1	Revision1
2	Equation of state of the high-pressure Fe ₃ O ₄ phase and a new structural
3	transition at 70 GPa
4	Angele Ricolleau ^{1,2} , Yingwei Fei ²
5	¹ Aix-Marseille Université, CNRS, Centre Interdisciplinaire de Nanoscience de Marseille,
6	UMR 7325, 13288 Marseille, France
7	² Geophysical Laboratory, Carnegie Institution of Washington, 5251 Broad Branch Road,
8	Washington, DC, 20015
9	
10	Abstract
11	We have investigated the high-pressure behavior of Fe ₃ O ₄ by in situ X-ray diffraction
12	measurements from 11 to 103 GPa. Up to 70 GPa, the previous observed high-pressure Fe_3O_4
13	phase (h-Fe ₃ O ₄) is stable, with a CaTi ₂ O ₄ -type structure. The compression curve shows an
14	abnormal volume contraction at about 50 GPa, likely associated with the magnetic moment
15	collapse observed at that pressure. Fitting the compression data up to 45 GPa to the Birch-
16	Murnaghan equation of state yields a bulk modulus, K_{T0} = 172 GPa, and V_0 = 277 Å ³ , with
17	fixed $K' = 4$. At a pressure between 64 and 73 GPa, a new structural transition was observed
18	in Fe ₃ O ₄ , which can be attributed to a martensitic transformation as described by Yamanaka et
19	al. (2008) for post-spinel structural transition. The diffraction data can be best fitted with a
20	Pnma space group. No breakdown of Fe ₃ O ₄ was observed up to at least 103 GPa. The new

high-pressure polymorph is about 6% denser than the $h-Fe_3O_4$ phase at 75 GPa.

23

24 Introduction

Iron exists in several oxidation states with ferrous (Fe^{2+}) and ferric (Fe^{3+}) iron being 25 the most common in the rock-forming minerals. Magnetite Fe₃O₄ has very interesting 26 characteristics because of the equal presence of Fe^{2+} and Fe^{3+} in its structure (Fleet, 1981). 27 The behavior of Fe_3O_4 at high pressure and temperature is fundamentally important for 28 understanding the oxidation state in the Earth's interior. A structural transition in magnetite 29 was observed at around 25 GPa (Mao et al., 1974) and its high-pressure phase was first 30 assumed to have a monoclinic structure. Fei et al. (1999) obtained X-ray diffraction data on 31 the high-pressure phase using an imaging plate detector and monochromatic synchrotron X-32 radiation and proposed an orthorhombic cell with a Pbcm space group. Further structure 33 analysis of the high-pressure phase (h-Fe₃O₄) by Haavik et al. (2000) indicated that the Cmcm 34 space group (CaTi₂O₄-type structure) would better fit the observed X-ray diffraction data. 35 Schollenbruch et al. (2011) studied the precise transition between magnetite and h-Fe₃O₄ and 36 37 observed the transition at 10 GPa and 1000 K.

The stability of h-Fe₃O₄ at high pressure has been debated. Because the predicted densities of the FeO + Fe₂O₃ assemblage would become higher than that of h-Fe₃O₄ at high pressure, it has been suggested that h-Fe₃O₄ could decompose into FeO and Fe₂O₃ at sufficiently high pressure (Haavik et al., 2000). Lazor et al. (2004) predicted a breakdown of the h-Fe₃O₄ at pressures higher than 50 GPa based on thermodynamical calculations. By combining experimental data and ab initio calculation, Dubrovinsky et al. (2003) predicted that the h-Fe₃O₄ phase would be stable up to 100 GPa.

The pressure effect on iron bearing oxides is further complicated because of possible spin transitions at high pressure (Badro et al., 1999). Depending of its concentration in the

oxide and valence state, the spin transitions from high spin to intermediate or to low spin state 47 can occur at different pressures (Badro et al., 2005). With ab initio calculation, Ju et al. (2012) 48 49 predicted a phase transition in magnetite from cubic to orthorhombic Pbcm space group at 30 50 GPa as previously observed and a second transition at 65 GPa to an orthorhombic Cmcm space group. They showed that the two transitions were caused by spin transitions in iron 51 52 from high spin to intermediate spin at 30 GPa and then to low spin at 65 GPa. The second transition at 65 GPa is associated with an abrupt decrease in the magnetic moment of iron in 53 one site of orthorhombic structure changing from intermediate to low spin state. Another ab 54 55 initio study concluded that Fe_3O_4 had no spin transition in iron up to 45 GPa (Bengston et al., 2013). An experimental study from Xu et al. (2004) showed an incipient metallic behavior in 56 Fe_3O_4 coupled with a magnetic moment collapse above 50 GPa, Fe_3O_4 becoming 57 nonmagnetic after 70 GPa. 58

In this study, we obtained X-ray diffraction (XRD) patterns on Fe_3O_4 at pressures from 11 GPa to 103 GPa. We observed the formation of h-Fe₃O₄ and determined its equation of state up to 70 GPa. After 70 GPa, we observed a structural transition to another orthorhombic structure with Pnma space group.

63

64 EXPERIMENTAL METHODS

We used Fe₃O₄ magnetite from Alfa Aesar (99.95% purity) as the starting material. Approximately 7 wt% of high purity gold powder was mixed with magnetite powder and gold served as the pressure calibrant. One diamond anvil cells (DAC) was loaded with the mixture in argon pressure transmitting medium (run #1) and a second DAC, in neon pressure medium (run #2). After performing the first two experiments in which a structure transition was observed, we loaded a third DAC with the mixture in neon medium to collected additional

11/4

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5409

diffraction data around the transition pressure. Anvils with 300 μ m culets and beveled anvils with 200 μ m culets were used. Re gasket were pre-indented to a thickness of 25-30 μ m and then drilled sample holes with diameters of 130 μ m to 200 μ m depending on the culet size.

