

1 **REVIEW 2**

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3 **Radiation damage in biotite mica by accelerated α -particles: a synchrotron microfocus**
4 **X-ray diffraction and X-ray absorption spectroscopy study**

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20 synchrotron, microfocus, X-ray diffraction, X-ray absorption, EXAFS, infrared,
21 spectroscopy, geodisposal, nuclear waste

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ABSTRACT

32 A critical radiation damage assessment of the materials that will be present in a Geological Disposal
33 Facility (GDF) for radioactive waste is a priority for building a safety case. Detailed analysis of the
34 effects of high energy α -particle damage in phyllosilicates such as mica is a necessity, as these are
35 model structures for both the clay-based backfill material and the highly sorbent components of a
36 crystalline host rock. The α -radiation stability of biotite mica (general formula:
37 $\text{K}(\text{Mg,Fe})_3(\text{Al,Si}_3\text{O}_{10})(\text{F,OH})_2$) has been investigated using the 5 MV tandem pelletron at the
38 University of Manchester's Dalton Cumbrian Facility (DCF) and both the microfocus spectroscopy
39 (I18) and core X-ray absorption spectroscopy (B18) beamlines at Diamond Light Source (UK).
40 Microfocus X-ray diffraction 'mapping' has demonstrated extensive structural aberrations in the mica
41 resulting from controlled exposure to the focused $^4\text{He}^{2+}$ ion (α -particle) beam. Delivered doses were
42 comparable to α -particle fluences expected in the highly active, near-field of a GDF. At doses up to
43 6.77 displacements per atom (dpa) in the region of highest particle fluence, biotite mica displays a
44 heterogeneous structural response to irradiation on a micron scale, with sequential dilation and
45 contraction of regions of the structure perpendicular to the sheets, as well as a general overall
46 contraction of the phyllosilicate layer spacing. At the peak of ion fluence, the structure collapses
47 under a high point defect density and amorphous areas are pervasive amongst altered domains of the
48 original lattice. Such structural alterations are likely to affect the material's capacity to sorb and retain
49 escaped radionuclides over long time scales; increased edge site availability may favour increased
50 sorption whilst interlayer uptake will likely be reduced due to collapse. Radiation-induced reduction
51 of structural iron at the region of highest structural damage across an α -particle's track has been
52 demonstrated by Fe *K*-edge X-ray absorption near edge spectroscopy (XANES) and local structural
53 disorder has been confirmed by analysis of both potassium *K*-edge XANES and Fe *K*-edge extended
54 X-ray absorption fine structure analysis. An infra-red absorption study of deformations in the OH⁻
55 stretching region, along with electronprobe microanalysis complements the synchrotron data
56 presented here.

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INTRODUCTION

61 Phyllosilicates, such as mica and clay mineral phases, are of critical importance to the international
62 safety case behind the construction of a long-term geological disposal facility (GDF) for radioactive
63 waste. Biotite is a ubiquitous component of the crystalline host rocks being considered for permanent
64 geological isolation and phyllosilicate structures are expected to retard radionuclide transport
65 following eventual waste canister failure, accumulating radiation damage in the process. In addition,
66 mica shares similar structural and chemical characteristics with components of the proposed bentonite
67 backfill (i.e. hydrated, metal rich layer-silicates with variable interlayer components) and an
68 investigation into the α -radiation stability of biotite will have important parallels for clay-based
69 barrier performance assessment in a repository (Lee and Tank, 1985; Allard and Calas, 2009; Savage
70 and Arthur, 2012). An examination of the effects of ionising radiation on these mineral structures and
71 the potential impact upon their durability as barrier materials is essential.

72

73 The microscale, naturally occurring radiation damage accumulation in biotite mica by α -particles has
74 been studied before. Commonly exhibiting darkened aureoles of α -radiation damage surrounding
75 micrometer-scale radioactive inclusions, biotite is a useful proxy for studying the complex,
76 heterogeneous nature of radiation damage in layered minerals over long timescales (Nasdala et al.,
77 2001, 2006; Patrick et al., 2013; Bower et al., 2015b). This study presents the results of controlled α -
78 radiation damage in biotite mica using accelerated (5 MeV) $^4\text{He}^{2+}$ ions to systematically examine the
79 effects of α -particle bombardment on the mineral structure and chemistry. With control over the ion
80 energy, flux and sample irradiation conditions, elucidating the mechanisms of radiation damage is
81 possible, in contrast to naturally radiation damaged samples which may have unconstrained thermal
82 histories. α -particles are identical to $^4\text{He}^{2+}$ ions, however $^4\text{He}^{2+}$ ions are conventionally termed ‘alpha’
83 particles only when they are a product of nuclear decay (Lieser, 1997). For the purposes of this study,
84 both α -particles and $^4\text{He}^{2+}$ ions are considered synonymous.

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METHODS

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Biotite Mica

90 The biotite samples irradiated in this study were extracted from a large (10 x 15 x 0.5 cm) single
91 crystal from Iveland, Southern Norway (Rickwood, 1981), displaying chemical homogeneity across
92 broad (cm scale) areas. Electron probe microanalysis (EPMA) yielded an average formula of
93 $K_{0.83}Na_{0.016}(Mg+Fe_{1.21})Ti_{0.17}Al_{1.13}Si_{2.60}O_{10}(F,OH)_2$ where $Mg/(Mg+Fe)=0.545$, siting the biotite near
94 the centre of the phlogopite-annite solid solution series. For the purposes of this study, the term
95 ‘biotite’ is employed to describe the sample, albeit it is a general series name for trioctahedral dark
96 mica (Rieder et al., 1999).

97

98 The biotite mica structure comprises sheets of metal rich octahedra (O), bounded by silica and
99 alumina rich sheets of tetrahedra (T) arranged in hexagonal rings; the apical oxygens of the tetrahedral
100 layers point inwards towards the octahedral layer. Biotite is a trioctahedral mica, indicating that all
101 three octahedral sites are occupied, however vacant sites are typical across the structure (Fleet et al.,
102 2003). Negatively charged TOT sheets are weakly bonded by interlayer potassium cations, a minor
103 lateral offset in sheet stacking results in a monoclinic crystal habit. Structural hydroxyl groups lie in
104 the centre of the hexagonal tetrahedral rings (see Figure 1). Such 2:1 phyllosilicates are typified by
105 their perfect (001) cleavage, easily splitting along interlayers.

106

Ion irradiation

107 Samples were irradiated using the University of Manchester’s newly commissioned 5 MV NEC
108 15SDH-4 tandem pelletron ion accelerator, equipped with a toroidal volume ion source (TORVIS) at
109 the Dalton Cumbrian Facility, UK (Leay et al., 2014). Samples were held in a customized target
110 station (Bower et al., 2015) and irradiated under vacuum with a focused $^4He^{2+}$ beam. The energy of
111 the ion beam on the sample was held at 5 MeV across all exposures, in line with α -particle energies
112 along the uranium decay chains (48 MeV) (see NNDC, 2013). Activation cross-sections were
113 calculated for incident helium ions accelerated into constituent elements of the biotite target chemistry

114 to ensure any activation of the mica would be kept as low as reasonably possible. An incident energy
115 of 5 MeV was therefore selected to minimise sample activation. Ion fluences were varied with
116 irradiation time and beam current, which was kept sufficiently low (< 300 nA) to minimise the
117 temperature increase on samples during exposure. Biotite samples were prepared as thin (< 40 μm)
118 cleavage parallel planes and mounted on glass slides for irradiation. As such, samples were irradiated
119 along a plane normal to (001), perpendicular to the phyllosilicate layers (see Figure 1). Post-exposure,
120 samples were further thinned (by peeling along the {001} cleavage) from the back to remove the
121 undamaged material and isolate the high defect density layer for analysis.

122 *Ion dose and radiation damage modeling*

123 Radiation damage in materials is expressed in units of displacements per atom (dpa) i.e. a dpa value of
124 0.1 signifies that 10% of the atoms in the structure have been displaced from their original sites at
125 least once. The Monte-Carlo based simulation software, the Stopping and Range of Ions in Matter
126 (SRIM) (Ziegler, 2013) was used to calculate necessary ion fluences to achieve sequentially
127 increasing ion doses across the biotite samples, up to a maximum ion density normalised dose of 6.77
128 dpa (average across dose gradient = 0.28 dpa; see Tables 1 & 2). The critical amorphisation threshold
129 for a mica structure has been calculated to be 0.15 dpa (Wang et al. 1998), however the X-ray
130 diffraction (XRD) study here demonstrates that some original structure is retained at doses far higher
131 than this, despite extensive point defect densities throughout the former lattice. SRIM modeling was
132 used to determine that a relatively light incident ion such as helium entering a mica sample at 5 MeV
133 creates on average 225 knock-on displacements (Frenkel defects) per ion track in the sample. Across
134 such a track, the interaction of the ions with the biotite displays a peak with respect to atomic
135 displacements at ca. 18 μm from sample surface. Energy is lost via ionisation through the region of
136 low nuclear interaction i.e. the first ~ 15 μm as the α -particle accepts electrons to render it a neutral
137 helium atom. For comparison, SRIM modeling using 8 MeV helium ions indicated an average of 249
138 'knock-on' displacements over a 36 μm ion track, thus an increase in ion energy does not denote a
139 proportional increase in radiation damage, as structural interaction is theoretically maximised only
140 after sufficient energy is lost to ionisation (helium ion energy ~ 100 keV or less).

141

142 Three biotite samples exposed to sequentially increasing ion doses have been analysed in this study
143 (Table 1). Across all samples, >97 % of the beam fluence was contained within 1cm^2 .

144

145 Beam intensity using the DCF Pelletron's focused ion beam displays a Gaussian profile in real space
146 with respect to particle fluence. In this manner, one sample can accumulate a range of doses along a
147 beam radius of ~ 6.75 mm and thereby allow analysis of relative changes in damage accumulation
148 across a single sample (see Figure 2). The ion fluence gradient across the beam spot for Sample 1,
149 and the corresponding estimates of dpa values based on the integrated ion beam densities at intervals
150 across the profile are given in Table 2.

151

152 Placed in broad context for a deep geological disposal facility, Reed et al., (1987) calculate that within
153 a high level waste repository, the total absorbed α -dose at the surface of a waste package will be \sim
154 1.4×10^8 gray. This model assumes a 10,000 year integration time and canister failure after 1000 years.
155 In comparison, the total α -dose across sample 1 (Table 1), surface-area normalized for Reed et al.,'s
156 calculations, equates to $\sim 1.5 \times 10^6$ gray; samples 2 and 3 received 0.9×10^6 and 0.3×10^9 gray
157 respectively. These doses are a large over estimate for that predicted to be experienced by the far-field
158 host rock, however the parallels between phyllosilicates and bentonite backfill constituent minerals
159 make this dose a reasonable proxy for near-canister materials, although not as high as the modeled
160 surficial dose under the assumptions presented by the 1987 study. Further work at higher doses is
161 currently underway. The concentration of radionuclides in a host rock following release must also be
162 considered (through fluid flow and sorption onto host rock constituents/precipitation of radionuclide
163 bearing phases) and may cause variations in expected doses in the far field. A 10,000 year model is
164 rather short; it is likely than in the near field, the host rock will be subject to radiation fields up to
165 100,000 years and radionuclides will be concentrated in specific locations, such as phyllosilicates in
166 permeable zones.

167 *Synchrotron analysis*

168 Structural and chemical analysis of the radiation damaged samples was undertaken on beamlines I18
169 and B18, Diamond Light Source, UK (Dent et al. 2009; Mosselmans et al. 2009).

170 *Microfocus X-ray diffraction*

171 Microfocus, high resolution X-ray diffraction (μ XRD) patterns were collected in transmission
172 imaging mode on beamline I18, with a beam energy of 12 keV and a spot size of $\sim 3 \mu\text{m}$ diameter.
173 The setup comprised a Photonic Science XDI-VHR 125 CCD ranging from $12\text{-}44^\circ 2\theta$ with image
174 acquisition times of ca. 10-30s, depending on sample thickness. Si powder was used to calibrate the
175 images. The software Igor Pro (Igor Pro, 2014) with the Nika plugin (Ilavsky, 2012) was used to
176 calibrate and reduce the diffraction data and MatLab (The MathWorks Inc. 2013) was used to compile
177 μ XRD ‘maps’.

178

179 Sequential μ XRD traverses across the $^4\text{He}^{2+}$ ion beam dose gradient have quantified changes in the
180 structural parameters of biotite mica with increasing radiation damage. The irradiated samples were
181 mounted relative to the X-ray beam in two orientations; one with a 10° offset from [001] (z-axis
182 parallel) and another with a similar offset from [100] (x-axis parallel) to yield information from both
183 normal to and parallel to the basal plane, respectively. In this way, structural analysis could be
184 performed both along the irradiation vector and also at 90° to this, the latter permitting analysis of the
185 depth of the ions’ penetration into the sample. A high-precision XY stage was used to trace the beam
186 across damaged regions; for the purposes of these experiments, a beam step size of between 5-30 μm
187 was used, depending on required resolution.

188

189 Cu $K\alpha$ (1.541 Å) diffraction peaks for the biotite crystal structure were modelled using the software
190 *Jems* (Stadelmann, 2012) and comparisons made with the synchrotron XRD peaks at 12 keV (1.033
191 Å). Where confident identification was possible, indexed peaks have been quoted within diffraction
192 patterns. Reduced diffraction patterns have been fitted with a combination of Gaussian and Voigt line
193 profiles to give more accurate peak positions in 2θ space.

