pressure calibration at high temperature (e.g., 1200 °C) was established by determining transitions for well-known reactions in Bi, SiO₂,

Fe₂SiO₄, Mg₂SiO₄, and CaGeO₃ (Bertka and Fei 1996).

After the runs, samples were embedded in epoxy and polished. Major element mineral analyzes were acquired using a JEOL JXM 8900 microprobe. The acceleration voltage was 15 kV and the beam current was 20 nA. The locations for analysis and imaging were obtained in a backscattered electron mode. X-ray patterns were obtained on a Scintag XDS-2000 powder diffractometer with a solid-state detector equipped with a Ge crystal using CuK α radiation.

Diamond-anvil cell experiments

Polychromatic (white) wiggler synchrotron x-radiation from X-ray beam line X17c at the National Synchrotron Light Source at Brookhaven National Laboratory was used for energy dispersive X-ray diffraction experiments. Samples were contained by a rhenium gasket between two gem-quality diamonds which applied pressure on the samples. The cells were heated by a Mo-wire resistance heater situated inside the cylinder of the cell around the diamond anvils and by a separate platinum wire resistance heater surrounding the cylinder segment of the cell. This combination of heaters allowed the whole sample to be heated uniformly up to ~800–900 °C; for details see Fei (1996). Pressure was determined by including small pieces of gold foil and referring to the *P-V-T* equation of state of gold by Anderson et al. (1989). Because no pressure medium was used and the pressure gradient across the cell could be as great as 10 GPa, spectra were taken including both gold and sample peaks. Temperature was measured with a Pt-Pt13%Rh thermocouple with its contact on the gasket near the sample.

Raman spectra

Raman spectra of MgTiO₃ samples were excited using the 514 nm line from an Ar⁺ ion argon laser, and recorded with an EG&G Model 4000 CCD with acquisition times up to 20 minutes. Spectra were collected from polished samples embedded in epoxy using a petrographic microscope attached to the micro-Raman system (625 \times magnification) to determine whether the samples had transformed from MgTiO₃ I to MgTiO₃ II in phase equilibrium studies.

RESULTS

Calorimetry

Transposed temperature drop measurements were made for FeTiO₃-MgTiO₃ solid solutions in both the ilmenite and lithium niobate structures. The heat effect for each ilmenite listed in Table 1 is simply the heat content of the sample, because ilmenite is stable at the calorimeter temperature of 801 °C. Our values are ~4% lower than the heat contents of MgTiO₃ and

TABLE 1. Heat content of FeTiO₃-MgTiO₃ I (ilmenite structure)

Formula	Heat content* (kJ/mol)
MgTiO ₃	88.58 \pm 2.03(7)
Fe _{0.2} Mg _{0.8} TiO ₃	89.63 \pm 1.06(4)
Fe _{0.5} Mg _{0.5} TiO ₃	91.22 \pm 0.63(6)
Fe _{0.8} Mg _{0.2} TiO ₃	92.06 \pm 1.11(6)
FeTiO ₃	92.56 \pm 0.50(6)

Notes: Error represents two standard deviations of the mean. The values in parentheses is number of experiments for each composition.

* ($H_{298K} - H_{1074K}$).

FeTiO₃ of Robie and Hemingway (1995). This difference appears to be systematic and is possibly due to a problem in our calibration, but it should not significantly affect our results for the enthalpy of transition. The heat effect for the lithium niobates contains the heat content and the heat of transformation from lithium niobate to ilmenite structure. The complete transformation with no amorphization of the sample was attained by rapidly removing samples from the calorimeter after forty minutes at 801 °C and confirmed through optical examination and X-ray diffraction of end-member composition samples. The agreement of the second transposed temperature drop measurement with the heat content of an ilmenite of the same composition is consistent with complete transformation of the lithium niobate to ilmenite structures (Table 2). To calculate the enthalpy of transition (Table 3, Fig. 1) for each composition, the difference was calculated between the transposed temperature drop measurement of the lithium niobate and the average heat content for ilmenite (although the difference between the first and second drop could also be used to give results that agree within error). There is some scatter, but a linear trend between the two end-members is evident. The present enthalpy of transition (I → II) for pure FeTiO₃, -14.52 (2.16) kJ/mol, is within error of the -13.5(1.2) kJ/mol obtained by Mehta et al. (1994), who used much larger pellets (10–30 mg).

TABLE 2. Transposed-temperature-drop calorimetric measurements on FeTiO₃-MgTiO₃ II (lithium niobate structure)

	First drop		Second drop	
	Mass (mg)	Heat content* + heat of transition (kJ/mol)	Mass (mg)	Heat content* (kJ/mol)
MgTiO ₃	1.098	60.52	0.992	87.27
	1.719	59.06	1.536	88.73
Fe _{0.2} Mg _{0.8} TiO ₃	0.862	69.00		
Fe _{0.5} Mg _{0.5} TiO ₃	1.756	69.78	1.499	92.23
	1.375	66.90	1.267	88.11
	0.889	71.11		
Fe _{0.8} Mg _{0.2} TiO ₃	1.239	78.88	0.665	89.92
	0.996	72.45	0.712	86.41
	1.458	79.73	2.137†	91.50
	0.958	71.03	2.137†	91.50
FeTiO ₃	2.481	80.09	2.345†	91.62
	1.419	77.56	2.345†	91.62
	1.445	76.44	1.315	89.09

* ($H_{295K} - H_{1074K}$).

† Pellet made from combination of two reacted MgTiO₃ pellets.

TABLE 3. Reaction enthalpy for I to II transition

	ΔH of transformation (kJ/mol)	No. of expts.
MgTiO ₃	-28.78 ± 1.45	2
Fe _{0.2} Mg _{0.8} TiO ₃	-20.63	1
Fe _{0.5} Mg _{0.5} TiO ₃	-21.96 ± 2.15	3
Fe _{0.8} Mg _{0.2} TiO ₃	-16.54 ± 4.42	4
FeTiO ₃	-14.52 ± 2.16	3

Notes: Calculated using the average difference between the lithium niobate heat content and the average ilmenite heat content of the same composition. Errors are two standard deviations from the mean.