In situ XRD measurements were carried out at the GSECARS (13IDD) sector of the 74 75 Advanced Photon Source (Argonne National Laboratory) with a fixed wavelength of 0.3344 Å. High temperature was achieved by double-sided laser-heating (Shen et al. 2001; 76 77 Prakapenka et al. 2008). The 20 µm laser heating spot with relatively uniform temperature 78 was carefully aligned with a 6-µm X-ray spot to obtain diffraction data at simultaneous high 79 pressure and temperature. Temperature measurements were acquired from both sides during 80 each XRD acquisition. In general, we first compressed the sample to a targeted pressure at 81 room temperature and then laser-heated sample to high temperatures. In situ XRD patterns were collected from 11 GPa to 103 GPa, at room temperature and high temperature up to 82 2300 K. Only room-temperature data obtained after annealing were presented in this study. 83

84 Pressures were calculated from the equation of state of gold (Fei et al., 2007a). Quasi-85 hydrostatic environment was achieved with neon pressure medium (Meng et al., 1993). We 86 obtained 17 XRD patterns at room temperature, integrated from two-dimensional diffraction patterns with the Fit2d program (Hammersley et al., 1996). Rietveld refinements of the in situ 87 XRD patterns were performed to determine the volumes of present phases with the GSAS 88 89 software package (Larson and Von Dreele, 1994) and the EXPGUI interface (Toby, 2001). After convergence, the values of R_{wp} were always lower than 0.03. The crystal structure of the 90 phases, atomic positions, and thermal parameters from the literature were used as the fixed 91 92 parameters to refine the unit cell parameters.

93

94 **RESULTS AND DISCUSSION**

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5409

We have conducted three experiments and obtained XRD spectra of Fe_3O_4 up to 103 95 96 GPa. During the first experiment, we collected diffraction data from 11 to 26 GPa (run #1). 97 Both magnetite and h-Fe₃O₄ were present in these patterns. For the second experiment, we 98 directly compressed Fe_3O_4 to 45 GPa, and then heated the sample to high temperatures using 99 the double-sided laser-heating system at the beamline. The sample pressures were gradually increased to 103 GPa (run #2). At each pressure, we carried out heating cycles to anneal the 100 sample and monitor the change of the XRD patterns. We observed the presence of h-Fe₃O₄ 101 only from 45 to 64 GPa. During the heating cycles, we did not observe any new diffraction 102 peaks besides those of h-Fe₃O₄, indicating no structure change upon heating. At pressures 103 higher than 64 GPa, several new diffraction peaks appeared, caused by a transition from h-104 Fe_3O_4 to a new high-pressure phase. We attempt to constrain this transition with a third 105 experiment (run #3), where Fe_3O_4 sample was directly compressed at 64 GPa and then at 76 106 107 GPa, using the same experimental procedure. We used neon as the pressure calibrant (Fei et al., 2007a) for run #3 because gold diffraction peaks were not observed. 108

109

Equation of state of high-pressure Fe₃O₄

110 First, we check the structure of the high-pressure phase of magnetite in our patterns 111 because the space group assignment is still debated. Fei et al. (1999) first suggested a Pbcm 112 space group. Haavik et al. (2000) indicated that the powder diffraction data are more consistent with a Cmcm space group assignment, whereas ab initio calculation of Ju et al. 113 114 (2012) showed that a structure with a Pbcm space group is stable. We tested the $CaTi_2O_4$ -type 115 structure (Cmcm) with the atomic position from Haavik et al. (2000) and a CaMn₂O₄-type structure (Pbcm) with atomic position from Fei et al. (1999). The CaTi₂O₄-type structure give 116 a very subtle better fit compare to the CaMn₂O₄-type as also observed by Haavik et al. (2000). 117 118 We also evaluated the atomic positions of the Cmcm structure type given by Dubrovinsky et al. (2003) for h-Fe₃O₄ and by Yamanaka et al. (2009) for Fe₂TiO₄. The fitting results are very 119

similar, which do not permit us to distinct between these atomic positions. Because Pnma is a 120 subgroup of Cmcm and of Pbcm (Hahn, 1983), we also evaluated if the Pnma space group 121 would yield a better fit. For Pnma space group, we used atomic positions of Cmcm structure 122 123 given by Haavik et al. (2000) and transformed it to a Pnma space group (Stokes and Hatch, 124 1988). The refinements by using the three space groups produced similar fitting results to the 125 observed XRD patterns. The fitting parameters are listed in Table 1 for the XRD pattern 126 obtained at 64 GPa in run #2 with no gold. Because Pnma is the lower symmetry space group, 127 it is expected that the error for Pnma is slightly smaller than that for Pbcm after refining the atomic positions. Although the refinement with a Pbcm showed a slightly smaller error than 128 the Cmcm refinement, the favored space group assignment is Cmcm because of its higher 129 130 symmetry space group. The same conclusion has been made by Sun et al. (2009) on AlH_3 where they showed that space groups R3c, Pbcm and Pnma gave the same satisfactory 131 Rietveld refinement results and the structure with the highest symmetry space group was 132 133 favored.

We choose to fit our XRD patterns by a Rietveld analysis method using the Cmcm 134 structure (i.e., the higher symmetry space group) with the atomic positions from Haavik et al. 135 136 (2000) for consistency since not all of our XRD patterns permit us to refine the atomic positions. The refined unit cell parameters for $h-Fe_3O_4$ are listed in Table 2. Figure 1 shows 137 the volume of magnetite and h-Fe₃O₄ up to 70 GPa obtained in this study compared to the 138 literature data. At pressures between 11 GPa and 26 GPa, magnetite coexists with the high-139 pressure phase. Our refined volumes of magnetite are in a good agreement with those 140 obtained by Haavik et al. (2000), whereas volumes of $h-Fe_3O_4$ show smaller values (Fig. 1). 141 We also compared our volumes with data from Dubrovinsky et al. (2003), Mao et al. (1974), 142 143 Lazor et al. (2004), and Fei et al. (1999). Our results are consistent with those of Lazor et al. (2004), Mao et al. (1974), and Fei et al. (1999). The data point of Haavik et al. (2000) at 40 144