194

195 *Microfocus X-ray absorption spectroscopy*

196 X-ray absorption spectroscopy (XAS) data were acquired at both the Fe and K *K*-edges across the
197 radiation damaged samples, yielding information on the oxidation state and short range structure of
198 the Fe-rich octahedral layers, as well as the local structure of the K-rich interlayers, respectively.
199 Microfocus, Fe *K*-edge XAS data were collected in fluorescence mode on beamline I18, along
200 traverses coincident with XRD data acquisition. The penetration depth of the X-rays at the Fe *K*-edge
201 is ~ 20 μm given the chemistry and density of the sample, this is broadly consistent with the modeled
202 penetration distance of the α -particles into biotite, thus a useful probe for the region of highest
203 structural damage across the α -particle's track. Given its lower edge energy, potassium data was
204 sampled from a shallower region (~ 10 μm from the surface). Potassium *K*-edge XANES data were
205 also collected in fluorescence mode in a helium atmosphere on beamline B18, the core extended X-
206 ray absorption fine structure (EXAFS) beamline at Diamond Light Source. In contrast to the
207 microfocus (~ 3 μm spot size) study, beamline B18 acquires an average signal from a broad area (~ 1
208 mm spot size) of sample. All XAS data were reduced and analysed with the *Demeter* suite (Ravel and
209 Newville, 2005). Spectra are presented as background subtracted, intensity normalized plots. In the
210 instance of the potassium data, multiple (~16 per sample) spectra have been merged to yield better
211 data quality.

212

213 *Fourier transform infra-red spectroscopy*

214 Fourier transform infra-red (FT-IR) spectra were collected using a SpotlightTM 400 FT-IR
215 spectrometer in attenuated total reflectance (ATR) mode. Single cleavage planes of biotite were
216 placed flat against the ATR crystal and scanned in high resolution across the characteristic absorption
217 window for the OH⁻ stretching region (3800-3300 cm^{-1}). Spectra were reduced and normalized in the
218 specialist FT-IR analysis software OMNIC (Thermo Electron Corporation, 2004).

219

220 *Electronprobe microanalysis*

221 Electronprobe microanalysis (EPMA) was undertaken using the Pacific Northwestern National
222 Laboratory's (PNNL) JEOL 8200. The instrument was calibrated against commercial P&H
223 Developments mineral standards: albite ($\text{NaAlSi}_3\text{O}_8$), apatite $\text{Ca}_5(\text{PO}_4)_3(\text{OH},\text{F},\text{Cl})$, hematite (Fe_2O_3),
224 periclase (MgO), orthoclase (KAlSi_3O_8), rutile (TiO_2) and wollastonite (CaSiO_3).

225

226 **RESULTS AND DISCUSSION**

227 *Microfocus XRD*

228 *Analysis perpendicular to basal plane (z-axis parallel)*

229 Two XRD traverses across sample 1 are presented in the following section (Figures 4-8). Initially, a
230 series of images were collected from the undamaged region into the zone of maximum ion dose with
231 an image acquisition step of 30 μm (233 patterns over 7000 μm) (Figures 3 & 4 i-iii). This was
232 complimented by a shorter, higher resolution traverse with a step size of 5 μm across a 300- μm region
233 of interest (60 patterns), selected to display a representative area of diffraction peak variability (Figure
234 4 iv). Both sets of diffraction patterns were collected along the Gaussian profile of the ion fluence
235 gradient shown in Figure 2. Representative CCD images from sequential positions across the
236 irradiated biotite are displayed in Figure 3.

237

238 At 7000 μm from the ion beam fluence maximum, the area analysed is outside the irradiated region
239 and sharp spots are visible in the diffraction pattern (Figure 3). Crossing over the ion beam limit,
240 minor reflection angle fluctuations and some diminished intensity occurs until ca. 4000 μm from the
241 peak of ion dose, however, the original structure is still evident and all reflections are represented by
242 sharp spots. At 4000 μm some diffraction spots have split and most high angle reflections have
243 become more diffuse. Peak splitting, line broadening and selective intensity reduction persist to 2000
244 μm , whereupon spots become smeared, inconsistent rings; indicative of multiple crystallites of this
245 pseudo-original structure in numerous orientations. Where the ion beam dose was highest (1000 μm

246 to 0 μm), nearly all reflections have disappeared and the resulting pattern is a series of highly diffuse,
247 minor reflections bearing only a slight resemblance to the original structure. The biotite is interpreted
248 to be extensively amorphous over the final $\sim 500 \mu\text{m}$ towards the beam centre, as broad, diffuse spots
249 become faint, minor reflections, signifying near complete loss of long range order.

250

251 Merging a series of 1D patterns (reduced via rotational averaging of the CCD images) yields the XRD
252 ‘map’ shown in Figure 4. The diffraction map has been separated into three regions of interest for
253 subsequent discussion. Panel i is the lowest dose region and displays the reflections outside the
254 irradiated area as well as the limit of the ion beam. Consequently, at this scale, reflections show both
255 unirradiated and irradiated regions and thus a high reflection angle variability. Panel ii is a region of
256 more intense dose across the Gaussian gradient whereby radiation effects are more pervasive. Panel iii
257 therefore represents the peak of ion dose across the biotite and is dominated by near-amorphous
258 regions.

259

260 Several assertions can be made from detailed analysis of these maps. Relative comparisons in XRD
261 patterns have been made across a single sample, owing to the difficulty in indexing patterns across
262 multiple single crystals using microfocus diffraction due to orientation and thickness variations. Panel
263 i (7000 μm – 4700 μm) shows a higher variation in 2θ values than any other point across the traverse
264 at this scale. Directly across the ion beam limit, most reflections shift to higher angles (lower d-
265 spacings) by up to 2 degrees 2θ and broaden, periodically fluctuating back towards their original
266 values. The line broadening observed is consistent with a contribution from point defect
267 accumulation, however there is no apparent peak splitting or loss of reflections in the initial low dose
268 area. It is likely that at these doses, radiation damage in the crystal is variable in intensity due to the
269 stochasticity of defect accumulation in the sample, thus lattice dimensions are extremely
270 heterogeneous on a micron scale. The map suggests areas of undamaged crystal are neighbouring
271 regions where bulk crystal parameters have shifted slightly (collapsed) to accommodate increased
272 point defects. Across the region of highest ion beam damage within panel i, most peaks have moved

273 to higher 2θ values and these peak positions remain relatively constant, indicating increasing defect
274 concentrations throughout the crystal, yielding a more homogeneous bulk structural response to
275 irradiation.

276

277 Panel ii (4700 μm – 2200 μm from peak of dose) displays a new reflection (also visible in the top of
278 panel i, ca. 5200 μm from the peak of ion dose), ‘growing in’ at ca. 24.9 degrees 2θ . This ‘new’ peak,
279 indexed as the biotite (158) reflection, remains the most intense reflection across the traverse as the
280 dose the biotite experienced increases. The emergence of this peak is consistent with the introduction
281 of rings in the diffraction patterns. It is proposed that a further increase in point defect density results
282 in the reorientation of crystallites within the structure, revealing new reflections in the view direction.

283

284 Of particular interest is the high resolution trace (panel iv) (~4300 μm – 4000 μm from the peak of
285 ion dose) across panel ii, displaying significant deviations in diffraction angle with smaller peaks
286 ‘splitting’ from the main reflection. These ‘satellite’ reflections display a sinusoidal movement in 2θ
287 space, arcing away from and then back towards the intense, primary reflection. Peak splitting is
288 indicative of a non-uniform structural response to higher defect accumulation and suggests the
289 presence of multiple crystallites with varying lattice parameters. Certain sinusoids appear
290 concomitant, moving in phase with a neighbouring reflection, whilst others have a shallower deviation
291 from the main peak and appear out of phase as regions of the crystal dilate and contract sequentially.
292 It is clear that radiation damage and structural accommodation have a consequential effect across
293 neighbouring areas in the structure.

294

295 In panel iii (2200 μm - 0 μm from peak of dose), reflection intensity across all peaks has diminished
296 and closely spaced peaks have merged and cannot be isolated for indexing. The peak at $\sim 27^\circ$ displays
297 a large progressive increase in 2θ of up to 5° with increasing ion beam intensity, before losing
298 intensity and becoming indistinguishable from background. Structural collapse is prevalent across the

299 region of highest dose and all reflections are sporadic and highly diffuse by the centre of the ion beam
300 spot.

301

302 Figure 5 displays the reduced 1D diffraction patterns at intervals across the traverse with peaks
303 indexed. In the context of the biotite crystal structure, reflection (131) (at ca. 22.8 degrees 2θ) denotes
304 a plane of metal cations in alternate octahedral sheets *across* phyllosilicate layers. With increasing ion
305 fluence, the (131) peak splits and both peaks display a shift to higher reflection angles, signifying a
306 contraction of 'domains' of the lattice along this plane. At the peak of radiation damage, this
307 reflection is lost, likely due to a variation in phyllosilicate layer offset such that the cations are no
308 longer in-plane.

309

310 A contractive shift is also observed for reflection (204) (at ca. 26.9 degrees 2θ), a plane of octahedral
311 metal cations in a similar orientation to the (131) plane, however at a more oblique angle to the basal
312 plane. The (204) reflection splits and displays multiple satellite peaks that vary sinusoidally in real
313 space, suggesting that structural change perpendicular to the phyllosilicate layers is non-uniform;
314 indeed contractions along certain planes of metal cations results in corresponding expansion
315 elsewhere in the crystal, as parts of the crystal bunch and relax to accommodate high defect densities.

316

317 Reflection (-225) (at ca. 28.8 degrees 2θ) also shifts to higher 2θ angles and represents a decrease in
318 d -spacing along this plane; an oblique series of basal tetrahedral oxygens. Contraction in this plane
319 will contribute to interlayer variations in biotite. Whilst not a contraction parallel to phyllosilicate
320 sheets, the misalignment of basal tetrahedral atoms will cause highly variable interlayer spacing as
321 domains within sheets become buckled.

322

323 Reflection (02,12) (at ca. 38.0 degrees 2θ) comprises planes of interlayer potassium atoms across
324 sheets. It is clear that the interlayers are the planes most susceptible to radiation damage, as the

325 reflection loses nearly all intensity almost immediately across the ion beam limit. The initial loss of
326 interlayer K^+ is in good agreement with previous studies of ion beam damaged powdered mica
327 (Chailley et al., 1994).

328

329 The new reflection, indexed to the (158) plane in biotite, only appears with increasing $^4\text{He}^{2+}$ ion
330 fluence (for distances $< 5300 \mu\text{m}$ from the peak of ion dose). Reflection (158) represents a series of
331 apical tetrahedral oxygen atoms in-plane across TOT layers. This region of the biotite unit cell has the
332 highest atomic density; apical oxygens of silica tetrahedra are shared with octahedrally coordinated
333 metal cations at the centre of the TOT sheets. Although the emergence of new reflections generally
334 represents new phases or increased crystallization, in this case it is unlikely to be the case here. Whilst, most
335 structural damage is manifested as interlayer collapse or structural offset, it is plausible to assume that
336 metal-rich octahedral layers are the most stable element of the mica structure and that changes in layer
337 offset will preferentially align regions with a higher percentage of crystallinity. As damage increases,
338 crystallites of pseudo-original structure will distort to accommodate lattice strain changes; such
339 distortions will also reveal new reflections, similar to tilting a sample in a transmission electron
340 microscope (TEM). At the height of radiation damage, even reflection (158) has broadened and
341 greatly diminished due to high levels of accumulated point defects. Additionally, only very minor,
342 diffuse reflections (± 2 degrees from original positions) are apparent, suggesting that small domains
343 of pseudo-original structure remain.

344

345 Relative changes in bulk lattice parameters were calculated using the X-ray diffraction analysis
346 software *UnitCell* (Holland and Redfern, 1997), results are displayed in Figure 6. The data is refined
347 from a series of peak-fitted patterns across points within the traverse presented in Figure 4. Only
348 dominant, indexable reflections, rather than diffuse satellite peaks, were used in the refinement, thus
349 small-scale changes in discrete crystallites are not represented. The biotite *a*- vector distance i.e. the
350 shorter basal edge of the monoclinic prism (coincident with the {010} plane) only fluctuates by up to
351 0.1 \AA , thereby remaining relatively stable across the irradiated surface, only decreasing at the height of

352 ion fluence. This is in contrast to the b - vector length, which remains consistent before the ion beam
353 limit and then becomes extremely variable with increasing radiation dose. The b - vector represents the
354 long basal edge of the monoclinic prism. This variability ultimately results in a sharp decrease in unit
355 cell volume with increasing radiation dose, as the structure collapses due to higher defect densities.
356 Due to the orientation of the crystal upon analysis (ca. 10 degrees offset from $\{001\}$), accurate c -
357 vector values were unattainable over the z -axis *parallel* XRD traverses, as no reflections yielded
358 accurate structural information parallel to layers. Further work upon characterising interlayer
359 behaviour is presented in the next section.

360

361 *Analysis parallel to basal plane (x-axis parallel)*

362 In comparison to the XRD analysis perpendicular to the basal plane, spectra collected parallel to the
363 basal plane display far more reflections. Inconsistencies in interlayer arrangement mean that not all of
364 these reflections have been accurately indexed, however a general trend can be observed across the
365 series of representative diffraction patterns shown in Figure 7. Reflections at 0 μm denote the surface
366 of the irradiated biotite; patterns were collected in 5 μm steps along the trajectory of the incident α -
367 particle to 30 μm , a depth which SRIM modeling predicts will be outside of the penetration depth of
368 the 5 MeV α -particles. The highest variation in peak position occurs within 20 μm from the exposed
369 biotite surface, although some loss in peak intensity is observed 25 μm from the surface. There is no
370 clear indication of a zone of ‘peak’ radiation damage at the end of the α -particles’ projected range (15
371 μm – 18 μm), however the data across this region is limited and peak broadening is pervasive across
372 the traverse. The (006) reflection (at ca. 17.9 degrees 2θ), is at its most intense at the irradiated
373 surface; this plane is one of the few elements of the structure that remains. It seems plausible to
374 assume some level of ‘packing’ of the interlayers as contractive effects become pervasive, potentially
375 retaining a certain degree of along layer crystallinity. Notably, peak (006) displays a shift to higher
376 reflection angles, consistent with interlayer collapse.