Phase stability

Although the high-pressure experiments are not reversals, they provide an upper limit for the phase boundary (Fig. 2). The reactions occur quickly (some experiments forming the high pressure phase were under 0.5 h). The boundary of the transition, calculated by a least squares fit to all the data, is $P(\text{GPa}) = -0.008(2)T(\text{K}) + 33(4)$. The slope of this boundary is very similar to that found for pure FeTiO₃: $P(\text{GPa}) = -0.01(3)T(\text{K}) + 25(4)$ (Mehta et al. 1994; Syono et al. 1980). The FeTiO₃ I-II transition occurs 8–10 GPa lower than that of MgTiO₃ at comparable temperatures.

Raman spectra

The Raman spectrum of MgTiO₃ II (Fig. 3) is similar to the spectrum of the lithium niobate polymorph of MnTiO₃ (Ko et al. 1989), and also that of its ilmenite polymorph, but the MgTiO₃ II spectrum exhibits more bands due to its lower symmetry. The higher frequency bands also occur at lower energies for the lithium niobate form, which can be associated with higher entropy for the lithium niobate structure (Ko et al. 1989; Kieffer 1985).

Diamond anvil-cell experiments

MgTiO₃ with the lithium niobate structure recovered from multi-anvil experiments was compressed in a DAC without a pressure medium. During compression of Mg_{0.5}Fe_{0.5}TiO₃ II, a phase transformation was observed at room temperature and 30 GPa. Based on comparison to the X-ray peaks of the new phase of FeTiO₃ III (Leinenweber et al. 1991) and MnTiO₃ III (Ross et al. 1989), the new phase was indexed as an orthorhombic perovskite (GdFeO₃-type, space group *Pbnm*). The

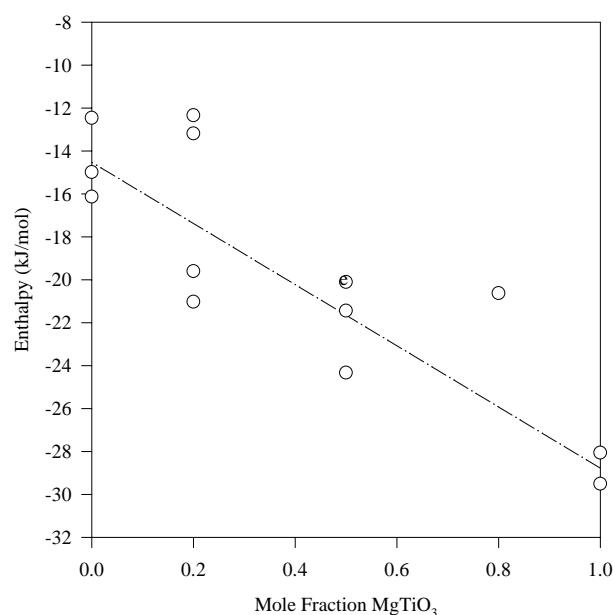


FIGURE 1. Enthalpy of the II to I transitions for FeTiO₃-MgTiO₃ for individual calorimetric experiments from five compositions. Dash-dotted line connects the end-members.

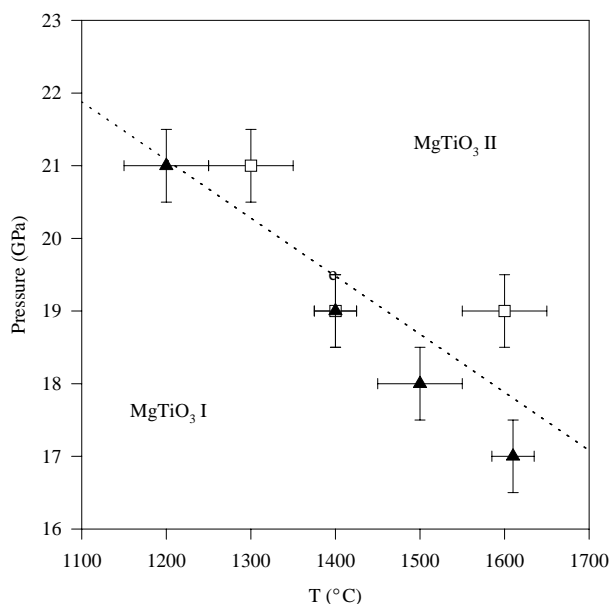


FIGURE 2. Phase diagram for MgTiO₃ I → MgTiO₃ II samples recovered after quench. Filled triangles = MgTiO₃ I; unfilled squares = formation of MgTiO₃ II; the dotted line is our calculated phase boundary: $P(\text{GPa}) = -0.008(2)T(\text{K}) + 33(4)$.

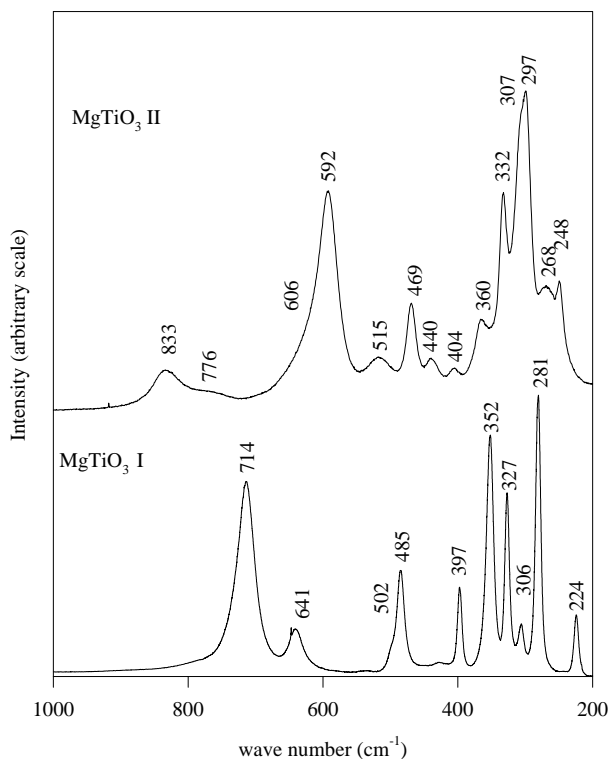


FIGURE 3. Raman spectra of MgTiO₃ I and MgTiO₃ II, taken from polished crystals at ambient pressure.