11/4

GPa refined with Cmcm structure is also consistent with our results. The difference in the 145 146 refined volumes between this study and Haavik et al. (2000) could come from fitting 147 procedure. For example, the volumes of $h-Fe_3O_4$ given in their Table 4 were obtained by using the CaMn₂O₄-type structure (Pbcm) with V= 239.8(1.6) Å³ at 40 GPa, whereas the 148 refined volume is V= 235.3(0.5) $Å^3$ with the Cmcm space group (see their Table 5) at the 149 same pressure which plots on our compression curve (Fig. 1). In addition, we heat the sample 150 up to about 2100 K and acquired patterns after heating, whereas Haavik et al. (2000) did not 151 152 heat their sample. There might be some structural distinctions between $h-Fe_3O_4$ transformed at 300 K and the annealed h-Fe₃O₄. Hazen and Navrotsky (1996) discussed the importance of 153 the volume of disordering in spinels with pressure. They mentioned a difference between 154 disordered volume and ordered volume up to 5% in oxides with cations involving mixed 155 valence and coordination. The h-Fe₃O₄ volume difference between those of Haavik et al. 156 (2000) and ours is about 2% and could be explained by cation ordering with annealing. 157

158 Lazor et al. (2004) predict that h-Fe₃O₄ becomes unstable above 50 GPa, favoring the breakdown to Fe₂O₃ and FeO. We have obtained diffraction data up to 103 GPa and did not 159 observe any breakdown. Figure 2 shows representative X-ray diffraction patterns at different 160 pressures. The patterns up to 64 GPa showed no new peaks other than the ones that belong to 161 h-Fe₃O₄ (Fig. 2). Upon heating the observed XRD patterns remain the same with just 162 systematic peak position shifts, indicating no oxidation change or breakdown reaction to form 163 Fe₂O₃ or FeO. New diffraction peaks did appear at pressures above 73 GPa which are caused 164 by a structure transition in Fe₃O₄ as discussed below. 165

We fitted the compression data of h-Fe₃O₄ to the Birch-Murnaghan equation and obtained a bulk modulus K₀ of 124 ±12 GPa and an initial volume V₀ of 287±4 Å³ with the derivative of the bulk modulus K' fixed to 4 by fitting the entire dataset up to 64 GPa. As shown in Figure 1, the calculated compression curve using these parameters does not reproduce the measured volumes well, particularly in the pressure range of 11 and 45 GPa. By examining the axial compression behavior, we found an abrupt contraction of *a*-axis at a pressure between 45 and 50 GPa (Table 2). The insert in Figure 1 shows the a/c ratio change as a function of pressure, indicating a clear discontinuity at about 50 GPa. The fit of the compression data up to 45 GPa yielded K_0 = 172±8 GPa and V_0 = 277±2 Å³. The derived bulk modulus is smaller than values reported in previous studies (Table 3). It is also smaller than the value of magnetite (K_0 = 186 GPa) (Reichmann and Jacobsen, 2004).

177 Phase transition in high pressure magnetite

At pressures above 70 GPa, we observed the appearance of new diffraction peaks in 178 the XRD patterns (Figure 2). Lazor et al. (2004) predicted the disproportion of Fe_3O_4 to FeO 179 180 and Fe_2O_3 above 50 GPa. We checked for a potential presence of FeO and ruled out this possibility based on the observed diffraction peaks which do not match the FeO diffraction 181 182 peaks at these pressures. Testing the presence of hematite is more complicated since the highpressure form of hematite is still uncertain (Ito et al., 2009; Shim et al., 2009). However, we 183 collected many diffraction patterns around the transition pressure at different temperatures 184 and observed consistent patterns of $h-Fe_3O_4$ and the new post $h-Fe_3O_4$ phase. We conclude 185 that no breakdown has taken place in Fe_3O_4 . The changes observed in the patterns must be 186 187 caused by a second structural transition in Fe₃O₄.

An ab initio study (Ju et al., 2012) proposed a structural transition in h-Fe₃O₄ at 65 GPa, which is in agreement with our observation. However, their proposed structure with a Cmcm space group does not fit the observed new diffraction patterns for the post h-Fe₃O₄ phase. We have researched possible structures of the post-spinel phases, particularly the highpressure structures obtained on different high-pressure post-spinel phases such as CaMn₂O₄, CaFe₂O₄ (Yamanaka et al., 2008) and Fe₂TiO₄ (Yamanaka et al., 2013). Yamanaka et al.

(2013) observed a transition in Fe₂TiO₄ at around 60 GPa. They proposed a Pmma space 194 195 group for the new high-pressure post-spinel phase in Fe_2TiO_4 . We used this orthorhombic structure to index the observed diffraction peaks, but failed to obtain a satisfactory solution. 196 We have further tested the post-spinel structure of $CaTi_2O_4$ and $CaFe_2O_4$ given by Yamanaka 197 et al. (2008), in which they observed a shift in atomic positions with pressure giving 198 orthorhombic cells with displacing atoms in every third layer perpendicular to the c axis. This 199 gives a three times bigger orthorhombic cell compared to the usual ones. Because $h-Fe_3O_4$ can 200 201 be fitted with a Cmcm space group, we tried the high-pressure form of CaTi₂O₄ post-spinel 202 which corresponds to a Cmcm orthorhombic cell that is three times larger than the $h-Fe_3O_4$ 203 unit cell. The Le bail method refinement gave a better fit compared to the one using the Pmma space group, but all diffraction peaks cannot be fitted with this structure. 204

Finally, we tested the high-pressure form of $CaFe_2O_4$ type structure (Yamanaka et al., 2008) which also corresponds to a three times larger cell of an orthorhombic form, but with a 2007 Pnma space group. The Le Bail refinement gives the best satisfactory result that indexes all 2018 the observed diffraction peaks. We conclude that the new post h-Fe₃O₄ phase at 70 GPa could 2029 have a Pnma space group with a spinel regular cell multiplied by 3. However, there is not 2030 enough information to constrain the atomic positions.