377

378 The intensity of most reflections is diminished by the irradiated surface of the crystal and most only
379 display an intense peak beyond 25 μm . Reflection (114) represents a series of in plane octahedral
380 cations across the sheets; applying these trends suggests severe buckling along phyllosilicate sheets,
381 throwing the metal cations out of plane.

382

383 Bulk X-ray diffraction across the cleavage planes was also performed on samples 2 & 3 using a
384 Bruker D8 lab source, yielding an average interlayer signal over the entire irradiated area for
385 comparison with unirradiated material (Figure 8). In broad agreement with the basal plane parallel
386 analysis, the bulk analysis shows a general trend of interlayer contraction (shift to higher 2θ angles)
387 across all layer-parallel reflections with increasing α -particle doses. Sample 3 (0.06 dpa) displays two
388 peaks either side of the original reflection angle, indicative of the development of two main structural
389 domains in the damaged lattice, as regions of the crystal dilate to accommodate collapse elsewhere.
390 This trend is not continued at higher doses (sample 2, 0.18 dpa), whereupon all plane-parallel lattice
391 planes present a contraction.

392

393 *X-ray absorption spectroscopy*

394 *Fe K-edge XANES*

395 Data collected up to 50 eV beyond the absorption edge (XANES region) yields information on the
396 core electronic state of the absorbing atom. Figure 9 displays the Fe *K*-edges collected in both an
397 unirradiated area (ca. 7000 μm from the ion beam centre) and an area of biotite in sample 1,
398 calculated to have accumulated ~ 0.28 dpa.

399

400 The reduction of Fe(III) adds another valence electron that partially reduces the interaction between
401 the Fe 1s electrons and the nucleus, leading to a reduced 1s binding energy and a shift of the white
402 line to lower energies, evident in the spectrum for irradiated biotite. Clear changes in the modulations
403 of the absorption coefficient beyond the edge are likely a product of the altered lattice dimensions

404 detailed earlier. Linear Combination Fitting (LCF) was possible for the two spectra presented in
405 Figure 9, using Fe(II) and Fe(III) Fe K-edge XANES standards wüstite (FeO) and hematite (Fe₂O₃)
406 respectively. LCF results indicate a 17.3 % increase in Fe(II) at the highest region of damage, at the
407 depth probed by the X-rays (see supplementary information).

408

409 *Fe K-edge EXAFS*

410 At energies above the absorption edge, the ‘extended’ region of the spectrum gives an insight into the
411 local environment of the absorbing atom. EXAFS data has been fitted to $k=11$, yielding structural
412 information to a third coordination shell in both the unirradiated and irradiated samples, aiding
413 quantification of radiation damage induced changes. Table 3 shows the results of the EXAFS fitting
414 across a range of ⁴He²⁺ ion doses; corresponding spectra are shown in Figure 10. Biotite mica, as a
415 trioctahedral phyllosilicate, should have divalent ions occupying all available octahedral sites. XAS
416 has shown that a proportion of iron is Fe(III) and that there are, therefore, local domains of
417 dioctahedral sites within the trioctahedral layers. Vedder (1969) notes that iron rich biotites exhibit a
418 high variability in the occupancy of octahedral layers.

419

420 Three paths have been fitted to the EXAFS spectra collected across the sample, (Fe – O), (Fe – Fe) &
421 (Fe – Si); each path assigned a fixed coordination number with independently variable Debye-Waller
422 factors, denoting disorder in the system. The amplitude reduction factors decrease into the irradiated
423 region, which may suggest the fixed coordination numbers are not wholly representative of the
424 damaged structures. Indeed, spectra C, D and E yielded better fits with reduced coordination in the
425 second and third shells, likely a product of defects in the structure. In general there are no significant
426 changes across the spectra, demonstrating high levels of local order at the scale analysed by EXAFS.

427

428 The irradiated spectra display a slight increase in average Fe - O distance, broadly increasing with
429 higher doses. A Fe – O bond length increase further supports Fe(III) to Fe(II) reduction. Consistent

430 with the XANES linear combination fitting (see supporting information), average bond length
431 increases across all spectra suggest up to a 20 % increase in Fe(II) across the irradiated region (at the
432 depth probed by the X-rays).

433

434 Fe-Fe distances (equating to the distance between neighbouring octahedra in the same sheet) and
435 coordination numbers are much more variable across the irradiated spectra. EXAFS analysis of the
436 unirradiated biotite yields a Fe-Fe distance of $3.11 \pm 0.03 \text{ \AA}$, comparing well to previous data for
437 biotite (Palmer and Conley, 1994). In contrast to the overall increase in Fe-O distance with
438 irradiation, Fe-Fe distances in irradiated biotite show either a slight increase or a more substantial
439 decrease in interatomic distance, to as low as $3.03 \pm 0.03 \text{ \AA}$. This is explained by variable, local
440 expansion and contraction within the metal layers demonstrated by XRD. Debye-Waller factors
441 across the Fe-Fe scattering paths are far higher than across other paths; this may be a product of
442 inherent 'disorder' in the octahedral layer, rather than a damage effect, as both trioctahedral and
443 dioctahedral coordination is present in the mica structure. This may also explain why the Fe-Fe
444 coordination number in the unirradiated sample is lower (4.7 ± 1.3) than the expected trioctahedral
445 CN of 6, although the unirradiated values are within refinement error (Table 3).

446

447 Fitting out to the third coordination shell, Fe-Si distances (octahedrally coordinated cation to
448 over/underlying tetrahedrally coordinated cation) show a general consistency with only small
449 variability ($\pm 0.02 \text{ \AA}$). This suggests a general stability of T-O sheets even at high doses, with a
450 slight contribution from within-layer collapse or buckling. The Fourier transform magnitude displays
451 a shoulder on the peak at $\sim 2.75 \text{ \AA}$, although it is not present across all spectra, possibly indicative of
452 inconsistent scattering contributions as the sheets are non-parallel. In the instances where the shoulder
453 is not present, the contractive shift of the second shell peak is greater. The heterogeneity of damage is
454 highlighted by the microfocus EXAFS at this resolution and it is difficult to interpret an overall trend.

455

456

K K-edge XANES

457 Potassium K-edge XANES data acquired on beamline B18 comprises a ‘bulk’ signal (beam spot = 3
458 mm diameter) and, therefore, provides only an average trend across a broad area of the radiation
459 damaged samples in comparison with the undamaged material. The location and coordination of
460 potassium in biotite presents challenges with XAS resolution. As a result, only XANES data are
461 presented as the data produced has an insufficient k range of high enough quality to yield reasonable
462 EXAFS fits. Figure 11 displays the potassium K-edge XANES and inset Fourier transforms for both
463 unirradiated and irradiated (average over beam area ~0.28 dpa) biotite. Small changes in the local
464 environment surrounding the potassium are evident in the irradiated sample in comparison to the
465 unirradiated standard, as shown by slight differences in the oscillations beyond the absorption edge. A
466 further qualitative assessment of this change is displayed in the Fourier transforms. Heterogeneous
467 local order around the absorbing atom (i.e. inconsistencies in layer spacing even in unirradiated
468 samples), as well as the extremely large first shell distances (~3 Å to the basal oxygens of the
469 tetrahedral layer) yield a high background, consequently the first peak in the transform is assigned as
470 background. The shells at ~3.2 Å and ~4.5 Å (non phase-corrected) represent basal oxygens and
471 cations of the surrounding tetrahedral layers, respectively. Both shells display a slight shift towards
472 shorter radial distances, in keeping with the assumption that radiation damage causes interlayer
473 contraction. This trend has been described in previous K K-edge EXAFS studies of phyllosulfates
474 (Strobel et al. 1993).

475

Fourier transform infra-red spectroscopy

477 Characteristic absorption bands across the OH⁻ stretching region of the biotite mica infra-red (IR)
478 spectrum display significant changes with increasing ion fluence (Figure 12). Structural OH⁻ groups in
479 mica lie in the voids at the centre of the hexagonal tetrahedral rings that make up the sheets and are
480 oriented towards the octahedral layer (Fleet et al., 2003). Strong orientation effects are typical of this
481 region of the mica spectrum, therefore all samples were analysed with cleavage planes perpendicular
482 to the beam.

483

484 Importantly, the irradiated samples were analysed from the back, not from the irradiated face, as the
485 FT-IR in ATR mode will only probe the initial 0.5-5 μm . By analysing the samples in this orientation,
486 the modeled region of highest damage is investigated (the end of the α -particles range), broadly
487 consistent with the region probed by synchrotron XAS. Two irradiated samples were analysed with
488 increasing dose, as well as a sample heated in air to 110°C; the calculated maximum temperature
489 reached over a 6 hour irradiation, to check for any thermally induced effects. Fe^{2+} oxidation through
490 heating is a well-documented mechanism in biotite (Rancourt et al., 2001 and references therein). As
491 no such trends were observed here, it is assumed that a sample temperature increase in-vacuum during
492 irradiation had a negligible effect on OH^- dissociation or redox reactions.

493

494 The unirradiated biotite spectrum displays a convincing fit to model IR spectra for ‘dark’ mica
495 presented by (Vedder, 1964; Wilkins, 1967; Chaussidon, 1971). A broad, intense peak at 3683cm^{-1}
496 can be attributed to hydroxyl groups neighbouring three divalent metal ions (N -group bands). This
497 region correlates to the N_C absorption band and denotes OH^- stretching neighbouring a trioctahedral
498 ($\text{Mg}^{2+} + \text{Fe}^{2+} + \text{Fe}^{2+}$) coordination (Wilkins, 1967). The adjacent, lower frequency bands (two poorly
499 defined peaks at 3585cm^{-1} and 3560cm^{-1}) denote a proposed vibration of hydroxyls adjacent to vacant
500 sites (V -group bands). The lower frequency peak indexes well to the V_C band; consistent with ($\text{Fe}^{3+} +$
501 $\text{Fe}^{3+} + \text{Vacancy}$) (Chaussidon, 1971), whilst the higher frequency peak remains unidentified.

502

503 At high $^4\text{He}^{2+}$ ion doses, significant changes are apparent. The broad, intense N_C peak has split into
504 two maxima at 3698cm^{-1} and 3664cm^{-1} . In keeping with the Vedder (1964) and Wilkins (1967)
505 models, these peaks index to characteristic N_B and N_D bands in biotite mica. A N_B band with a peak at
506 3698cm^{-1} is attributable to OH^- stretching neighbouring an ($\text{Mg}^{2+} + \text{Mg}^{2+} + \text{Mg}^{2+}/\text{Fe}^{2+}$) coordination,
507 whilst a N_D band with a maximum at 3664cm^{-1} denotes a pure ($\text{Fe}^{2+} + \text{Fe}^{2+} + \text{Fe}^{2+}$) octahedral
508 environment. The V bands become one broad peak, shifting to higher wavenumbers with increased

509 $^4\text{He}^{2+}$ ion dose, centred on 3569cm^{-1} . Shifts towards higher wavenumbers of the V-group bands denote
510 an increased divalent component of the octahedral coordination i.e. ($\text{Fe}^{2+} + \text{Fe}^{3+} + \text{Vacancy}$).

511

512 Consistent with the XANES data, these FTIR results may be further evidence of Fe^{3+} reduction at the
513 region of highest damage in the structure. Changes in Fe oxidation state have previously shown to
514 yield shifts in OH⁻ stretching band frequency in clay minerals (Fialips et al. 2002). Coupled with this,
515 the increase in intensity of the V-group bands denotes a higher signal from vacancy coordination, as
516 the volume of point defects in the structure increases. It is worth noting that the observed contraction
517 in the phyllosilicate layers as demonstrated by XRD may have an impact upon the orientation of the
518 O-H bond, which will have a consequent impact upon absorption across the stretching region. In
519 dioctahedral micas, the OH bond is oriented sub-normal to the (100) plane, with an inclination
520 towards the vacant octahedral site; any increase in layer charge from the tetrahedral layer will force
521 the bond to angle more steeply towards the octahedral layer (Fleet et al., 2003). Shifted frequencies
522 across V-group bands may denote varied OH⁻ group orientation, a product of high vacancy
523 accumulation. OH⁻ groups can act as charge transport ‘bridges’ between *M* sites in the octahedral layer
524 in mica (Rosso and Ilton., 2005); reorientation and/or dissociation at high α -fluences may have
525 detrimental consequences for electron transfer within the Fe sublattice.

526

527 *Electronprobe microanalysis*

528 Sample 1 was analysed via a step traverse from the unirradiated biotite into the discoloured damage
529 ‘zone’, coincident with the X-ray diffraction traverses described in the manuscript. Table 3 shows the
530 EMPA data across the traverse recalculated on an (OH)-free basis to 22 oxygens per formula unit
531 (pfu). Some relatively minor chemical changes were apparent across the major elements: a slight
532 increase in point to point variability, particularly across the Si values was present in the irradiated
533 region (along with a notable decrease in K). Radiation damage is clearly destabilising the interlayer
534 cations and thus Na + K will be more susceptible to volatilisation. A loss or relocation of poorly-
535 bound Na + K may be induced by slight local heating during irradiation or indeed by EPMA analysis

536 (see Nielsen and Sigurdsson, 1981). However the scale of analysis of the EPMA analysis will likely
537 not reveal the highly localised chemical changes produced by α -particle bombardment. In this sample,
538 as in many other natural biotites, the interlayer and octahedral sites are not completely filled, and, as a
539 result, the excess cations that are generated by treating total Fe as FeO do not exceed the
540 stoichiometric limit of 16 pfu (Table 4), therefore it is not possible to calculate model $\text{Fe}^{2+} - \text{Fe}^{3+}$
541 ratios. However, a decrease in the sum of Na + K (Table 3) into the irradiated region (ideal cation
542 total = 2.0) may also suggest an increased Fe^{2+} component is required to retain charge balance,
543 consistent with the XAS data.