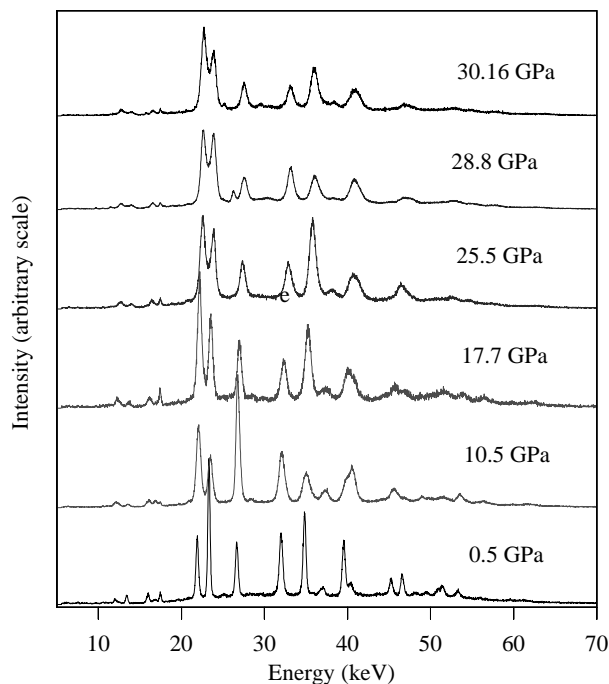


FIGURE 4. Energy-dispersive X-ray diffraction pattern of (Fe_{0.5}Mg_{0.5})TiO₃ II as pressure increases during compression in a DAC at room temperature with no pressure medium. At over 28 GPa the region between ~22 and ~24 E/keV can no longer be fit as just two peaks; a new peak is appearing between the original peaks.

volume of the perovskite is 46.43 Å³ per formula unit at 30 GPa. The most noticeable indication of a phase transition, the appearance of the 112 perovskite peak between the 104 and 110 peaks of the lithium niobate structure, occurs at around 28 GPa at room temperature (Fig. 4).

Upon heating the Mg_{0.5}Fe_{0.5}TiO₃ sample to 802 °C, the transformation became more complete and a single phase perovskite was formed (Fig. 5). The volume of the perovskite at 802 °C and 21 GPa is 46.94(11) Å³ per formula unit (Table 4). After heating, the sample was quenched at pressure from 802 to 300 °C quickly (~30 s), by turning off the small heater inside the cylinder of the DAC, and from 300 °C to room temperature more slowly (greater than one hour for complete cooling) by turning off the large heater external to the cylinder. After complete cooling, the cell was gradually decompressed. During the thermal quenching a two phase mixture of perovskite and ilmenite was formed, the perovskite in this mixture persisted through decompression down to ~8.3 GPa (Fig. 6). Microprobe analysis of the recovered sample from the diamond cell showed segregation into Fe and Mg rich areas. The average compositions were Mg_{0.622(2)}Fe_{0.418(8)}Ti_{0.980(2)}O₃ and Mg_{0.510(20)}Fe_{0.512(20)}Ti_{0.99(10)}O₃.

Phase diagram

From the thermodynamic data determined through calorimetry (Table 5) and some estimated values of physical parameters (Table 6), the phase relations in the FeTiO₃-MgTiO₃ system can be calculated. The estimated values are tested by compar-

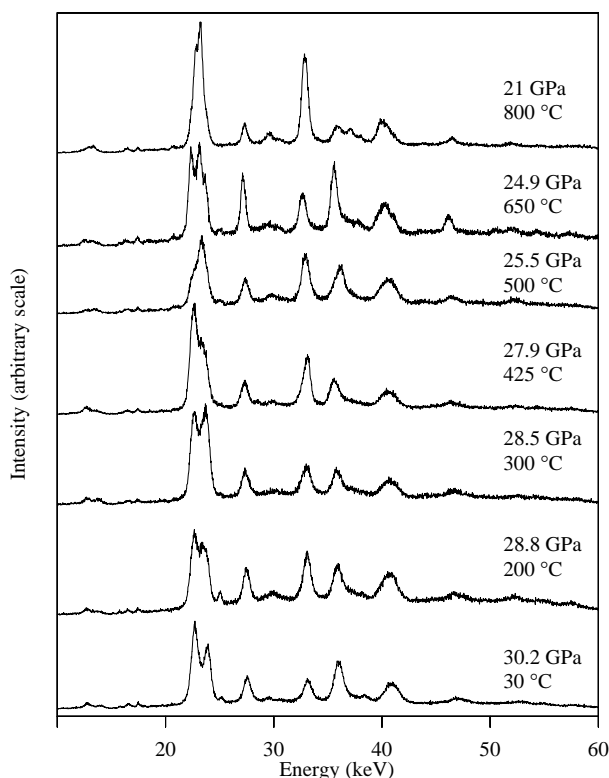


FIGURE 5. Energy-dispersive X-ray diffraction pattern of (Fe_{0.5}Mg_{0.5})TiO₃ II as temperature increases and pressure decreases in the diamond cell. Above 21 GPa and 800 °C the (Fe_{0.5}Mg_{0.5})TiO₃ II has completely transformed to a GdFeO₃-type orthorhombic perovskite [(Fe_{0.5}Mg_{0.5})TiO₃ III].