In the post h-Fe₃O₄ phase at pressures above 70 GPa, there are so many observed peaks and the LeBail refinement is not constrained enough and could give several volume values. To be more accurate in the volume determination, we used a Rietveld refinement with the atomic positions of Haavik et al. (2000) for a Cmcm space group transformed to a Pnma space group with an initial cell size multiplied by 3. Although it gives a high chi-value due to uncertainties in the estimated intensities, it permits us to obtain a better volume estimate of the new structure. Using this structure model, we then fitted the diffraction patterns by Le Bail

11/4

218 method. Figure 3 shows a representative refinement result at 76.5 GPa. The refined volumes219 and unit cell parameters are listed in Table 2.

We have examined the phase transition boundary by two laser-heating cycles started at 64 and 73 GPa at room temperature respectively. Figure 4 plots the stability of h-Fe₃O₄ and the new post h-Fe₃O₄ phase at high temperature and pressure. During the heating cycle at 64 GPa up to 2050 K, we observed only h-Fe₃O₄. After quenching to room temperature, we increased pressure to 73 GPa where the new post h-Fe₃O₄ phase appeared. During the heating cycle at 73 GPa up to 2050 K, only the post h-Fe₃O₄ phase was observed. The inferred Clapeyron slope of the phase boundary may be slightly positive.

 Fe_3O_4 shows complex behavior in its electronic and magnetic properties at high 227 228 pressure. Ding et al. (2008) proposed a magnetic transition attributed to a high spin to intermediate spin transition of Fe^{2+} between 12 and 16 GPa, whereas Baudelet et al. (2010) 229 argued for no electronic and magnetic abrupt transition up to 41 GPa. Very few experimental 230 studies have been focused on high-pressure behavior of Fe₃O₄. Xu et al. (2004) showed three 231 distinct ranges in the resistivity of Fe₃O₄: an increase between 20 to 40-50 GPa, a rapid 232 decrease up to 70 GPa, and a regular decrease up to 140 GPa. They suggested a gradual Fe³⁺ 233 moment collapse starting at 50 GPa and becoming nonmagnetic at 80 GPa. The observed 234 235 abrupt change in *a*-axis in this study (Fig. 1 insert) could be associated with observed change in resistivity of Fe_3O_4 at 50 GPa (Xu et al., 2004). The change in compression and resistivity 236 237 behavior after 50 GPa could be followed by a spin transition in iron, as observed in many iron oxides at high pressure (e.g., Badro et al., 1999; Fei et al., 2007b; Badro et al., 2002; Shim et 238 al., 2009; Merlini et al., 2010; Lavina et al., 2010). 239

Figure 5 plots our compression data up to 103 GPa, together with the calculated compression curves, compared with results on different spinels. The comparison may provide

insights to Fe^{2+} and Fe^{3+} spin transitions at high pressure. Fei et al. (2007b) show that the ion 242 radii of the low-spin Fe^{2+} and Mg^{2+} are similar. By studying the compression behavior of 243 (Mg, Fe)O with increasing pressure, they showed that the compression curves of (Mg, Fe)O 244 merge with that of MgO once Fe^{2+} is in the low spin state. The volumes of h-Fe₃O₄ at 52 GPa 245 and 64 GPa are plotted on the compression curve of post-spinel MgFe₂O₄ reported by 246 Andrault and Bolfan-Casanova (2001). This is consistent with Fe^{2+} in h-Fe₃O₄ being at low-247 spin state, supporting the hypothesis that Fe^{2+} in h-Fe₃O₄ undergoes a spin transition at about 248 50 GPa. 249

We also compare our data with CaFe₂O₄ compression curves for iron in high and low 250 251 spin states (Merlini et al., 2010) in which the spin transition takes place around 50 GPa. The volume difference between their two compression curves is about 6%. The transition from h-252 Fe₃O₄ to the post h-Fe₃O₄ phase results in comparable volume reduction, which may be 253 associated with the spin transition of Fe^{3+} . An ab initio study (Ju et al., 2012) showed gradual 254 spin transitions of ions in Fe₃O₄ from high spin to intermediate spin of iron at about 65 GPa. 255 Spin transition of Fe³⁺ in high-pressure iron oxides have been observed, such as, in Fe₂O₃ and 256 in CaFe₂O₄ (Badro et al., 2002; Merlini et al., 2010). It is often associated with a magnetism 257 collapsing, observed at high pressure in Fe_2O_3 (Pasternak et al., 1999) and in Fe_3O_4 (Xu et al., 258 2004). The observed volume reduction in h-Fe₃O₄ and the structure transition to the post h-259 Fe_3O_4 phase are likely associated with the spin transitions of Fe^{2+} and subsequently Fe^{3+} in the 260 high-pressure phases. Additional spectroscopic studies at high pressure are needed to gain 261 insights to the iron spin transitions in Fe₃O₄. 262

263

264 IMPLICATIONS

We have confirmed the high-pressure Fe_3O_4 phase (h-Fe_3O_4) has an orthorhombic 265 266 structure. The most probable space group for h-Fe₃O₄ is Cmcm. A least-squares fit of the compression data of h-Fe₃O₄ up to 45 GPa yields the EOS parameters of K_0 = 172±8 GPa and 267 $V_0 = 277 \pm 2$ Å³, with fixed K' = 4. We observed a volume contraction without structure change 268 at about 50 GPa, which is likely caused by a spin transition of Fe²⁺ in h-Fe₃O₄. This 269 suggestion is also supported by the observed changes in the magnetic behavior of Fe_3O_4 at 50 270 GPa (Xu et al., 2004) and the DFT calculations (Siberchicot, 2013; Ju et al., 2012). In 271 addition, we observed a new structural transition in Fe₃O₄ at a pressure between 64 and 73 272 GPa. This new post h-Fe₃O₄ phase has an orthorhombic structure with a likely Pnma space 273 group and the Clapeyron slope of the phase boundary between h-Fe₃O₄ and the post h-Fe₃O₄ 274 phase is positive. 275