544

545 *Radiation damage mechanisms*

546 Whilst this is the first detailed microfocus X-ray diffraction analysis of controlled α -particle radiation
547 damage in biotite mica, some studies of swift heavy ion irradiations on similar mineral phases exist
548 for comparison. (Chailley et al. 1994) document X-ray diffraction analysis of high energy (850 MeV)
549 xenon and (170 MeV) oxygen irradiation of mica. Xenon irradiations yielded an overall unit cell
550 expansion with increased defect accumulation, manifested as XRD peaks shifting to lower 2θ angles,
551 whilst oxygen irradiations yielded only peak broadening and no overall lattice parameter change. Both
552 this study and later work (Chailley et al., 1996) conclude the presence of 3 phases; amorphous,
553 pseudo-crystalline and expanded phases across an irradiated mineral. Biro et al. (1997) also observed
554 this trend using 209 MeV krypton ions, asserting that the “neutralisation” of interlayer potassium ions
555 causes the negatively charged phyllosilicate layers to repel each other and expand. Both previous
556 studies conclude there is a threshold of point defect accumulation below which radiation damage in
557 mica does not affect the overall structure. These studies present the opposite structural effect to those
558 observed with α -particles; they maintain there is an expansion of the lattice along phyllosilicate layers
559 as a result of radiation damage. It is clear that damage manifestation with very high energy, heavy
560 ions is wholly different than the radiation damage presented in the present study, the difference being
561 that swift heavy ions induce multiple defect ‘cascades’, resulting in far higher point defect densities

562 along a single ion track. Comparing the differences in damage accumulation produced by heavier ions
563 in phyllosilicates will form the basis of a future study.

564

565 Lattice contractions as a product of radiation damage have also been documented. (Lu et al. 2012)
566 present a 1.42 % lattice contraction in cubic ZrN, manifested as a shift to higher 2θ values of XRD
567 peaks, as a result of irradiation by 350 keV oxygen atoms. The study suggests that lower energy ions
568 can cause unit cell contraction; the number of displacements per incident ion must be low enough
569 such that interstitials move more readily than vacancies in a damaged structure. The migration of
570 interstitials through a structure leaves voids (vacancy nucleation) into which the lattice collapses.
571 Increased vacancy density thereby causes a relaxation of a crystal's inherent strain vectors. In Lu et
572 al's example, interstitial migration is aided by the nano-scale crystal size.

573

574 Under high-fluence α -particle irradiation, both dilation and contraction are apparent *across the TOT*
575 *sheets*, with extreme variability in cell parameters on a micron scale. As point defect densities
576 increase (manifested by peak broadening), XRD analysis indicates that biotite mica forms randomly
577 oriented crystallites of pseudo-original structure (as observed by broad 'rings' in the diffraction
578 patterns). Such crystallites are subject to heavy distortion with increased dose, as the structure
579 responds to the changing strain field of the increasingly damaged crystal. The likely loss of interlayer
580 potassium from its interlayer site and coincident layer offset causes *basal plane parallel* sheet
581 buckling and eventual structural collapse, potentially aided by the 'interstitial migration' model
582 presented earlier (Lu et al., 2012). Of particular note is the potential for helium gas build-up in a
583 sample, as neutral helium atoms diffuse readily through a structure and become trapped by (often self-
584 induced) vacancies. The formation of helium 'bubbles' has been documented in plutonium metal
585 (Wolfer 2000; Martz and Schwartz, 2003) and is likely to be a contributing factor within these
586 samples, the presence of 'sinusoidal' movements in XRD patterns may denote periodic bubble
587 formation at or near the surface of the biotite and such gas build up may contribute to interlayer
588 contraction. Further work on bubble formation in these materials is underway.

589

590 Structural Fe(III) undergoes radiation-induced reduction at the region of highest structural radiation
591 damage along an α -particle track; potentially a result of OH⁻ group radiolysis liberating reducing
592 hydrogen (Patrick et al., 2013), as well as increased electron density at depth due to displacement
593 cascades and charge cycling by the penetrative alpha particles. Radical oxidizing species formed from
594 hydroxyl group dissociation will be free to migrate from the structure, resulting in net reduction. In
595 addition, the lighter oxygen atoms within this high-damage region will be more readily lost from the
596 former lattice (largely controlled by kinetics, the probability of displacement will follow: O > Mg ~
597 Al > Si > Fe).

598

599 Transient displaced oxygen species, potentially leaving the structure as a neutral atom or radical anion
600 following electron donation from the α -particles, will result in ‘under coordination’ of the iron (and
601 other cations) which will be stabilized as Fe(II). It is plausible to hypothesize net chemical reduction
602 within the discrete, highly damaged region of the α -particle’s track, as increasing electron density,
603 combined with a higher migration potential of the oxidizing species, will prevent redox balance.
604 Oxygen displacement and loss of local order surrounding cations has been shown to be commonplace
605 across irradiated materials (Noda et al., 1993; Morono et al., 2009; Walsh 2011; Enterkin et al., 2011;
606 Peugot et al. 2014a, 2014b). Whilst the Fe – O distances in the EXAFS data presented here display
607 only a small increase, the increase is not as large as that expected for octahedrally coordinated Fe
608 reduction. The decrease in Fe coordination by the loss of oxygen may account for this smaller than
609 expected bond lengthening.

610

611 Net reduction may not be the sole mechanism across the entire interaction range of the α -particle;
612 most of the analysis here has focused on the region of highest structural damage as an α -particle loses
613 its energy towards the end of its track. Indeed, the potential for oxidation effects to dominate in the
614 initial region of interaction, as the α -particle strips surface electrons to render it a neutral helium atom,

615 must also be considered. A number of mechanisms are potentially active in these damaged systems;
616 the influence of structural water is an area of required extensive further research. Alongside the
617 potential for radical production by OH⁻ radiolysis, it has been shown that the presence of water within
618 an irradiated system can inhibit structural recovery and further promote defect concentration
619 (Lockwood and Garofalini 2010).

620

621 Layer charge variations must also be considered, as it is highly unlikely that the radiation damaged
622 phases are charge balanced. Increased electron density within certain regions may yield an overall
623 higher negative charge of the TOT layers; spatial variations in layer charge may therefore also
624 contribute to interlayer buckling, particularly if interlayer K⁺ has been relocated.

625

626 *Implications*

627 With particular relevance to geological disposal, the radiation damage mechanisms presented here
628 expose both advantages and problems for phyllosilicate performance under relatively high α -particle
629 doses. It is apparent that mica interlayers are the most susceptible to radiation damage; this XRD
630 study demonstrate the instability of interlayers even at low doses and it is likely that interlayer K⁺ ions
631 are liberated or relocated as a result. Such interlayer distortion will have detrimental consequences for
632 uranium (and other radionuclides) uptake by sheet silicate minerals as contraction of the interlayer
633 spacing below 1 nm will likely retard uptake efficiency. For example, uranium (VI) has been
634 demonstrated to be efficiently taken up by biotite mica in the form of uranyl hydroxide (Idemitsu et
635 al., 1995) along basal sheets, crystal edges (Ames et al., 1983) as well as within the interlayers.
636 Phyllosilicate sheet spacing showed expansions of up to 47 % when uranyl ions were sited within
637 mica interlayers, aided by interlayer dilation from Ca²⁺ and Na⁺ substitution for K⁺ (Lee et al., 2009).
638 A consistent sheet spacing of > 1 nm is optimal for efficient uranium uptake within interlayers; such
639 consistency is not retained at high α -doses. This will also be detrimental for larger ions such as
640 caesium (although decay timescales must be considered in the context of escape to the geosphere). In
641 contrast, assuming a randomly oriented crystallite breakdown mechanism for damage accumulation,

642 an increase in available ‘edge’ sites will be beneficial to the crystal’s sorption capacity up to a certain
643 threshold. Further work on damaged mineral sorption and reactivity is required. It is likely that edge
644 site sorption will be more important than interlayer uptake in this context (Sawhney, 1972),
645 particularly for uranium.

646

647 Reduction of the solid phase iron as a result of irradiation (discussed earlier) may be advantageous in
648 a disposal case. Brookshaw et al (2013, 2014) observed considerable enhancement in cation uptake
649 (and consequent reductive transformation) by bio-reduced phyllosilicates; whilst net reduction in this
650 study is not on the same scale as bio-reduction, redox mechanisms clearly will have a crucial role in
651 GDF models. Fe(II) in sorbing minerals can reduce edge sited uranium from the mobile uranyl
652 $[U(VI)O_2]^{2+}$ cation to its immobile U(IV) oxidation state (Ilton et al., 2004; Myllykyla, 2008). Indeed,
653 electron hopping along the octahedral plane (Fe sublattice) has been proposed to reduce uranyl (or
654 similar) sorbed to edge sites in annite, thereby Fe²⁺ at ‘deeper’ sites within the structure can still
655 influence redox activity at the near-surface of a crystal (Rosso & Ilton, 2003, 2005), which provides
656 an Fe reduction mechanism by α -particle damage at depth in a crystal. With increasing dose however,
657 any perceived ‘benefits’ of Fe reduction by α -particles may be lost as the mica loses its structure
658 entirely, depending on the sorption capacity of highly damaged regions. Extensive further work on the
659 effect of radiation damage on sorption to sheet silicates with subsequent reductive precipitation is
660 underway.

661

662 Biotite mica is clearly highly susceptible to chemical and structural aberrations as a result of high
663 fluence α -particle bombardment. Indeed, at the maximum dose delivered in this study (~ 6.8 dpa)
664 amorphous areas dominate the structure; the crystalline regions that remain are heterogeneously
665 distorted over areas < 60 μ m. Whilst mica phases like biotite are unlikely to experience a near-field
666 radiation dose, the structure shares important parallels with the phyllosilicate bentonite (comprising
667 dominantly montmorillonite), which forms a crucial component of the engineered barrier system in a
668 GDF to protect the waste container against corrosion and limit the release of radionuclides. A limiting

669 factor for using mica as a proxy for bentonite is the expandability of clay phases with increased
670 hydration; a property not shared by micas excepting under weathering reactions ('hydrobiotite'
671 consists of interstratification of biotite and its hydrated alteration product, vermiculite (Brindley,
672 1983)). Despite this, many of the structural aberrations that biotite mica undergoes during α -particle
673 irradiation will be shared by layered clays; chiefly interlayer collapse, increased mosaicity and the
674 potential for Fe(III) reduction. Indeed, Ferrage et al., (2005) comment that a thermal 'plume'
675 surrounding waste canisters in the short term may cause smectite (a 2:1 expandable clay) to transform
676 into non-expandable illite. Further radiation damage studies of both hydrated and dehydrated clay
677 systems are needed to full understand their changing properties and performance in a developing
678 GDF.

679

680 Whilst many radiation damage models display bulk, time averaged lattice responses to increased point
681 defect densities; the data presented here demonstrates the necessity to consider the micro-scale spatial
682 evolution of radiation damage as a result of both heterogeneous ion bombardment, which is of
683 particular relevance when considering canister rupture and the subsequent complex irradiation effects
684 on the buffer material and the heterogeneous nature of the resulting damage. α -particles are not the
685 only radiation field present in the near-field of a GDF; combined irradiation studies (α, γ and heavy
686 ions to simulate α -recoil from adsorbed actinides) upon similar structures are currently underway.

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737 **References**

- 738 Allard, T., and Calas, G. (2009) Radiation effects on clay mineral properties. *Applied Clay Science*,
739 43, 143–149.
- 740 Ames, L.L., McGarrah, J. E, and Walker, B.A (1983) Sorption of uranium and radium by biotite,
741 muscovite and phlogopite. *Clays and Clay Minerals*, 31, 343–351.
- 742 Biro, L.P, Gyulai, J., and Havancsak, K. (1997) Comparison of damage produced by 209MeV Kr
743 irradiation in muscovite mica, graphite and silicon. *Nuclear Instruments and Methods in*
744 *Physics Research B*, 122, 476–480.
- 745 Bower, W.R., Smith, A.D., Pattrick, R.A.D., and Pimblott, S.M. (2015) Establishing α -particle
746 radiation damage experiments using the Dalton Cumbrian Facility's 5 MV tandem pelletron.
747 *Reviews of Scientific Instruments*, 86, 046105.
- 748 Bower W. R., Pearce, C.I., Droop, G. T. R., Mosselmans, J. F. W., Geraki, K., Pattrick, .R .A D.
749 (2015) Radiation damage from long term alpha particle bombardment of silicates - a
750 microfocus XRD and Fe K-edge XANES study. *Mineralogical Magazine*, In Press
- 751 Chailley, V., Dooryh e, E., Bouffard, S., Balanzat, E., and Levalois, M. (1994) Observations by X-ray
752 diffraction of structural changes in mica irradiated by swift heavy ions. *Nuclear Instruments*
753 *and Methods in Physics Research Section B: Beam Interactions with Materials and Atoms*,
754 91, 162–167.
- 755 Chailley, V., Dooryhee, E., and Levalois, M. (1996) Amorphisation of mica through the formation of
756 GeV heavy ion tracks. *Nuclear Instruments and Methods in Physics Research B*, 107, 199–
757 203.
- 758 Chaussidon, J. (1971) The I.R spectrum of structural hydroxyls of K⁺ depleted biotites. *Clays and*
759 *Clay Minerals*, 20, 59–67.
- 760 Dent, A.J., Cibin, G., Ramos, S., Smith, A.D., Scott, S.M., Varandas, L., Pearson, M.R., Krumpa,
761 N.A., Jones, C.P., and Robbins, P.E. (2009) B18: A core XAS spectroscopy beamline for
762 Diamond. *Journal of Physics: Conference Series*, 190, 012039.
- 763 Fialips, C. I, Huo, D., Yan, L., Wu, J., and Stucki, J.W. (2002) Effect of Fe oxidation state on the IR
764 spectra of Garfield nontronite. *American Mineralogist*, 87, 630–641.
- 765 Fleet, M.E, and Deer, W.A. Howie, R.A. Zussman, J. (2003) *Rock-forming Minerals: Micas, Second.*,
766 765 p. Vol. 3A. The Geological Society.
- 767 Holland, T.J.B., and Redfern, S.A.T (1997) Unit cell refinement from powder diffraction data: the use
768 of regression diagnostics. *Mineralogical Magazine*, 61, 65–77.
- 769 Idemitsu, K., Obata, K., Furuya, H., and Inagaki, Y. (1995) Sorption behaviour of uranium (VI) on a
770 biotite mineral. *Scientific basis for nuclear waste management*, 18.