TABLE 4. X-ray diffraction data for MgTiO₃* at 802 °C and 21 GPa

<i>h</i>	<i>k</i>	<i>l</i>	<i>d</i> (obs) (Å)	<i>d</i> (calc) (Å)	<i>d</i> (diff)	<i>h</i> / <i>l</i> ₀
0	0	2	3.611	3.614	-0.003	3
0	2	0	2.608	2.603	0.005	76
1	1	2	2.553	2.551	0.002	92
2	0	0	2.492	2.495	-0.003	6
1	0	3	2.173	2.170	0.003	19
1	1	3	2.007	2.003	0.004	15
1	2	2	1.943	1.945	-0.003	3
0	0	4	1.805	1.807	-0.002	100
1	2	3	1.666	1.667	-0.002	3
1	3	0	1.646	1.640	0.006	12
1	3	1	1.599	1.599	0.001	15
0	3	2	1.560	1.564	-0.004	3
0	2	4	1.486	1.484	0.001	18
2	0	4	1.464	1.463	0.001	3
3	1	2	1.449	1.451	-0.002	8
2	2	4	1.277	1.276	0.001	8
1	1	6	1.143	1.143	0.000	6

Notes: *a* = 4.9852(69) Å, *b* = 5.2104(43) Å, *c* = 7.2305(48) Å, *V* = 187.75(45) Å³

* GdFeO₃-type perovskite structure.

son to the phase stability data for the end-members.

For the MgTiO₃ end-member, thermodynamic data were constrained by the phase stability and the enthalpy of transition from MgTiO₃ I to MgTiO₃ II, both from this study. It was first necessary to establish whether the phase boundary, determined through multi-anvil experiments, represented a MgTiO₃ I to MgTiO₃ II transition or a MgTiO₃ I to MgTiO₃ III transition

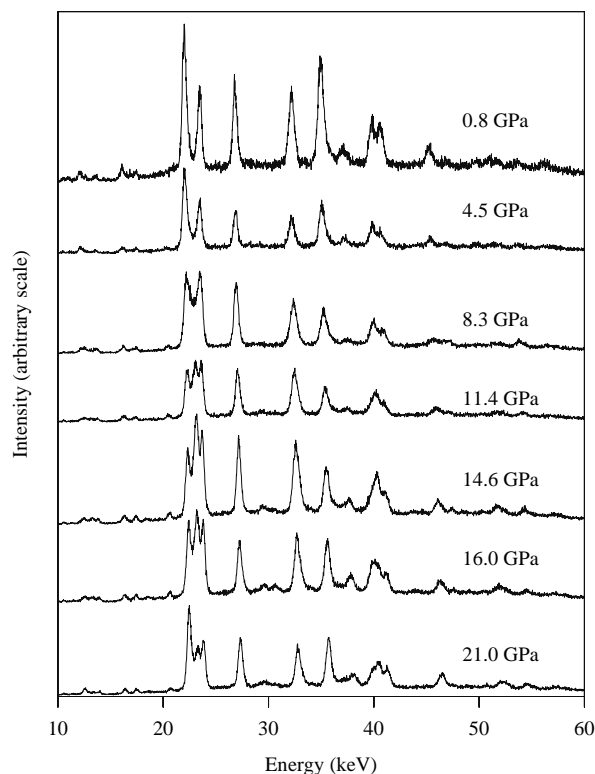


FIGURE 6. Energy-dispersive diffraction spectrum of (Fe_{0.5}Mg_{0.5})TiO₃ II as pressure decreases at room temperature in the DAC. Two phases a perovskite and an ilmenite are distinguishable down to ~8.3 GPa.

TABLE 5. Thermochemical data used in phase equilibrium calculations

Transition	ΔH_{298}^0 (kJ/mol)	S^0 (J/K-mol)
FeTiO ₃ (I) → FeTiO ₃ (II)	14.5 ± 2.2*	4.9 ± 3.1
MgTiO ₃ (I) → MgTiO ₃ (II)	28.8 ± 1.0†	4.0 ± 2.6
FeTiO ₃ (I) → FeTiO ₃ (III)	27.1 ± 4.0‡	8.9 ± 3.1 ^c
MgTiO ₃ (I) → MgTiO ₃ (III)	33.1 ± 5.1§	5.5 ± 2.6 ^d
FeTiO ₃ (II) → FeTiO ₃ (III)	12.3 ± 5.9*,‡	4.0 ± 4
MgTiO ₃ (II) → MgTiO ₃ (III)	4.4 ± 6.1†,§	1.5 ± 2

* Calorimetry (Mehta et al. 1994).

† Calorimetry, this study.

‡ High pressure-temperature phase equilibrium, (Syono et al. 1980).

§ High pressure-temperature phase equilibrium, this study.

TABLE 6. Physical properties used in phase equilibrium calculations

Compound	Structure	<i>V</i> ₂₉₈ ⁰ (cm ³ /mol)	<i>K</i> (GPa)
MgTiO ₃	ilmenite	30.87*	169#
FeTiO ₃	ilmenite	31.70†	177(3)†
MgTiO ₃	lithium niobate	30.71‡	172
FeTiO ₃	lithium niobate	31.40§	182(7)**
MgTiO ₃	perovskite	29.31	225
FeTiO ₃	perovskite	30.15	236††

Notes: *dK/dP* assumed to be 4. Thermal expansion was assumed equal to that of FeTiO₃ ilm, 27.9×10^{-6} /K [Wechsler and Prewitt (1984)].

* Wechsler and Von Dreele (1989).

† Wechsler and Prewitt (1984).