276 Knowledge of compression behavior, spin state of iron ions, and structure transitions of iron oxides at high pressure and temperature is essential for understanding the role of iron 277 in the Earth's mantle. The bulk iron content in the Earth's mantle is about 8 wt%, distributed 278 between ferropericlase and mantle silicates in the forms of Fe^{2+} and Fe^{3+} . Iron spin transitions 279 in ferropericlase and Fe-bearing bridgmanite have been extensively studied (e.g., Badro et al., 280 2005; Fei et al., 2007b; Li et al. 2004; Lin et al. 2013; Jackson et al., 2005). The observed 281 transitions in Fe_3O_4 provide further understanding of spin transitions of Fe^{2+} and Fe^{3+} at high 282 pressure and their effect on the density and element partitioning in lower mantle minerals. 283

Fe₃O₄ contains both Fe²⁺ and Fe³⁺ with different spin states at high pressure. As demonstrated, the electronic and structural transitions could significantly change the compression behavior of Fe₃O₄. As an important component in the iron oxide buffer system, accurate equation of state of Fe₃O₄ over a large pressure and temperature range is essential for calibrating the buffers involving Fe₃O₄ under mantle conditions. Campbell et al. (2009) showed the effect of pressure on the iron-wüstite oxygen fugacity buffer at high pressure. With the measured equations of state of different Fe_3O_4 phases, we can now calculate the FeO-Fe₃O₄ buffer as a function of pressure. The new structural transition shows a volume drop of about 6 % relative to the compression curve of h-Fe₃O₄. This volume change could have a strong impact on the buffer involving Fe₃O₄. Further experiments are required to better constrain the space group and atomic positions, and its equation of state.

295

296 Acknowledgements

This work was supported by NSF geophysics grant to YF and by the Carnegie Institution ofWashington. We thank V. Prakapenka and P. Dera for technical assistance and C. Seagle for

collecting part of the experimental data. The X-ray diffraction data were collected at APS

300 GSECARS beamline supported by NSF, DOE, and the State of Illinois.

301

302 **References**

Andrault, D., and Bolfan-Casanova, N. (2001) High-pressure phase transformations in the
 MgFe₂O₄ and Fe₂O₃-MgSiO₃ systems. Physics and Chemistry of Minerals, 28, 211-217.

Badro, J., Struzhkin, V.V., Shu, J., Hemley, R.J., and Mao, H-K. (1999) Magnetism in FeO at
megabar pressures from X-ray emission spectroscopy. Physical Review Letters, 83, 20,
4101-4104.

- Badro, J., Fiquet, G., Struzhkin, V.V., Somayazulu, M., Mao, H-K., Shen, G., and Le Bihan,
 T. (2002) Nature of the High-pressure transition in Fe₂O₃ hematite. Physical Review
- 310 Letters, 89, 20, doi: 10.1103/PhysRevLett.89.205504.

- Badro, J., Figuet, G., and Guyot, F. (2005) Thermochemical state of the lower mantle: new
- insights from mineral physics. Earth's Deep Mantle: Structure, Composition, and
 Evolution, Geophysical Monograph Series 160, p 241-260.
- Baudelet, F., Pascarelli, S., Mathon, O., Itié, J-P., Polian, A., and Chervin, J-C. (2010)
- Absence of abrupt pressure-induced magnetic transitions in magnetite. Physical Review B,
 82, 140412.
- Bengston, A., Morgan, D., and Becker, U. (2013) Spin state of iron in Fe₃O₄ magnetite and hFe₃O₄. Physical review B, 87, 155141.
- 319 Campbell, A.J., Danielson, L., Righter, K., Seagle, C.T., Wang, Y., and Prakapenka, V.B.
- (2009) High pressure effects on the iron-iron oxide and nickel-nickel oxide oxygen
 fugacity buffers. Earth and Planetary Science Letters, 286, 556-564.
- 322 Ding, Y., Haskel, D., Ovchinnikov, S.G., Tseng, Y-T., Orlov, Y.S., Lang, J.C., and Mao, H-
- K. (2008) Novel pressure-induced magnetic transition in magnetite (Fe3O4). Physical
 Review Letters, 100, 045508.
- 325 Dubrovinsky, L.S., Dubrovinskaia, N.A., McCammon, C., Rozenberg, G.Kh., Ahuja, R.,
- 326 Osorio-Guillen, J.M., Dmitriev, V., Weber, H.P., LeBihan, T., and Johanson, B. (2003)
- The structure of the metallic high-pressure Fe₃O₄ polymorph: experimental and theoretical
 study. Journal of Physic: Condensed Matter, 15, 7697-7706.
- Fei, Y., Frost, D.J., Mao, H-K., Prewitt, C., and Hausermann, D. (1999) In situ structure
 determination of the high-pressure phase of Fe₃O₄. American Mineralogist, 84, 203-206.
- 331 Fei, Y., Ricolleau, A., Frank, M., Mibe, K., Shen, G., and Prakapenka, V. (2007a) Toward an
- internally consistent pressure scale. Proceeding of the National Academy of Sciences of
- the United States of America, doi:10.1073/pnas.0609013104.