- 771 Igor Pro (2014). WaveMetrics, Inc.
- 772 Ilavsky, J. (2012) Nika - software for 2D data reduction. *J. Appl. Cryst.*, 45, 324–328.
- 773 Ilton, E.S., Haiduc, A., Moses, C.O., Heald, S.M., Elbert, D.C., and Veblen, D.R. (2004)
774 Heterogeneous reduction of uranyl by micas: Crystal chemical and solution controls.
775 *Geochimica et Cosmochimica Acta*, 68, 2417–2435.
- 776 Leay, L., Bower, W., Horne, G., Wady, P., Pottinger, M., Nancekievill, M., Smith, A.D., Watson, S.,
777 Green, P.R., Lennox, B., and others (2015) Development of irradiation capabilities to address
778 the challenges of the nuclear industry. *Nuclear Instruments and Methods in Physics Research*
779 *B*, 343, 62–69.
- 780 Lee, S.Y., and Tank, R.W. (1985) Role of clays in the disposal of nuclear waste: A review. *Applied*
781 *clay science*, 1, 145–162.
- 782 Lee, S.Y, Baik, M. H., Lee, Y. J., and Lee, Y. B. (2009) Adsorption of U(VI) ions on biotite from
783 aqueous solutions. *Applied Clay Science*, 46, 255–259.
- 784 Lieser, K.H. (1997) *Nuclear and Radiochemistry: Fundamentals and Applications*, First. Vols. 1-1,
785 Vol. 1. VCH.
- 786 Lockwood, G.K., and Garofalini, S.H. (2010) Effect of moisture on the self-healing of vitreous silica
787 under irradiation. *Journal of Nuclear Materials*, 400, 73–78.
- 788 Lu, F., Huan, M., Yaqoob, F., Lang, M., Namavar, F., Trautmann, C., Sun, H., Ewing, C., and Lian, J.
789 (2012) Displacive radiation-induced structural contraction in nanocrystalline ZrN. *Applied*
790 *Physics Letters*, 101.
- 791 Mosselmans, J.F.W., Quinn, P.D., Dent, A.J., Cavill, S.A., Moreno, S.D., Peach, A., Leicester, P.J.,
792 Keylock, S.J., Gregory, S.R., Atkinson, K.D., and others (2009) I18 – the microfocus
793 spectroscopy beamline at the Diamond Light Source. *Journal of Synchrotron Radiation*, 16,
794 818–824.
- 795 Myllykylä, E. (2008) Reduction of uranium in disposal conditions of spent nuclear fuel pp. 1–111.
796 Working Report, Posiva.
- 797 Nasdala, L., Wenzel, M., Andrut, M., Wirth, R., and Blaum, P. (2001) The nature of radiohaloes in
798 biotite: Experimental studies and modeling. *American Mineralogist*, 86, 498–512.
- 799 Nasdala, L., Wildner, M., Wirth, R., Groschopf, N., Pal, D.C., and Möller, A. (2006) Alpha particle
800 haloes in chlorite and cordierite. *Mineralogy and Petrology*, 86, 1–27.
- 801 Nielsen, C. H. and Sigurdsson, H., (1981) Quantitative methods for electron microprobe
802 analysis of sodium in natural and synthetic glasses, *American Mineralogist*, 66, 557-552
803
- 804 Palmer, D., and Conley, M. (1994) *CrystalMaker*. English, CrystalMaker Software Ltd.
- 805 Patrick, R. A. D. Charnock, J. M., Geraki, T., Mosselmans, J. F. W., Pearce, C. I., Pimblott, S. and
806 Droop, G. T. R. (2013) Alpha particle damage in biotite characterised by microfocus X-ray
807 diffraction and Fe K-edge X-ray absorption spectroscopy. *Mineralogical Magazine*.
- 808 Ravel, B., and Newville, M. (2005) ATHENA, ARTEMIS, HEPHAESTUS: data analysis for X-ray
809 absorption spectroscopy using IFEFFIT. *Journal of Synchrotron Radiation*, 12, 537–541.

- 810 Reed, D.T, Donar, S.D, and Weiner, M.F (1987) Gamma and alpha radiation levels in a basalt high
811 level waste repository: potential impact upon container corrosion and packing properties.
812 Coupled Processes Associated with Nuclear Waste Repositories, 1, 325–338.
- 813 Rickwood, P.C. (1981) The largest crystals. American Mineralogist, 66, 885–907.
- 814 Rieder, M., Cavazzini, G., D'Yakonov, Y., Gottardi, G., Guggenheim, S., Koval, P., Muller, G.,
815 Neiva, A.M., Radoslovich, E., Robert, J., and others (1999) Nomenclature of the micas.
816 Mineralogical Magazine, 63, 267–279.
- 817 Rosso KM, and ES Ilton. 2003. "Charge transport in micas: The kinetics of FeII/III electron transfer
818 in the octahedral sheet." Journal of Chemical Physics 119(17):9207-9218.
- 819 Rosso KM, and ES Ilton. 2005. "Effects of Compositional Defects on Small Polaron Hopping in
820 Micas." *Journal of Chemical Physics* 122(24):244709
- 821 Savage, D., and Arthur, R. (2012) Exchangeability of bentonite buffer and backfill materials p. 57.
822 STUK Radiation and Nuclear Safety Authority.
- 823 Sawhney, B.L. (1972) Selective sorption and fixation of cations by clay minerals: a review. Clays and
824 Clay Minerals, 20, 93–100.
- 825 Stadelmann, P. (2012) Jems Electron Microscopy Software.
- 826 Strobel, P., Durr, J., Tuilier, M.-H., and Charenton, J.-C. (1993) Extended X-ray absorption fine
827 structure study of potassium and caesium phyllosulfates. J. Mater. Chem., 3, 453–458.
- 828 The MathWorks Inc. (2013) MATLAB R2013a.
- 829 Thermo Electron Corporation (2004) OMNIC.
- 830 Vedder, W. (1969) Dehydroxylation and rehydroxylation, oxidation and reduction of micas. American
831 Mineralogist, 54, 482–509.
- 832 Wang, L.M., Wang, S.X., Gong, W.L., and Ewing, R.C. (1998) Temperature dependence of Kr ion-
833 induced amorphization of mica minerals. Nuclear Instruments and Methods in Physics
834 Research Section B: Beam Interactions with Materials and Atoms, 141, 501–508.
- 835 Wilkins, R.W.T (1967) The hydroxyl-stretching region of the biotite mica spectrum. Mineralogical
836 Magazine, 36, 279–325.
- 837 Wolfer, W.G. (2000) Radiation effects in plutonium: What is known? Where should we go from here?
838 Los Alamos Science, 26, 274.
- 839 Ziegler, J.F. (2013) The Stopping and Range of Ions in Matter (SRIM).
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848 **Figure Captions**

849

850 **Figure 1.** CrystalMaker (Palmer, D. and Conley, M. 1994) image of biotite mica structure, viewed
851 along the phyllosilicate layers, illustrating the sequential arrangement of the TOT sheets. Large,
852 potassium-rich interlayers create planes of weakness accounting for the micas perfect cleavage. OH
853 groups (green label) occupy the voids between hexagonal silica rings and are oriented towards the
854 octahedral layer. Samples were mounted such that incident ions penetrated across the parallel layers,
855 as shown by the dashed arrow.

856

857 **Figure 2. A.** Optical photograph of a cleavage parallel plane of biotite mica, sectioned for analysis,
858 following exposure to 9.2×10^{15} $^4\text{He}^{2+}$ ions across the beam profile in **B**. Note the mottled, silver
859 discolouration of the irradiated area, marked by the dashed line. **B.** Cross-section through the
860 Gaussian profile of the $^4\text{He}^{2+}$ ion beam footprint shown in Figure 2A. Solid black lines (sections **i**, **ii**
861 and **iii**) represent areas selected for sequential XRD traverses (each with a track step of 30 μm as
862 shown in Figure 4). Dashed black lines (section **iv**) display the area analysed at higher resolution (5
863 μm track steps), the dashed grey line indicates the ion beam limit, note that the area to the left of this
864 line in panel **i** has not been irradiated.

865

866 **Figure 3.** Representative high resolution 2D X-ray diffraction patterns (CCD images) from across the
867 irradiated biotite sample. Distances (μm) denote proximity from $^4\text{He}^{2+}$ ion beam centre, as illustrated
868 in Figure 2. The upper left panel displays the calibration pattern taken from powdered silicon, on
869 which three major reflections are marked (2θ angles).

870

871 **Figure 4.** Transmission X-ray diffraction ‘map’ displaying X-ray diffraction angle changes (2θ)
872 across the [001] oriented biotite, plotted as a function of distance from the ion beam centre (0 μm
873 denotes height of beam intensity, see Table 2). The dashed line denotes the limit of the ion beam,
874 below which the sample has not been irradiated. Panels **i** – **iv** represent areas displayed in Figure 2
875 and correspond to relative changes in beam intensity across the Gaussian beam profile. Panels **i**, **ii** &
876 **iii** show a 7000 μm trace with diffraction patterns collected at 30 μm intervals, whilst panel **iv** denotes
877 a 300 μm trace with patterns collected at 5 μm intervals over a region of interest (4000-4300 μm).

878

879 **Figure 5.** 1D X-ray diffraction patterns showing the changes in peak position and intensity as a
880 function of ion beam dose. Patterns have been selected at 1000 μm intervals across the traverse shown
881 in Figure 4. Dotted lines display the reduced 1D data from the CCD images, solid lines represent
882 fitted peaks overlaid onto the 1D data.

883

884 **Figure 6.** Relative changes in *a* and *b* lattice dimensions of irradiated biotite as a function of distance
885 from beam centre (0 μm represents maximum dose) calculated using *UnitCell* (Holland and Redfern,
886 1997).

887

888 **Figure 7. A.** Representative *basal plane parallel*, 1D diffraction patterns at increasing depths into the
889 sample from the irradiated surface of the crystal. Dotted lines display the reduced 1D data from the
890 CCD images, solid lines represent fitted peaks overlaid onto the 1D data. **B.** SRIM modeling predicts
891 that $^4\text{He}^{2+}$ ions have a penetration distance of 18 μm in biotite mica, thus patterns at 25 and 30 μm are
892 assumed to be unirradiated. The bar chart shows simulated knock-on displacements caused by
893 incident $^4\text{He}^{2+}$ ions into model biotite density and chemistry (simulations do not account for mineral
894 structure). The highest displacement density occurs between 15-20 μm (with a peak at 18 μm) as the
895 incident ions lose enough energy to permit structural interaction within the target. Simulations were
896 run using SRIM (Ziegler, 2013).

897

898 **Figure 8.** Single crystal X-ray diffraction patterns acquired with a Bruker D8 CuK α lab source across
899 two irradiated biotite samples and one standard. Δd denotes the change in d spacing (\AA) between the
900 unirradiated (0 dpa) and the highest dose (0.18 dpa) patterns for each plane. Due to the nature of the
901 prepared cleavage parallel samples, only reflections parallel to the phyllosilicate layers were acquired.
902 Dotted lines display the 1D data, solid lines represent fitted peaks overlaid.

903

904 **Figure 9. A.** Fe-K edge X-ray absorption near edge spectra of irradiated ($\sim 2000 \mu\text{m}$ from $^4\text{He}^{2+}$ beam
905 centre) and unirradiated biotite, spaced for clarity. Note the $\sim 2\text{eV}$ energy shift of the absorption edge
906 following exposure to the ion beam.

907

908 **Figure 10.** Iron K-edge Extended X-ray Absorption Fine Structure (EXAFS) oscillations (k^3) for
909 biotite irradiated with sequentially increasing fluences of $^4\text{He}^{2+}$ ions. Black solid lines show the
910 EXAFS oscillations whilst red dashed lines display the fit to the data. Adjacent is the non phase-
911 corrected Fourier transform magnitude for each of the corresponding EXAFS plots with fits (A-F). All
912 spectra have been spaced for clarity.

913

914 **Figure 11.** Potassium K-edge edge X-ray Absorption Near Edge Spectra (XANES) for both
915 unirradiated and irradiated biotite. Inset is the non phase-corrected Fourier transform magnitude for
916 the corresponding spectra, note the slight contraction of the first and second ‘shells’.

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918 **Figure 12.** Intensity-normalized Fourier transform infra-red absorption spectra, centred over the OH $^-$
919 stretching region of unirradiated, heated and irradiated biotite.

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926 **Table 1.** Irradiated sample details.

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Sample	Current (nA)	Time (mins)	Ion fluence*	dpa**
1	160	324	9.6×10^{15}	0.28
2	100	288	6×10^{15}	0.18
3	75	175	2×10^{15}	0.06

928

929 *total accumulated $^4\text{He}^{2+}$ ions over the beam footprint ($r = 6.6$ mm)

930 ** Displacements per atom calculated based on the average ion beam density over the beam spot,

931 based on SRIM modeling and extrapolation of ion densities. Differences in dose across a single

932 sample is shown in Table 2.

933 **Table 2.** Dose estimations across beam profile for sample 1.