‡ Linton et al. (1997).

§ Leinenweber et al. (1995).

|| estimated, see text.

Liebermann (1976).

** Mehta et al. (1994).

†† Leinenweber et al. (1991).

with MgTiO₃ II formed metastably during the quench. If we assume that the molar volume change is not a strong function of temperature or pressure for the MgTiO₃ I to MgTiO₃ II transition [Mehta et al. (1994) showed that ΔV is independent of pressure for FeTiO₃], then we can simplify the equation for the free energy of a phase change to:

$$\Delta G(P,T) \approx \Delta H^0 - T\Delta S^0 + P\Delta V^0 \quad (1)$$

At the phase boundary, where $\Delta G = 0$:

$$P \approx \left(\frac{\Delta S}{\Delta V} \right) T - \left(\frac{\Delta H}{\Delta V} \right) \quad (2)$$

For MgTiO₃ I \rightarrow MgTiO₃ II, $\Delta H_{\text{ilm} \rightarrow \text{lb}} = 28.78 \pm 1.45$ kJ/mol (this study), $\Delta V_{\text{ilm} \rightarrow \text{lb}}^0 = -0.16$ cm³/mol (Linton et al. 1997), and we calculate the $T = 0$ K intercept of the transition to be $(\Delta H/\Delta V) = 179 \pm 1$ GPa. This value can be compared to the $T = 0$ K intercept from the equation of the phase boundary obtained in this study: $P(\text{GPa}) = -0.008(2)T(\text{K}) + 33(4)$ (GPa). Clearly the 179 ± 1 GPa calculated is much higher than the 33 ± 4 GPa observed, and the phase boundary from multi-anvil experiments is not the MgTiO₃ I to MgTiO₃ II boundary; rather, analogously to FeTiO₃ (Mehta et al. 1994), it is most likely the MgTiO₃ I to MgTiO₃ III boundary with MgTiO₃ II as a quench product.

The phase diagram of the FeTiO₃-MgTiO₃ join at pressure can be calculated by simultaneously solving the following equations;

$$\Delta H_{\text{MgTiO}_3}^0(\text{il} \rightarrow \text{pv}) - T\Delta S_{\text{MgTiO}_3}^0(\text{il} \rightarrow \text{pv}) + \int_{\text{latm}}^P \Delta V_{\text{MgTiO}_3}^0(\text{il} \rightarrow \text{pv})dP + RT \ln \left(\frac{a_{\text{MgTiO}_3}^{\text{pv}}}{a_{\text{MgTiO}_3}^{\text{il}}} \right) = 0 \quad (3)$$

$$\Delta H_{\text{FeTiO}_3}^0(\text{il} \rightarrow \text{pv}) - T\Delta S_{\text{FeTiO}_3}^0(\text{il} \rightarrow \text{pv}) + \int_{\text{latm}}^P \Delta V_{\text{FeTiO}_3}^0(\text{il} \rightarrow \text{pv})dP + RT \ln \left(\frac{a_{\text{FeTiO}_3}^{\text{pv}}}{a_{\text{FeTiO}_3}^{\text{il}}} \right) = 0 \quad (4)$$

where S is entropy and a is activity. We model the activity by assuming the ideal case $\gamma = 1$, where the activity, a_i , simply equals the mole fraction, X_j , or another approach employs nonideal mixing models. The MgTiO₃-FeTiO₃ join for the ilmenite structure has been described by a symmetrical regular mixing model, $RT \ln \gamma_i = W_G(1 - X\alpha_i)^2$, where $W_G = 4300$ J/mol, based on partitioning of Mg and Fe between coexisting ilmenites and olivines (Andersen et al. 1991). Using the method of Davies and Navrotsky (1983), which correlates the interaction parameter with the volume mismatch of the two end-members, we estimated symmetric W_G values as 2278 J/mol, 2163 J/mol, and 1824 J/mol, respectively, for Mg-Fe mixing in the ilmenite, perovskite, and lithium niobate structures. With these nonideal mixing models, Equations 3 and 4 become:

$$\Delta H_{\text{MgTiO}_3}^0(\text{il} \rightarrow \text{pv}) - T\Delta S_{\text{MgTiO}_3}^0(\text{il} \rightarrow \text{pv}) + \int_{\text{latm}}^P \Delta V_{\text{MgTiO}_3}^0(\text{il} \rightarrow \text{pv})dP + RT \ln \left(\frac{X_{\text{MgTiO}_3}^{\text{pv}}}{X_{\text{MgTiO}_3}^{\text{il}}} \right) - W_G(X_{\text{FeTiO}_3}^{\text{il}})^2 + W_{G_{\text{pv}}}(X_{\text{FeTiO}_3}^{\text{pv}})^2 = 0 \quad (5)$$

and:

$$\Delta H_{\text{FeTiO}_3}^0(\text{il} \rightarrow \text{pv}) - T\Delta S_{\text{FeTiO}_3}^0(\text{il} \rightarrow \text{pv}) + \int_{\text{latm}}^P \Delta V_{\text{FeTiO}_3}^0(\text{il} \rightarrow \text{pv})dP + RT \ln \left(\frac{X_{\text{FeTiO}_3}^{\text{pv}}}{X_{\text{FeTiO}_3}^{\text{il}}} \right) - W_G(X_{\text{MgTiO}_3}^{\text{il}})^2 + W_{G_{\text{pv}}}(X_{\text{MgTiO}_3}^{\text{pv}})^2 = 0 \quad (6)$$

These two equations were solved for different values of $W_G(\text{il})$ and $W_G(\text{pv})$ using Newton's method, with starting values for composition from the exact solutions of Equation 3 assuming ideal mixing, to obtain the phase diagram at 802 °C (Fig. 7). Ideal mixing best predicts the compositions of the two phases in the recovered sample from the DAC, although that experiment does not necessarily define an equilibrium pair of compositions at any P and T .