- Fei, Y., Zhang, L., Corgne, A., Watson, H., Ricolleau, A., Meng, Y., and Prakapenka, V.
- 335 (2007b) Spin transition and equations of state of (Mg, Fe)O solid solutions. Geophysical
- Research Letters, 34, L17307, doi:10.1029/2007GL030712.
- Fleet, M.E. (1981) The structure of magnetite. Acta Crystallographica section B-Structural
 science, 37, 917-920, doi:10.1107/S0567740881004597.
- Haavik, C., Stolen, S., Fjellvag, H., Hanfland, M., and Häusermann, D. (2000) Equation of
- 340 state of magnetite and its high-pressure modification: Thermodynamics of the Fe-O system
- at high pressure. American Mineralogist, 85, 514-523.
- Hahn, T.D. (1983) International Tables for Crystallography, volume A: Space Group
 Symmetry. Reidel Publishing Company, Dordrecht.
- Hammersley, A.P., Svensson, S.O., Hanfland, M., Fitch, A.N., and Hausermann, D. (1996)
- Two-dimmensional detector software: From real detector to idealized image or two-theta scan. High pressure Research, 14, 235-245, doi:10.1016/S0012-821X(03)00361-3.
- Hazen, R.M., and Navrotsky, A. (1996) Effects of pressure on order-disorder reactions.
 American Mineralogist, 81, 1021-1035.
- 349 Ito, E., Fukui, H., Katsura, T., Yamazaki, D., Yoshino, T., Aizawa, Y., Kubo, A., Yokoshi, S.,
- Kawabe, K., Zhai, S., Shatzkiy, A., Okube, M., Nozawa, A., and Funakoshi, K-I. (2009)
- 351 Determination of high-pressure phase equilibria of Fe_2O_3 using the Kawai-type apparatus
- equipped with sintered diamond anvils. American Mineralogist, 94, 205-209, doi:
- 353 10.2138/am.2009.2913.
- Jackson, J.M., Sturhahn, W., Shen, G., Zhao, J., Hu, M.Y., Errandonea, D., Bass, J.D., and
- Fei, Y. (2005) A synchrotron Mösbauer spectroscopy study of (Mg,Fe)SiO₃ perovskite up
- to 120 GPa. American Mineralogist, 90, 199-205.

- Ju, S., Cai, T.Y., Lu, H.S., and Gong, C.D. (2012) Pressure-induced crystal structure and
- spin-state transitions in magnetite (Fe_3O_4). Journal of the American Chemical society, 134,
- 359 13780-13786, doi: 10.1021/ja305167h.
- 360 Knittle, E., and Jeanloz, R. (1986) High-pressure metallization of FeO and implications for
- the Earth's core. Geophysical Research Letters, 13, 1541-1544.
- Larson, A. C., and R. B. Von Dreele (2000) General structure analysis system (GSAS), Los
 Alamos National Laboratory Report LAUR, 86-748.
- Lavina, B., Dera, P., Downs, R.T., Yang, W., Sinogeikin, S., Meng, Y., Shen, G., and
- Schiferl, D. (2010) Structure of siderite FeCO3 to 56 GPa and hysteresis of its spin-pairing
 transition. Physical Review B, 82, 064110.
- Lazor, P., Shebanova, O.N., and Annersten, H. (2004) High-pressure study of stability of
 magnetite by thermodynamic analysis and synchrotron X-ray diffraction. Journal of
 Geophysical Research, 109, B05201, doi:10.1029/2003JB002600.
- Li, J., Struzhkin, V.V., Mao, H.K., Shu, J., Hemley, R.J., Fei, Y., Mysen, B., Dera, P.,
- Prakapenka, V., and Shen, G. (2004) Electronic spin state of iron in lower mantle
- perovskite. Proceeding of the National Academy of Sciences of the United States of
 America, 101, 14027–14030.
- Lin, J. F., Speziale, S., Mao, Z., and Marquardt, H. (2013) Effects of the electronic spin transitions of iron in lower-mantle minerals: implications to deep-mantle geophysics and geochemistry. Reviews of Geophysics, 51, 244–275.
- Mao, H-K., Takahashi, T., Bassett, W.A., Kinsland, G.L., and Merrill, L. (1974) Isothermal
 compression of magnetite to 320 kbar and pressure-induced phase transformation. Journal
- of geophysical research, 79, 8, 1165-1170.

- Meng, Y., Weidner, D.J., and Fei, Y. (1993) Deviatoric stress in a quasi-hydrostatic diamond
 anvil cell: Effect on the volume-based pressure calibration. Geophysical Research Letter,
 20, 1147-1150.
- 383 Merlini, M., Hanfland, M., Gemmi, M., Huotari, S., Simonelli, L., and Strobel, P. (2010) Fe³⁺

spin transition in $CaFe_2O_4$ at high pressure. American Mineralogist, 95, 200-213.

- 385 Murakami, M., Hirose, K., Ono, S., Tshuchiya, T., Isshiki, M., and Watanuki, T. (2004) High
- pressure and high temperature phase transitions of FeO. Physics of the Earth and Planetary
 Interiors, 146, 273-282.
- 388 Ono, S., Hirose, K., Kikegawa, T., and Saito, Y. (2002) The compressibility of a natural
- composition calcium ferrite-type aluminous phase to 70 GPa. Physics of the Earth andPlanetary Interiors, 131, 311-318.
- Pasternak, M.P., Rozenberg, G.Kh., Machavariani, G.Yu., Naaman, O., Taylor, R.D., and
 Jeanloz, R. (1999) Breakdown of the Mott-Hubbard state in Fe₂O₃: a first-order insulatormetal transition with collapse of magnetism at 50 GPa. Physical Review Letters, 82, 46634666.
- 395 Prakapenka, V. B., Kubo, A., Kuznetsov, A., Laskin, A., Shkurikhin, O., Dera, P., Rivers,
- 396 M.L., and Sutton, S.R. (2008) Advanced flat top laser heating system for high pressure
- research at GSECARS: application to the melting behavior of germanium. High Pressure
 Research, 28, 225-235.
- Reichmann, H.J., and Jacobsen, S.D. (2004) High-pressure elasticity of a natural magnetite
 crystal. American Mineralogist, 89, 1061-1066.
- 401 Schollenbruch, K., Woodland, A.B., Frost, D.J., Wang, Y., Sanehira, T., and Lagenhorst, F.
- 402 (2011) In situ determination of the spinel-post spinel transition in Fe_3O_4 at high pressure
- and high temperature by synchrotron X-ray diffraction. American Mineralogist, 96, 820-
- 404 827.