934

Dist. from beam centre (μm)	Ion fluence*	dpa**
7000-6750	Unirradiated	0
6750-5750	2.67×10^{13}	0.01
5750-4730	1.67×10^{14}	0.02
4730-3720	4.85×10^{14}	0.10
3720-2700	1.05×10^{15}	0.20
2720-1690	1.86×10^{15}	0.40
1690-1010	1.31×10^{15}	0.70
1010-690	1.49×10^{15}	2.60
690-0	3.20×10^{15}	6.77

935

936

937 *total accumulated $^4\text{He}^{2+}$ ions over the area specified integrated over a 323 minute exposure to a

938 ~ 200 nA beam, estimated by on-sample current monitoring.

939 ** Displacements per atom calculated based on the ion beam density within the area specified, note

940 the differences in area width accounting for non-linear increase in dpa into the beam centre. Atomic

941 displacement parameters for biotite calculated using SRIM (Ziegler, 2013).

942

943

944 **Table 3.** Fe *K*-edge EXAFS fit results for unirradiated and irradiated biotite. Dist. denotes approximate distance
 945 from the ion beam centre (height of α -particle fluence). CN denotes coordination number; R denotes interatomic
 946 distance; σ^2 denotes Debye-Waller factor; $S0^2$ denotes the amplitude reduction factor; *r* denotes the ‘goodness
 947 of fit’ factor. ‘*’ denotes a fixed parameter.

Dist. (μm)	Spectrum	Path	CN	R (\AA)	σ^2 (\AA^2)	$S0^2$	<i>r</i>
Unirradiated	A	Fe-O	6.0*	2.07 ± 0.01	0.016 ± 0.001	0.90	0.004
		Fe-Fe	6.0*	3.11 ± 0.03	0.020*		
		Fe-Si	4.0*	3.38 ± 0.01	0.010 ± 0.002		
~ 3000 \pm 100	B	Fe-O	6.0*	2.10 ± 0.01	0.010 ± 0.001	0.70	0.012
		Fe-Fe	6.0*	3.12 ± 0.02	0.020 ± 0.005		
		Fe-Si	4.0*	3.40 ± 0.01	0.006 ± 0.002		
~ 2500 \pm 100	C	Fe-O	6.0*	2.09 ± 0.01	0.013 ± 0.001	0.78	0.006
		Fe-Fe	6.0*	3.03 ± 0.03	0.020 ± 0.005		
		Fe-Si	3.0*	3.37 ± 0.01	0.010 ± 0.002		
~ 2000 \pm 100	D	Fe-O	6.0*	2.11 ± 0.01	0.009 ± 0.001	0.70	0.016
		Fe-Fe	6.0*	3.12 ± 0.03	0.013 ± 0.005		
		Fe-Si	3.0*	3.39 ± 0.03	0.007 ± 0.002		
~ 1500 \pm 100	E	Fe-O	6.0*	2.08 ± 0.01	0.013 ± 0.001	0.75	0.006
		Fe-Fe	5.0*	3.01 ± 0.03	0.030 ± 0.005		
		Fe-Si	3.0*	3.38 ± 0.02	0.010 ± 0.002		
~ 1000 \pm 100	F	Fe-O	6.0*	2.11 ± 0.01	0.008 ± 0.001	0.75	0.009
		Fe-Fe	6.0*	3.12 ± 0.03	0.012 ± 0.005		
		Fe-Si	4.0*	3.39 ± 0.01	0.005 ± 0.002		

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964 **Table 4.** Formula-normalized major element results of the EPMA traverse into the irradiated biotite
965 (Std error \pm 0.02). ‘Dist.’ represents the distance from the centre of the ion beam spot (peak of
966 radiation damage). The point identified by a ‘*’ denotes the limit of the ion beam as marked by the
967 colour change on the biotite surface. ‘Total’ denotes the sum of all cations. The sum of Na + K has
968 also been expressed for charge balance illustration (see text for further details).

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Dist. (mm)	Al	Fe^{TOT}	Si	Ti	Na	Mg	K	Total	Na + K
9.0	2.35	2.46	5.48	0.37	0.03	2.96	1.95	15.60	1.99
8.5	2.45	2.54	5.27	0.37	0.04	3.13	1.98	15.78	2.02
8.0	2.41	2.46	5.43	0.36	0.04	2.99	1.94	15.63	1.98
7.5	2.39	2.48	5.45	0.36	0.03	2.97	1.93	15.61	1.96
7.0	2.47	2.48	5.41	0.36	0.03	2.92	1.94	15.61	1.97
6.5	2.45	2.46	5.42	0.35	0.02	2.99	1.93	15.62	1.96
*6.0	2.50	2.46	5.34	0.36	0.03	3.02	1.96	15.67	1.99
5.5	2.44	2.43	5.46	0.37	0.02	2.93	1.91	15.56	1.93
5.0	2.37	2.41	5.58	0.35	0.03	2.85	1.88	15.48	1.91
4.5	2.41	2.43	5.51	0.36	0.04	2.88	1.91	15.53	1.95
4.0	2.43	2.46	5.46	0.35	0.03	2.93	1.91	15.58	1.94
3.5	2.27	2.42	5.60	0.36	0.04	2.92	1.91	15.51	1.95
3.0	2.36	2.40	5.63	0.36	0.04	2.78	1.85	15.42	1.89
2.0	2.34	2.43	5.60	0.36	0.02	2.83	1.88	15.47	1.90
1.5	2.32	2.39	5.56	0.36	0.05	2.95	1.88	15.52	1.93
1.0	2.37	2.45	5.54	0.36	0.04	2.86	1.88	15.52	1.93
0.0	2.44	2.41	5.46	0.36	0.03	2.94	1.93	15.57	1.96

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SUPPORTING INFORMATION

989

990 *Linear Combination Fitting*

991 Linear combination fitting (LCF) was performed using Athena, as part of the Demeter XAS data
992 processing suite (Ravel & Newville, 2005). Hematite Fe(III)₂O₃ and Wüstite Fe(II)O standards were
993 used for relative comparisons between the two XANES spectra presented in the main manuscript
994 (Figure 9). Results are presented below; whilst there are errors in the fits, relative comparisons
995 between the two spectra will be valid.

996

997 **1. Standard Biotite**

998

999 LCF fit of biotite_std_1.dat as flattened mu(E) from 7114.4808 to 7144.4808

1000

1001 Fit included 49 data points and 3 variables, and approximately 22.517 measurements

1002 Weights sum to 1: yes

1003 Weights forced between 0 and 1: yes

1004 Overall e0 shift used: no

1005 Noise added to data: 0

1006 R-factor = 0.0069107

1007 Chi-square = 0.08485

1008 Reduced chi-square = 0.0018053

1009

1010 . standard weight e0

1011

1012 Fe(III) Fe₂O₃-Hematite 0.542 (0.048) -1.637 (0.261)

1013 Fe(II) FeO- Wüstite 0.458 (0.048) 4.986 (0.184)

1014

1015

1016

1017 . sum 1.000

1018

1019

1020 **2. Irradiated Biotite**

1021

1022 LCF fit of biotite_irradiated_1.dat as flattened mu(E) from 7111.5 to 7141.5

1023

1024 Fit included 72 data points and 3 variables, and approximately 22.517 measurements

1025 Weights sum to 1: yes

1026 Weights forced between 0 and 1: yes

1027 Overall e0 shift used: no

1028 Noise added to data: 0

1029 R-factor = 0.0037686
1030 Chi-square = 0.07826
1031 Reduced chi-square = 0.0011342

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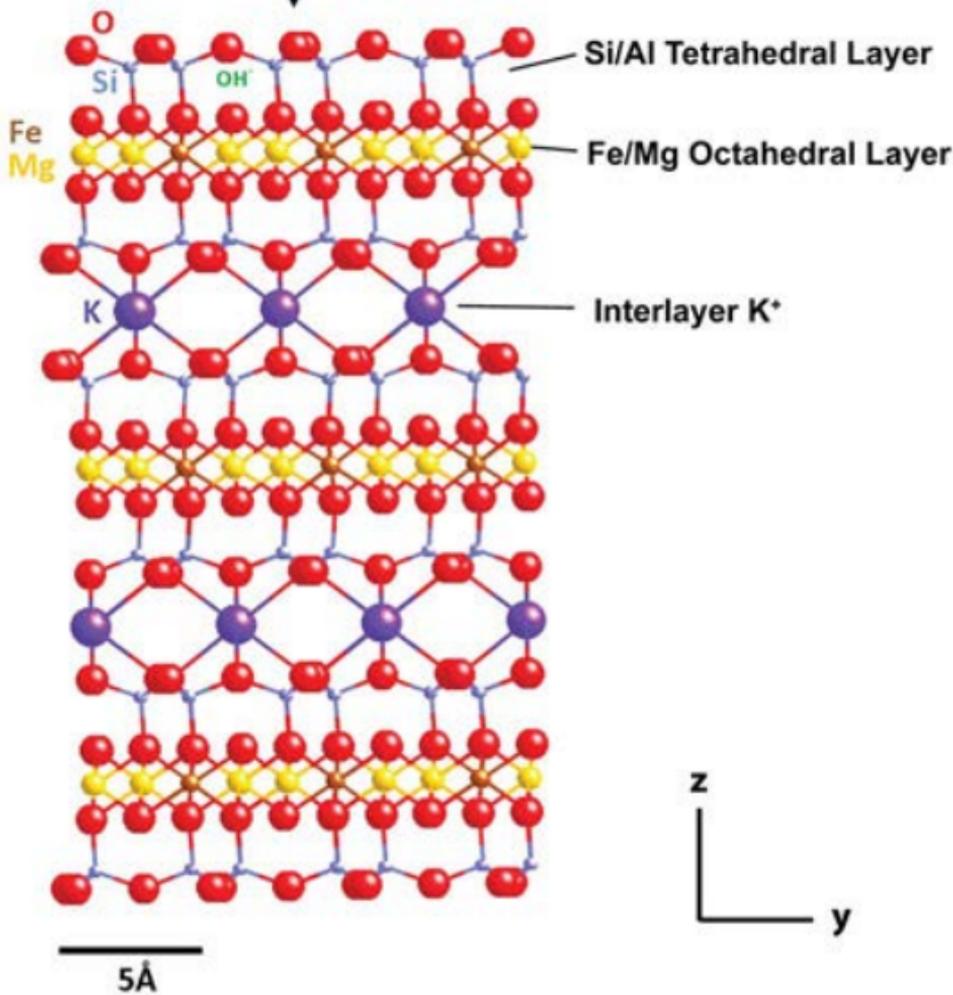
. standard	weight	e0
=====		
1035 Fe(III) Fe2O3-Hematite	0.369 (0.040)	-1.800 (0.279)
1036 Fe(II) FeO- Wüstite	0.631 (0.040)	2.560 (0.133)
1037		
1038 . sum	1.000	

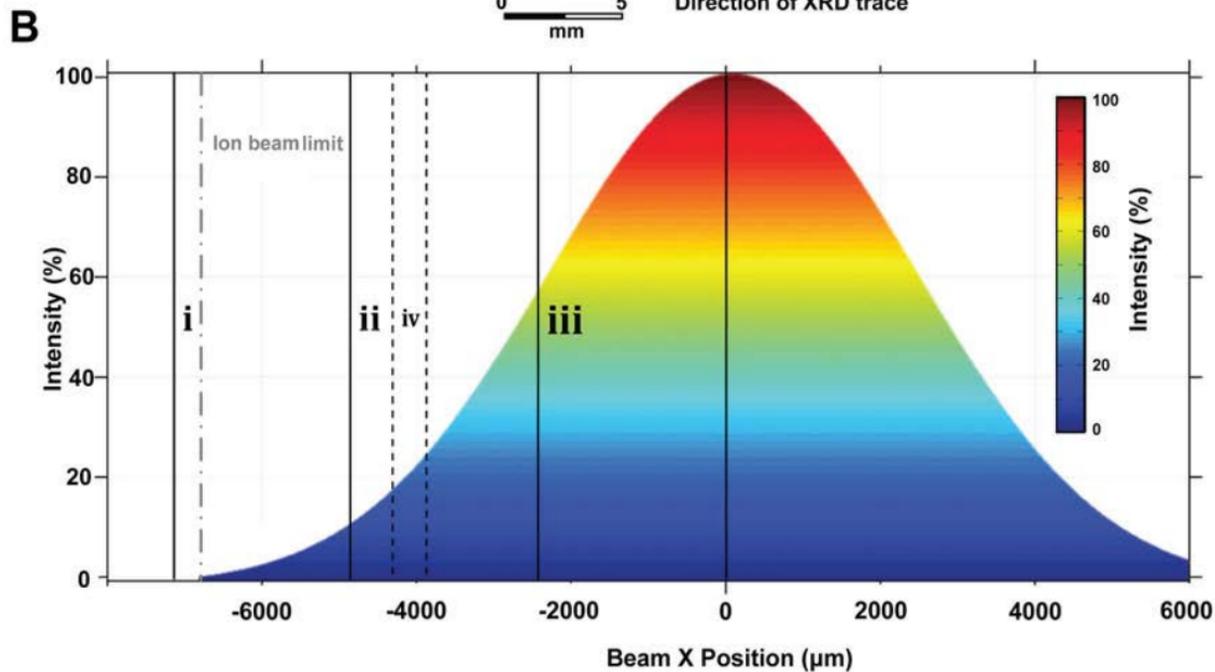
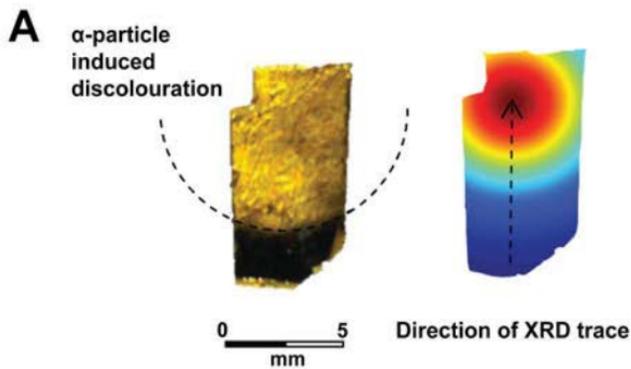
1039 **Figure S1.** Linear combination fit results for irradiated (above) vs standard (below) biotite. The