DISCUSSION

Petrological implications

The high-pressure phase diagram of the FeTiO₃-MgTiO₃ join can be applied to ultra-high pressure metamorphic rocks and kimberlites. Dobrzhinetskaya et al. (1996) found ilmenite inclu-

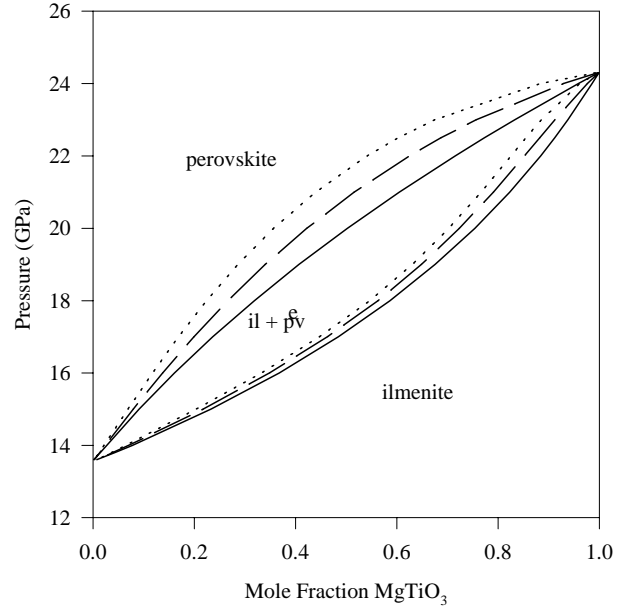


FIGURE 7. Calculated phase diagram for the FeTiO₃-MgTiO₃ join at 800 °C. The two phase loops for the ilmenite to perovskite transition were calculated using $W_G(\text{il}) = 4300$ J/mol and $W_G(\text{pv}) = 4300$ J/mol (dotted curve), $W_G(\text{il}) = 2278$ J/mol and $W_G(\text{pv}) = 2163$ J/mol (dashed curve). The solid curve represents ideal mixing.

sions with unique structures in olivine from ultra-high-pressure rocks from the Alpe Arami peridotite; garnet-orthopyroxene-clinopyroxene equilibration in these rocks indicates pressures from 4 to 5 GPa and temperatures from 900 to 1000 °C. However, the volume of exsolved chromite and FeTiO₃ in olivines was greater than could be explained by any documented substitution of Cr³⁺ and Ti³⁺ into the olivine structure, and Dobrzhinetskaya et al. (1996) postulated that the inclusions reflect an even higher pressure precursor phase, such as (Mg_{0.9}Fe_{0.1})₂SiO₄ in the β-phase, which accommodates a higher substitution of trivalent cations. This implies that these rocks reached pressures above 14 GPa, high enough to stabilize iron-rich (Fe,Mg)TiO₃ III. The topotaxially oriented (Fe,Mg)TiO₃ rod-shaped precipitates in olivine have orthorhombic crystal structures with lattice parameters: $a = 0.470$ nm, $b = 1.028$ nm, $c = 0.574$ nm. Dobrzhinetskaya et al. (1996) propose that during uplift β-phase transformed to olivine from which crystallographically oriented (Fe,Mg)TiO₃ III exsolved [(Fe,Mg)TiO₃ I exsolves from olivine with no preferred orientation]. The observed orthorhombic crystals would then be metastable decomposition structures from (Fe,Mg)TiO₃ III.

(Fe,Mg)TiO₃ II was not observed, possibly due to topotaxial constraints and/or slower cooling rates which might avoid this metastable phase. For the set of reactions to occur as Dobrzhinetskaya et al. (1996) propose, the transformation from β-phase to olivine must occur before the transformation of (Fe,Mg)TiO₃ III to ilmenite (I). Using the phase boundary of Syono et al. (1980) for pure FeTiO₃, this is a certainty, but from our study it becomes plain that which transformation occurs first during uplift is very much a function of the MgTiO₃ content of the (Fe,Mg)TiO₃ III, which Dobrzhinetskaya et al. (1996) were not able to accurately determine. Whether or not the (Fe,Mg)TiO₃ structures of the Alpe Arami are remnants of the (Fe,Mg)TiO₃ III phase, they are a good example of the petrological importance of the high *P* and *T* phase relations of the FeTiO₃-MgTiO₃ join.

Two other (Fe,Mg)TiO₃ phases exist in the natural olivines that Dobrzhinetskaya et al. (1996) conjecture are intermediates between the orthorhombic (Fe,Mg)TiO₃ and (Fe,Mg)TiO₃ I. One of these intermediates has reported lattice parameters of $a \cong 0.470$ nm, $b \cong 0.5014$ nm, $c \cong 0.574$ nm. This phase has an even smaller unit cell volume than the Mg_{0.5}Fe_{0.5}TiO₃ III observed in our DAC studies. The persistence of perovskite during the slow cooling of the diamond cell from 300 °C to room temperature indicates that at low temperatures perovskite may persist metastably to pressures below the ilmenite-perovskite transition. If the olivine were exerting a confining pressure on the (Fe,Mg)TiO₃ phase, it might be worthwhile to reexamine the TEM data of Dobrzhinetskaya et al. (1996) keeping in mind the possibility of metastable perovskite structures.

Systematics of ilmenite-lithium niobate-perovskite phase transitions

Previous X-ray diffraction of MnTiO₃ and FeTiO₃ in a DAC showed that the lithium niobate structure (II) phase transformed to a perovskite (III) at high pressure, but some uncertainty remained as to which structure was stable at high temperature and high pressure (Ross et al. 1989; Leinenweber et al. 1991).