- Shen, G., Rivers, M.L., Wang, Y., and Sutton, S. R. (2001) A laser heated diamond cell
 system at the Advanced Photon Source for in situ x-ray measurements at high pressure and
 temperature. Review of Scientific Instruments, 72, 1273-1282.
- 408 Sherman, D.M. (1989) The nature of the pressure-induced metallization of FeO and its
- implications to the core-mantle boundary. Geophysical Research Letters, 16, 515-518.
- 410 Shim, S-H., Bengtson, A., Morgan, D., Sturhahn, W., Catalli, K., Zhao, J., Lerche, M., and
- 411 Prakapenka, V. (2009) Electronic and magnetic structures of the postperovskite-type Fe₂O₃
- and implications for planetary magnetic records and deep interiors. Proceeding of the
- 413 National Academy of Sciences of the United States of America, 106, 14, 5508-5512, doi:
- 414 10.1073/pnas.0808549106.
- Siberchicot, B. (2013) On the optical properties and low pressure-induced spin transition in
 magnetite from ab initio calculations. Journal of magnetism and magnetic materials, 335,
 86-89.
- 418 Stokes, H.T., Hatch, D.M. (1988) Isotropy subgroup of the 230 crystallographic space groups.

419 World Scientific Publishing Co. Pte. Ltd., p 9, p 1-56.

- 420 Sun, S., Ke, X., Chen, C., and Tanaka, I. (2009) First-principles prediction of low-energy
- 421 structures for AlH₃. Physical Review B, 79, 024104.
- Toby, B.H. (2001) EXPGUI, a graphical user interface for GSAS. Journal of Applied
 Crystallography, 34, 210-213.
- 424 Vanpeteghem, C.B., Angel, R.J., Ross, N.L., Jacobsen, S.D., Dobson, D.P., Litasov, K.D.,
- 425 and Ohtani, E. (2006) Al, Fe substitution in the MgSiO₃ perovskite structure: A single-
- 426 crystal X-ray diffraction study. Physics of the Earth and Planetary Interiors, 155, 96-103.
- 427 Xu, W.M., Machavariani, G.Y., Rozenberg, G.K., and Pasternak, M.P. (2004) Mössbauer and
- resistivity studies of the magnetic and electronic properties of the high-pressure phase of
- 429 Fe₃O₄. Physical Review B, 70, 174106.

This is a preprint, the final version is subject to change, of the American Mineralogist (MSA) Cite as Authors (Year) Title. American Mineralogist, in press. (DOI will not work until issue is live.) DOI: http://dx.doi.org/10.2138/am-2016-5409

430	Yamanaka, T., Uchida, A., and Nakamoto, Y. (2008) Structural transition of post-spinel
431	phases CaMn ₂ O ₄ , CaFe ₂ O ₄ , and CaTi ₂ O ₄ under high pressures up to 80 GPa. American
432	Mineralogist, 93, 1874-1881.
433	Yamanaka, T., Mine, T., Asogawa, S., and Nakamoto, Y. (2009) Jahn-Teller transition of
434	Fe2TiO4 observed by maximum entropy method at high pressure and low temperature.
435	Physical Review B, 80, 134120.
436	Yamanaka, T., Kyono, A., Nakamoto, Y., Meng, Y., Kharlamova, S., Struzhkin, V.V., and
437	Mao, H.K. (2013) High-pressure phase transitions of Fe_{3-x} Ti _x O ₄ solid solution up to 60
438	GPa correlated with electronic spin transition. American Mineralogist, 98, 736-744.
439	
440	
441	
442	Figures captions
443	
444	Figure 1: volume of magnetite and h-Fe ₃ O ₄ obtained at high pressure (using gold pressure
445	scale) and room temperature in this study are shown together with data from the literature.
446	Compression curves obtained using dataset up to 45 GPa (1) and 64 GPa (2) and obtained by
447	Haavik et al. (2000) on h-Fe ₃ O ₄ are shown for comparison. The inserted figure shows the
448	change of a/c ratio as a function of pressure.
449	Figure 2: XRD patterns of magnetite with pressure showing h-Fe ₃ O ₄ diffraction peaks and the
450	new structure after 70 GPa. Black and grey tick marks correspond to Neon and gold
451	diffraction peaks respectively.
452	Figure 3: XRD pattern measured at 76.5 GPa (Neon pressure) fitted by Le Bail method with a
453	Pnma space group obtained from the Cmcm unit cell of Haavik et al. (2000) transformed to

454 Pnma and multiplied by three. Cell parameters are a = 8.851(1) Å, b = 7.8370(6) Å and c = 8.660(1) Å.

456 Figure 4: Presence of h-Fe₃O₄ or the new structure with pressure and temperature as observed
457 in our XRD patterns.

- Figure 5: Volume of magnetite and its high pressure forms as a function of pressure plotted
 with compression curves obtained in this study using data up to 45 GPa. For comparison,
 compression curves from several spinel has been plotted; CF type from MORB from Ono et
- 461 al. (2002); CaFe₂O₄ LS and HS, low spin and high spin EOS parameters from Merlini et al.
- 462 (2010); MgFe₂O₄, compression curve from Andrault and Bolfan-Casanova (2001).

- 464
- 465
- 466
- 467 468
- 469
- 470
- 471
- 472
- 473
- 474
- 475
- -
- 476

477 Table 1: Result of h-Fe₃O₄ cell parameters and atomic positions fitted with three

478 orthorhombic space groups with the XRD pattern at 64 GPa.