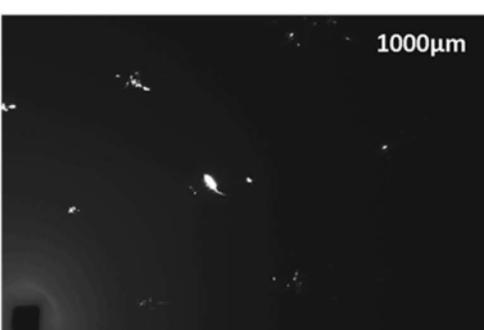
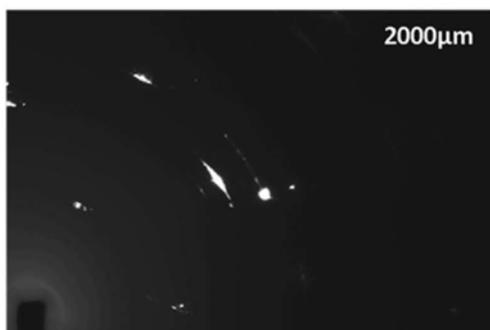
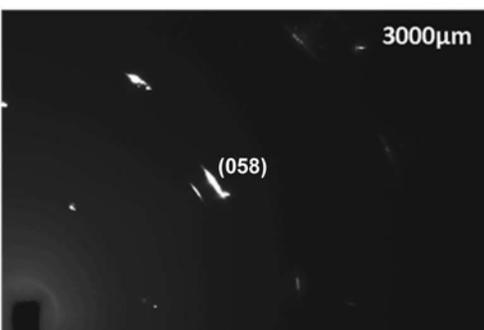
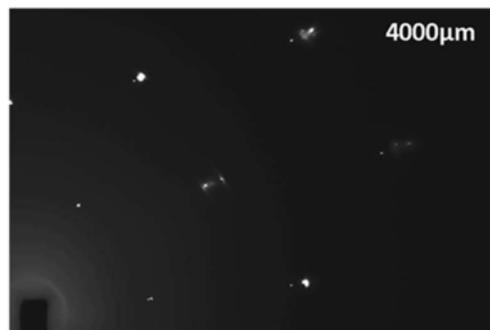
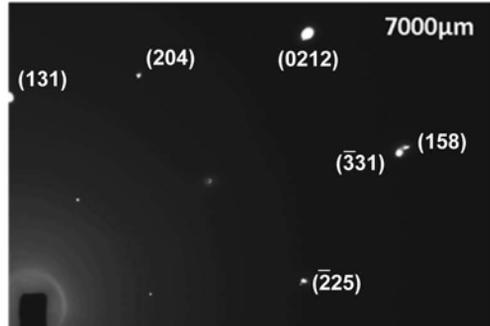
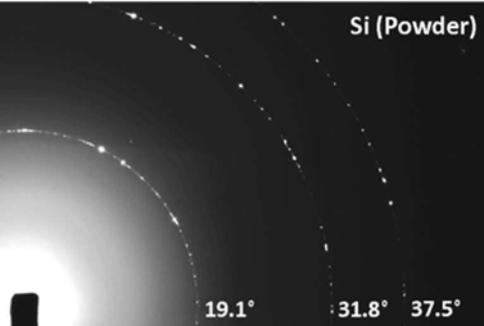
1040 irradiated spectra was collected at ~ 2000 μm from the centre of the $^4\text{He}^{2+}$ ion beam spot on sample 1.

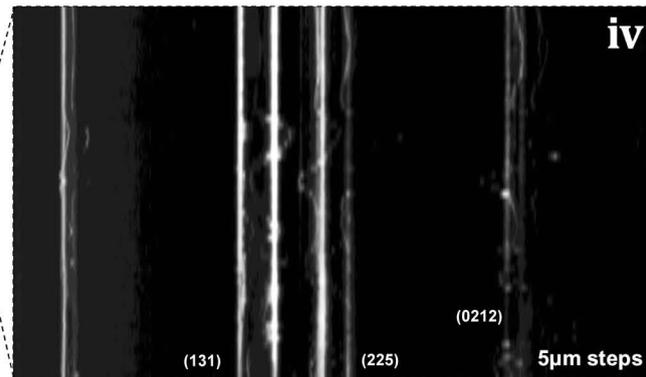
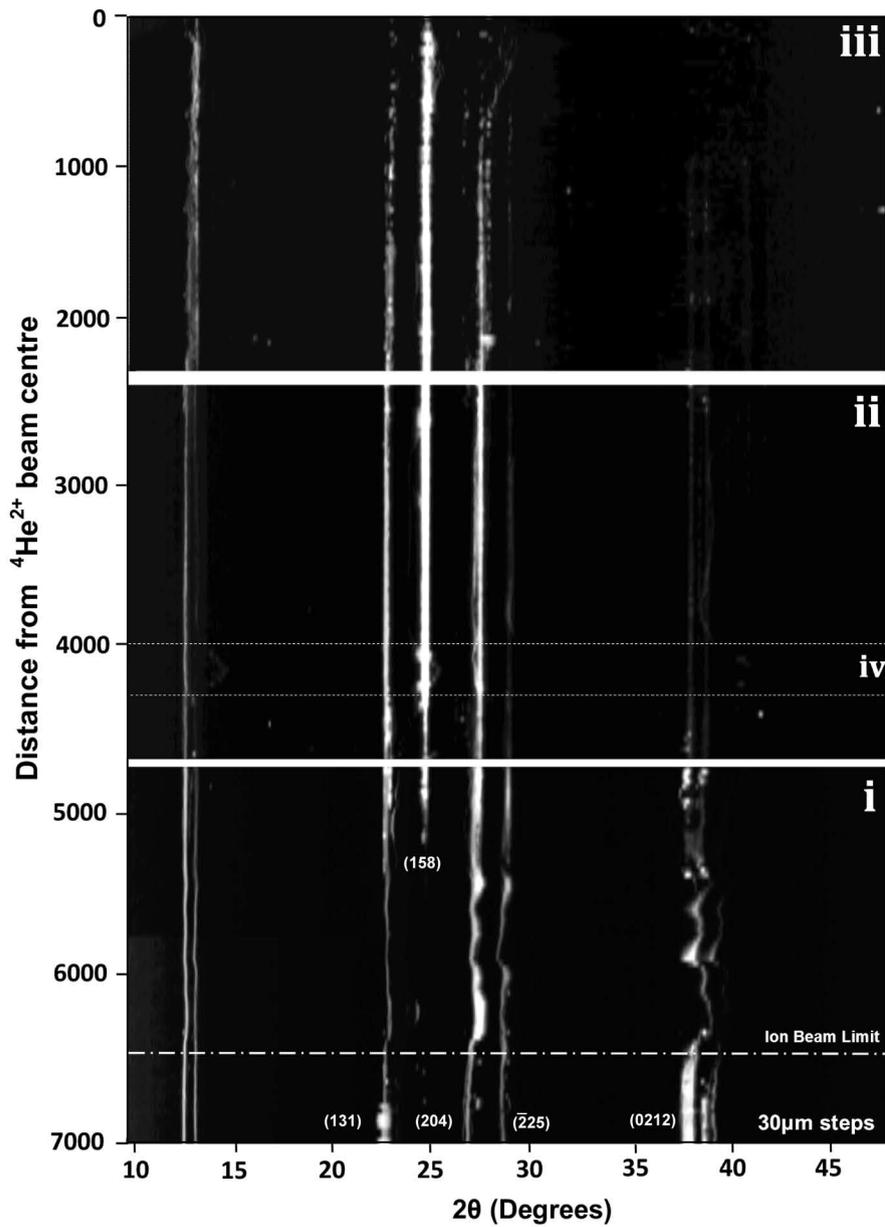
1041 See corresponding Figures 9 & 10 in the main manuscript.

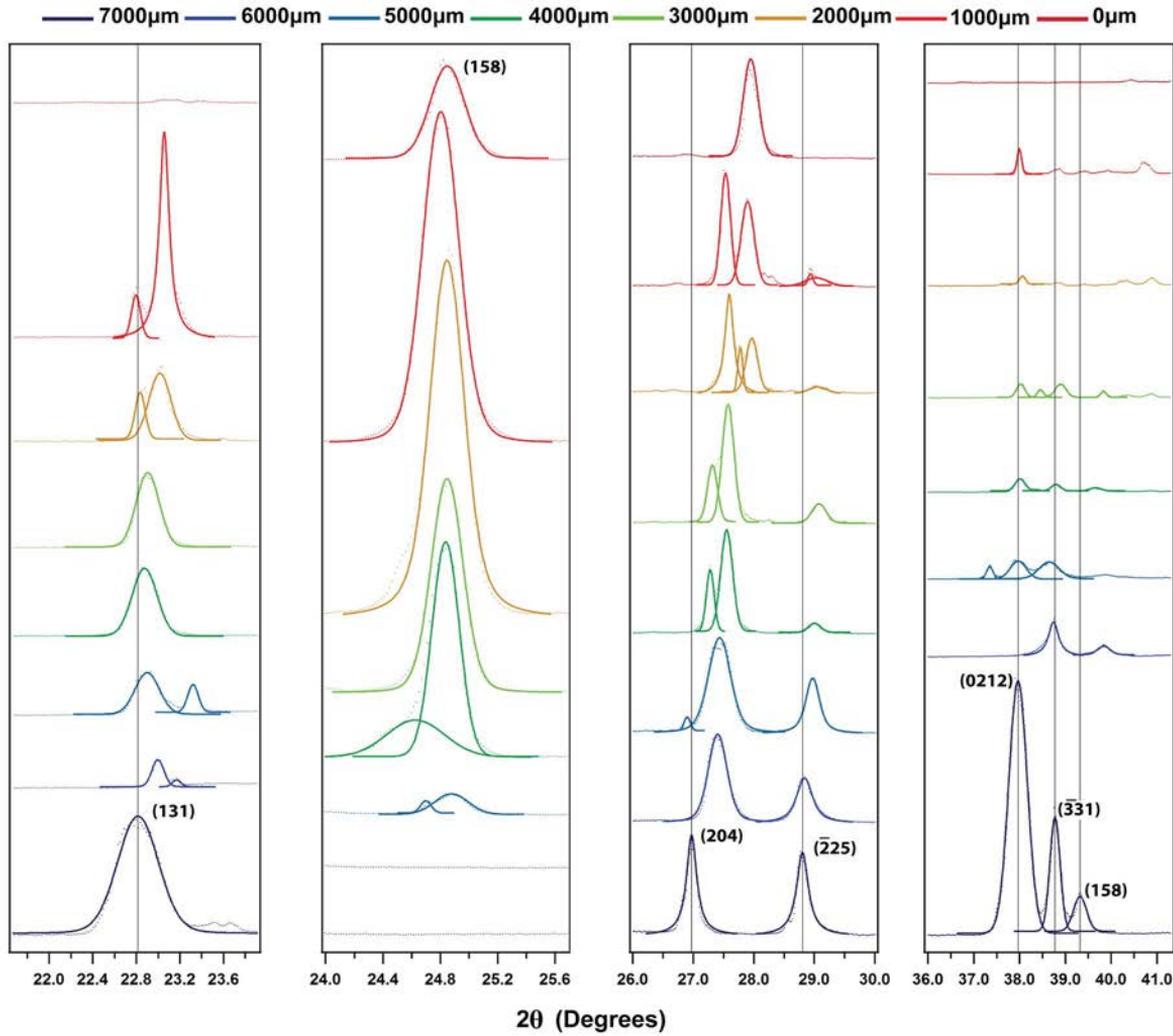
${}^4\text{He}^{2+}$ ion irradiation

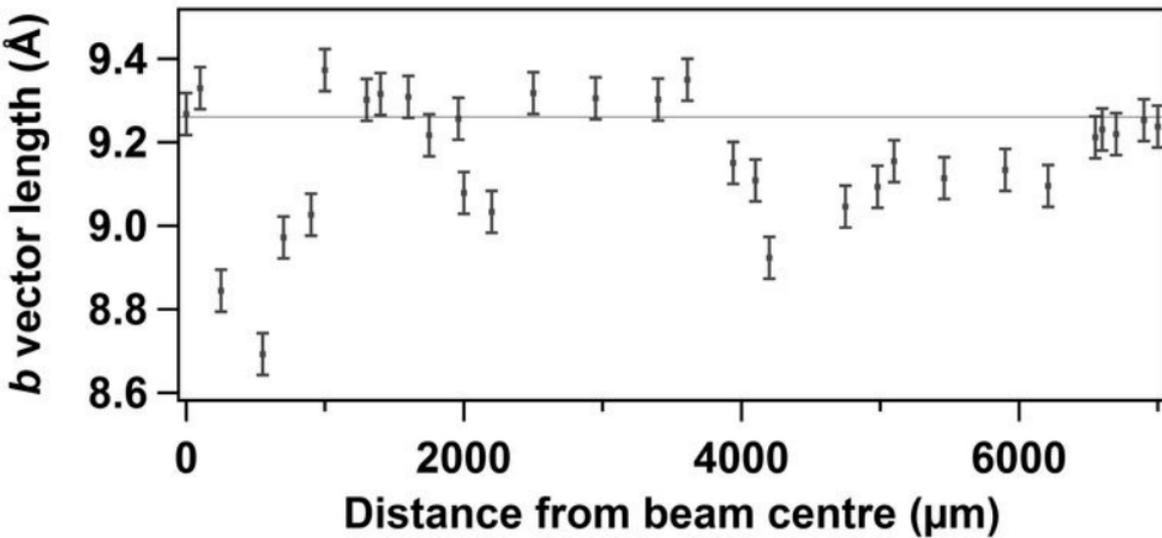
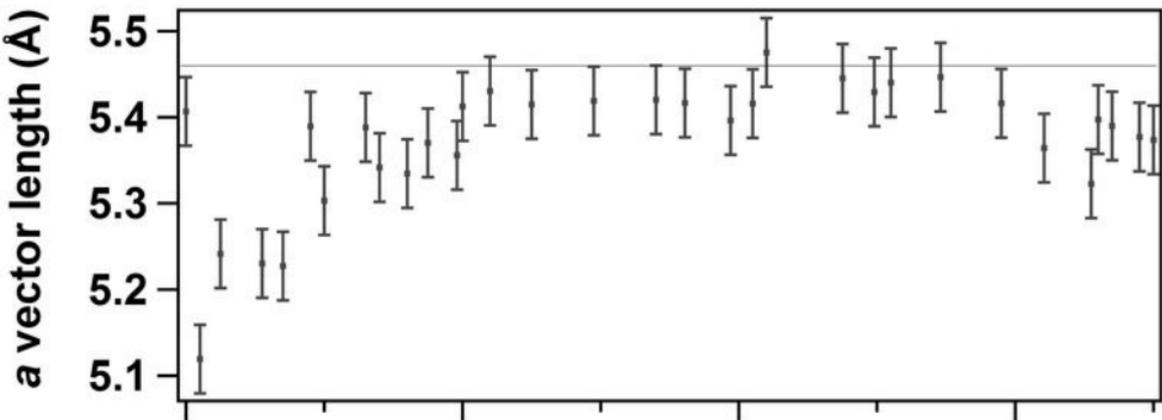