Whereas Mehta et al. (1994) concluded FeTiO₃ III was the stable phase at high pressure and temperature, Ko et al. (1989) suggested that MnTiO₃ II, which was stable under high-temperature and high-pressure conditions and that MnTiO₃ III was only stable under high-pressure, low-temperature conditions. Here, Fe_{0.5}Mg_{0.5}TiO₃ sample show, Fe_{0.5}Mg_{0.5}TiO₃ III is the stable phase at high temperature and high pressure. Reexamining the MnTiO₃ data with the assumption, based our observations of the FeTiO₃-MgTiO₃ join, that the I to III transition must have a positive ΔS leads us to the conclusion that MnTiO₃ III is also the stable high-temperature and high-pressure phase in this system. Our new understanding of the energetics of I-II-III transitions in FeTiO₃-MgTiO₃ may also allow predictions to be made for other titanates.

Figure 8 shows the relationship between “A” cation size and the energetic barrier to forming a perovskite. The “A” cation radius of Ca is 1.00 Å, and CaTiO₃ is stable as a perovskite at room temperature and pressure. BaTiO₃ has “A” cation radii greater than 1 Å, is stable at room temperature and pressure as a perovskite, and has a high temperature hexagonal polymorph (not an ilmenite structure). The $\Delta H_{298}^0(\text{hex} \rightarrow \text{pv})$ of this transition does not fall on the line for ilmenite-perovskite transitions.

The ΔH^0 of transition between II and I increases linearly with Mg content (Fig. 9). Similarly, $\Delta H_{\text{I} \rightarrow \text{II}}^0$ can be shown to be a linear function of “A” cation radius (in sixfold coordination). This relationship holds even when other titanates, such as MnTiO₃ (Ko et al. 1989), are included. The titanate “A” cation radius also depends linearly on the 0 K pressure of transition from ilmenite to perovskite (Fig. 10), except that CdTiO₃, with an “A” cation radius of 0.95 Å, falls off the trend at negative ΔH^0 . CdTiO₃ is the only one of the four titanates

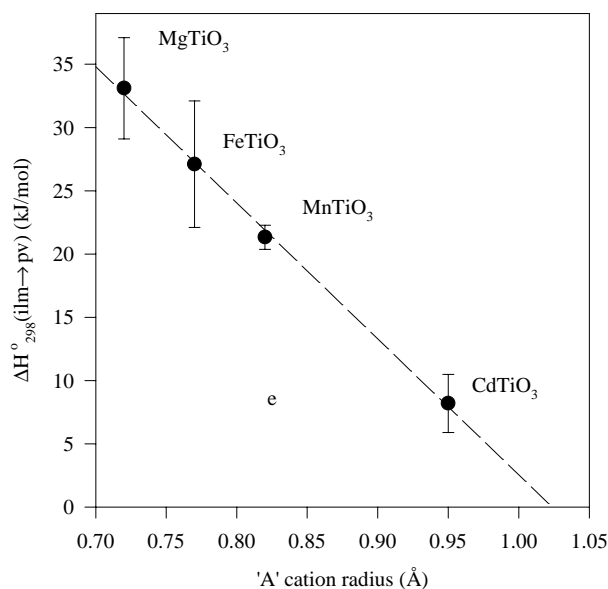


FIGURE 8. $\Delta H_{98}^0(\text{ilm} \rightarrow \text{pv})$ vs. the “A” cation ionic radius in sixfold coordination of the titanates from this study, plus CdTiO₃ (Takayama-Muromachi and Navrotsky 1988) and MnTiO₃ (Ko et al. 1989). The line is a best fit with a $\Delta H_{98}^0 = 0$ intercept of 1.02 ± 0.04 Å.

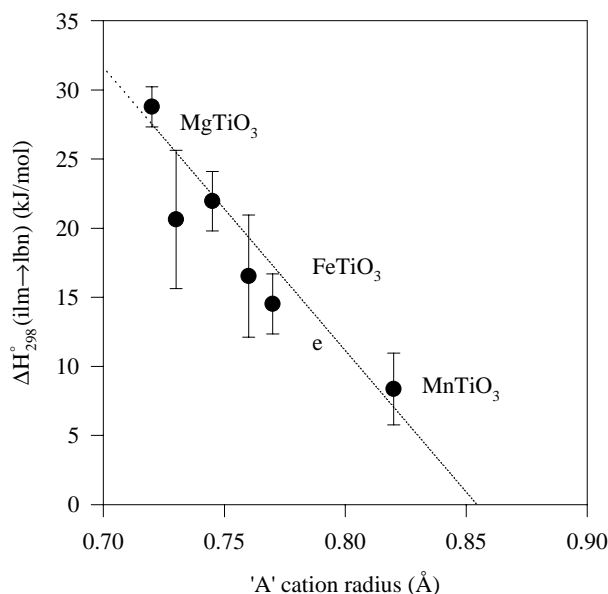


FIGURE 9. $\Delta H_{298}^{\circ}(\text{ilm} \rightarrow \text{lbn})$ vs. the “A” cation ionic radius in sixfold coordination of the titanates from this study, plus MnTiO₃ (Ko et al. 1989). Line is best fit line with a $\Delta H_{298}^{\circ} = 0$ intercept of 0.85 ± 0.05 Å.

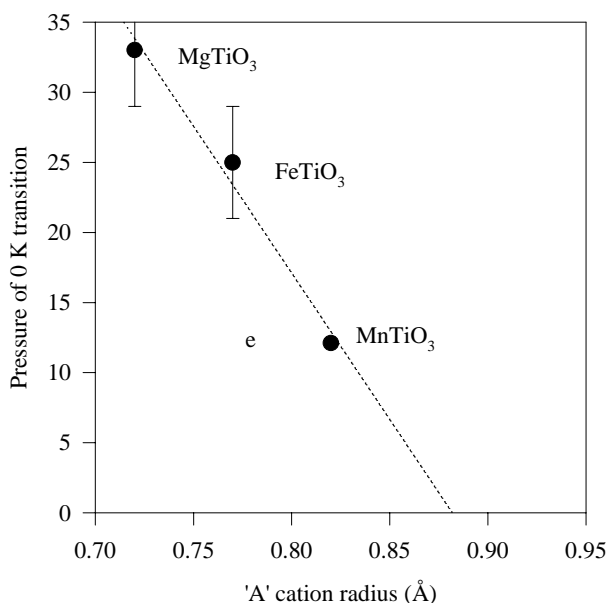


FIGURE 10. The “A” cation ionic radius in sixfold coordination vs. the 0 K pressure of transition from ilmenite to perovskite from phase equilibria studies: MgTiO₃ from this study, FeTiO₃ from Syono et al. (1980), MnTiO₃ from Ko et al. (1989). The dotted line is a best fit with a $P = 0$ intercept at 0.88 ± 0.04 Å.

(CdTiO₃, FeTiO₃, MgTiO₃, and MnTiO₃) which can be synthesized at atmospheric pressure in the perovskite structure, and CdTiO₃, unlike FeTiO₃, MgTiO₃, and MnTiO₃, does not have a lithium niobate quench phase.

The similar intercepts for the $\Delta H_{298}^{\circ}(\text{ilm} \rightarrow \text{lb})$ and 0 K pressure of transition plots (Figs. 8–10) suggest that for titanates

the upper limit of “A” cation radius for compounds with a lithium niobate quench phase is $\sim 0.86 \pm 0.15$. We can predict on this basis that NiTiO₃ (Ni ionic radius = 0.70) and CoTiO₃ (Co ionic radius = 0.735) will have high-pressure perovskite polymorphs that will quench to lithium niobate structures on release of pressure. This correlation, if it holds for minerals other than titanates, could be used as a useful indicator of the stability of lithium niobate type minerals. This result also indicates that Fe is not acting differently from Mg due to its electronic structure, as had previously been suggested by Reynard and Guyot (1994); the size of the “A” cation governs stability.

Different A²⁺B³⁺O₃ oxides with perovskite structure types are compared using a Goldschmidt diagram, Figure 11 (Ringwood 1970). Because this type of diagram is typically used to designate fields for structures such as garnet, pyroxene, and ilmenite where the “A” and “B” cations are in sixfold coordination, the “A” and “B” cation ionic radii (r_A and r_B) plotted are with a sixfold coordination with values taken from Shannon and Prewitt (1969). The tolerance factors:

$$t = (r_A + r_O) / [2(r_B + r_O)]^{1/2}$$

shown are a measure of the size mismatch for sixfold coordination. Tolerance factors are approximately five (more or less, depending on the A cation) less than the Goldschmidt tolerance factors for perovskites, which are calculated with the “A”

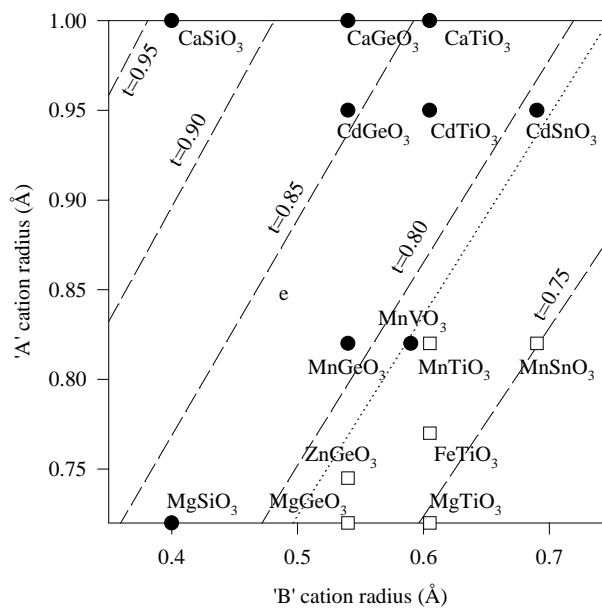


FIGURE 11. Goldschmidt diagram, based on Leinenweber et al. (1991), plots the “B” cation radius in sixfold coordination vs. the “A” cation radius also in sixfold coordination of selected ABO₃ compounds. Filled circles are perovskite forming compounds with no lithium niobate quench phase. Open squares have lithium niobate quench structure. Dashed lines are tolerance factors. Dotted line is our calculated upper tolerance limit for compounds to have lithium niobate quench structure (see text).

cation in eightfold coordination, as it would be in perovskite. The ideal cubic perovskite structure would have a Goldschmidt tolerance factor equal to one. However, for ilmenites and lithium niobates, where the "A" cation is in sixfold coordination, plotting a tolerance factor based on the "A" cation radii from sixfold coordination is more convenient. The upper limit for "A" cation radius for a titanate which will form a lithium niobate quench product corresponds to a tolerance factor of 0.79 ± 0.01 Å (with "A" in a sixfold coordination). This value agrees well with the behavior of other ABO₃ compounds plotted (Fig. 11).

These results can also be compared to the arguments in Funamori et al. (1997) for the existence of a LiNbO₃ quench phase of natural garnet with a composition between Mg₄Si₄O₁₂ and Mg₃Al₂Si₃O₁₂. Funamori et al. (1997) calculate tolerance factors for Mg₄Si₄O₁₂ and Mg₃Al₂Si₃O₁₂ using "A" cations in sixfold coordination. They claim the *t* values they calculate, 0.83 and 0.80 respectively, are below the expected *t* values, *t* > 0.85, for samples quenching to rhombohedral perovskites, but they are just above our limit for lithium niobate quench products. A sample with a composition between Mg₄Si₄O₁₂ and Mg₃Al₂Si₃O₁₂ might have too high a tolerance to yield a lithium niobate quench product; however, as they predict, with increasing Al content the likelihood of quenching to lithium niobate would increase.

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