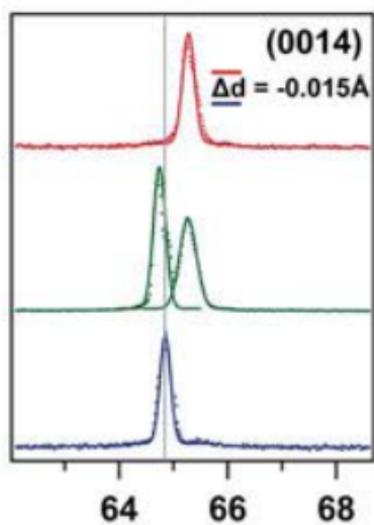
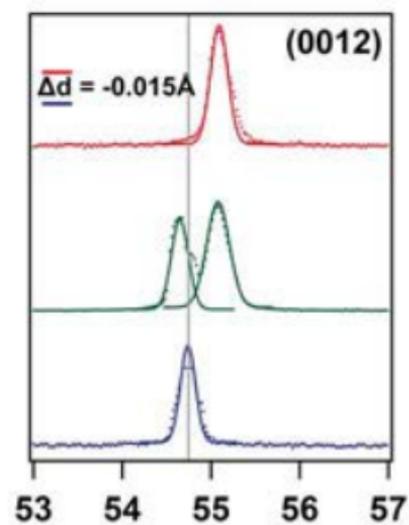
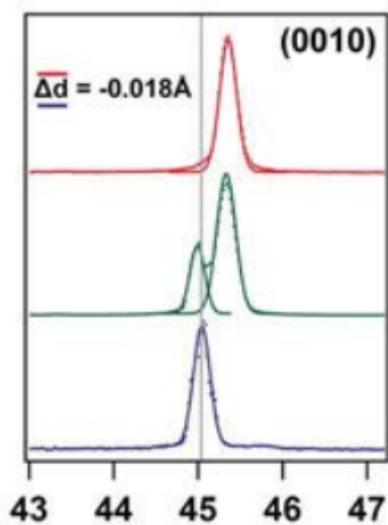
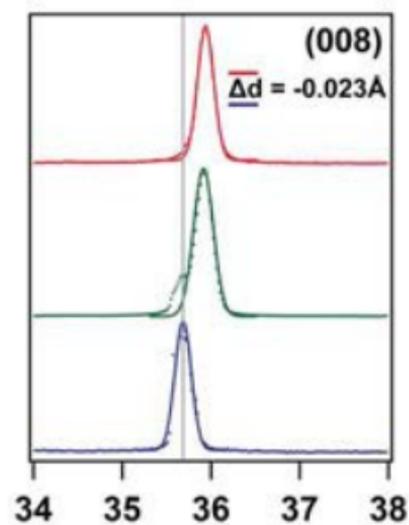
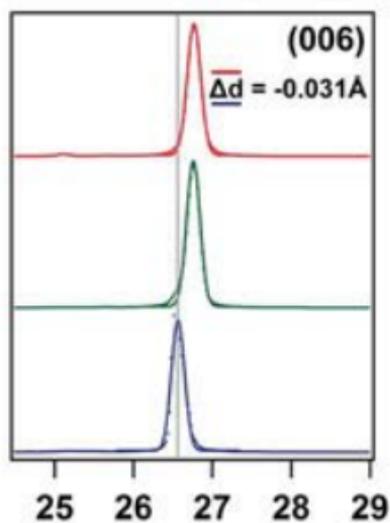
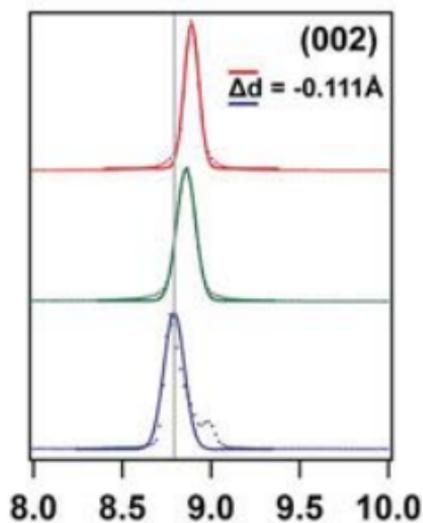


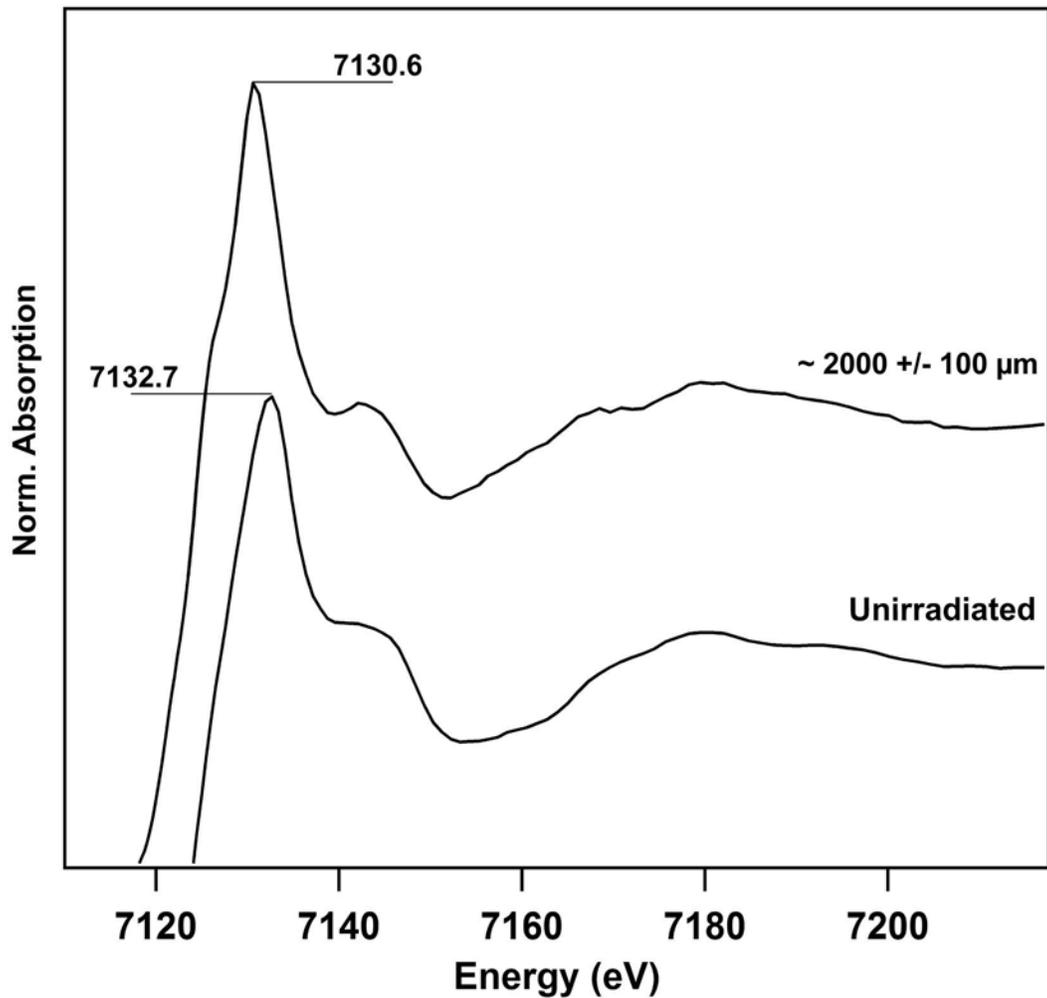


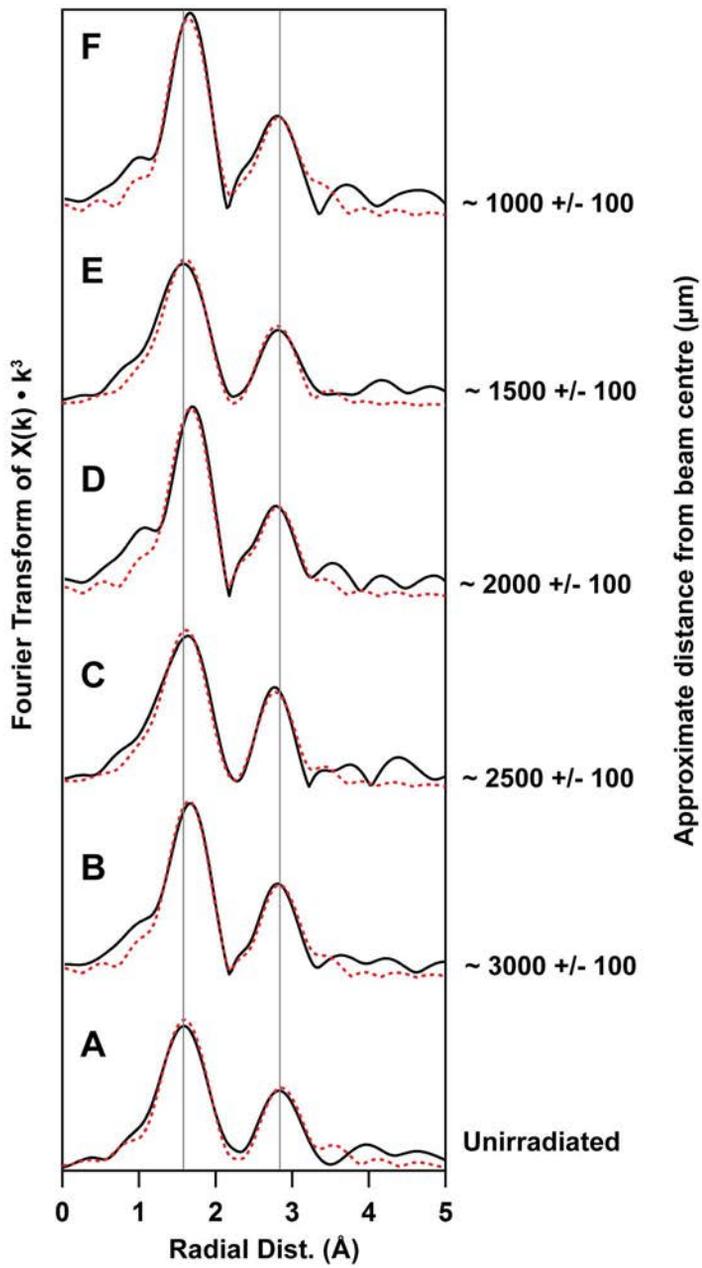
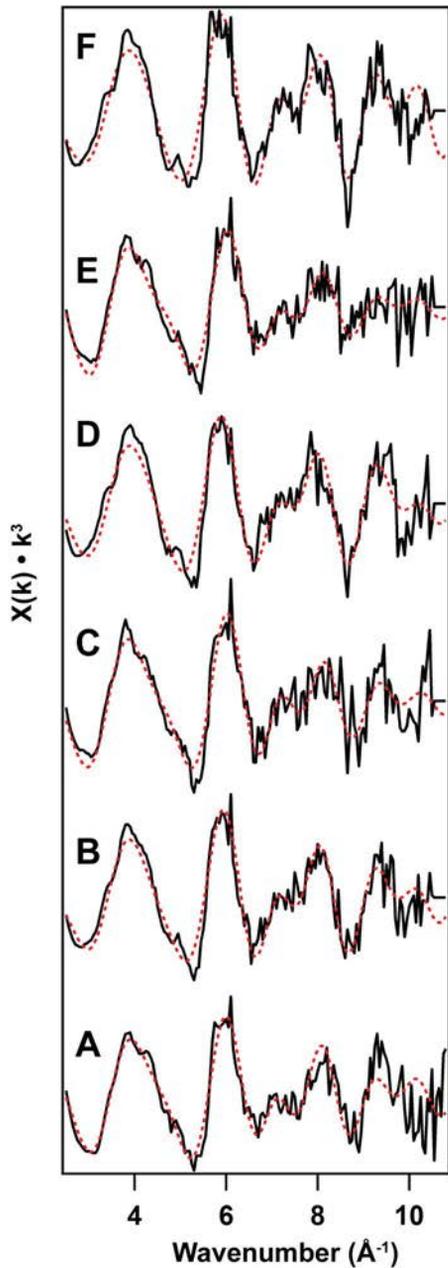
Unirradiated

0.06 dpa

0.18 dpa

 2θ (Degrees)





Norm Absorption