Cmcm		Pbcm				Pnma						
Rwp = 0.0052 Reduced $\chi 2 = 0.2976E-01$ a = $2.63880(33)$ b = $8.88490(142)$ c = $9.05426(99)$			Rwp = Reduc a = 2. b = 8. c = 9.0	Rwp = 0.0050 Reduced $\chi 2 = 0.2829E-01$ a = 2.63886(33) b = 8.88442(139) c = 9.05423(100)			Rwp = Redu a = 9. b = 2. c = 8.	Rwp = 0.0048 Reduced χ^2 = 0.2581E-01 a = 9.05338(99) b = 2.63839(31) c = 8.88668(140)				
Aton	nic positio	ns		Atom	ic positions			Atom	Atomic positions			
	x	у	Z		x	у	Z		x	у	z	
Fe1	0.0	0.3830(7)	0.25	Fe1	0.782(8)	0.3675(7)	0.25	Fe1	0.0592(12)	0.75	0.8794(13)	
Fe2	0.0	0.1365(7)	0.0713(5)	Fe2	0.277(6)	0.1128(7)	0.0726(6)	Fe2	0.2416(21)	0.75	0.1346(7)	
01	0.0	0.0486(28)	0.25	01	0.718(30)	0.25	0.0	Fe3	0.9124(14)	0.75	0.6062(14)	
02	0.0	0.2430(22)	0.6014(15)	02	0.209(24)	0.2001(29)	0.25	01	0.609(7)	0.75	0.987(8)	
03	0.0	0.5	0.0	03	0.260(24)	0.4903(22)	0.1022(15)	02	0.027(5)	0.75	0.277(4)	
								03	0.257(9)	0.75	0.7978(28)	
								04	0.398(6)	0.75	0.508(9)	
	Bond	distances			Bond	distances			Bond o	listances		
	Fe1-Fe1	2.63880(33)	x2		Fe1-Fe1	2.63886(33)	x2		Fe1-Fe1	2.63839(31)	x2	
	Fe1-Fe2	2.722(7)	x2		Fe1-Fe2	2.712(7)	x2		Fe1-Fe1	2.735(20)	x2	
	Fe1-01	1.976(20)	x2		Fe1-01	2.498(6)	x2		Fe1-Fe2	2.806(18)		
	Fe1-02	2.192(14)	x4		Fe1-O2	2.12(5)			Fe1-Fe3	2.768(12)		
	Fe1-03	2.4910(25)	x2		Fe1-O2	1.87(4)			Fe1-O2	2.067(26)	x2	
					Fe1-O3	2.21(4)	x2		Fe1-O3	1.93(8)		
	Fe2-Fe2	2.63880(33)	x2		Fe1-O3	2.14(4)	x2		Fe1-04	1.79(5)	x2	
	Fe2-Fe2	2.749(11)							Fe1-O4	1.77(7)		
	Fe2-Fe2	2.734(11)	x2		Fe2-Fe2	2.63886(33)	x2					
	Fe2-01	1.797(10)			Fe2-Fe2	2.809(20)			Fe2-Fe2	2.63839(31)	x2	
	Fe2-02	1.828(13)			Fe2-Fe2	2.670(18)			Fe2-Fe3	2.640(18)		
	Fe2-02	1.721(14)	x2		Fe2-Fe2	2.769(12)			Fe2-01	2.17(5)	x2	
	Fe2-O3	1.905(4)	x2		Fe2-01	2.02(6)			Fe2-O2	2.32(6)		
					Fe2-01	1.81(5)			Fe2-O2	2.70(6)		
					Fe2-O2	1.793(12)			Fe2-O3	1.960(20)	x2	
					Fe2-O3	1.81(5)			Fe2-O4	2.14(5)	x2	
					Fe2-O3	1.66(5)						
					Fe2-O3	1.829(13)			Fe3-Fe3	2.796(25)	x2	
									Fe3-Fe3	2.63839(31)	x2	
									Fe3-01	1.70(5)	x2	
									Fe3-01	1.97(6)		
									Fe3-O2	1.768(22)	x2	
									Fe3-O3	1.64(8)		

479

480

482 Table 2. Volume and unit-cell parameters for h-Fe₃O₄ and for the new high pressure phase

P (GPa)	V (Å ³)	a (Å)	b (Å)	c (Å)	a (Å) Gold
11.0(1)	260.7(4)	9.512(17)	9.635(11)	2.845(4)	4.003(0)
18.4(2)	254.4(3)	9.405(11)	9.529(8)	2.839(3)	3.964(1)
25.7(9)	245.9(2)	9.364(8)	9.421(5)	2.787(2)	3.929(4)
26.3(9)	245.9(2)	9.355(8)	9.421(5)	2.790(2)	3.927(4)
44.7(6)	231.1(1)	9.248(3)	9.231(2)	2.707(1)	3.859(2)
44.8(4)	231.1(1)	9.246(3)	9.234(2)	2.707(1)	3.859(1)
53.9(4)	222.4(1)	9.095(4)	9.147(2)	2.674(1)	3.831(1)
51.6(1.3)	222.3(1)	9.063(3)	9.161(2)	2.678(1)	3.838(4)
53.1(5)	222.3(5)	9.054(3)	9.168(2)	2.678(1)	3.833(2)
64.2(1)	213.4(1)	8.908(4)	9.075(3)	2.639(1)	3.803(0)
64.4(1)	213.4(1)	8.909(4)	9.075(3)	2.639(1)	3.802(0)
64.3 ^(a)	212.3(0)	8.881(2)	9.055(1)	2.640(0)	3.084(1) ^(a)
73.2(1.1)	613.2(5)	8.841(6)	7.933(6)	8.743(7)	3.780(3)
76.4 ^(a)	600.7(9)	8.851(1)	7.837(1)	8.660(1)	3.038(1) ^(a)
87.2(6)	592.7(4)	8.742(6)	7.873(6)	8.612(7)	3.749(1)
88.5(2)	586.6(6)	8.671(9)	7.822(11)	8.649(14)	3.746(0)
103.1(2.0)	577.9(8)	8.517(12)	7.771(15)	8.731(18)	3.717(4)

483 with pressure obtained with gold unit-cell parameter.

484

^(a)data corresponding to the run 3 where pressure value is obtained with Neon.

485

486	Table 3.	Equation	of state	parameters	for h-	Fe ₃ O ₄	compared	with the	literature.
			01 00000		101 11		••••••••		

	V ₀ (Å ³)	K ₀ (GPa)	К'
This study (up to 64 Gpa)	287 (4)	124 (12)	4
This study (up to 46 Gpa)	277 (2)	172 (8)	4
Haavik et al. (2000)	278.24	202 (7)	4
Dubrovinsky et al. (2003)	277.9 (4)	198 (5)	4
Bengston et al. (2013)	270.7	189	4.02
Ju et al. (2012)	-	234.2 (2.5)	2.70 (09)

487

488

489



Figure 1



Figure 2



Figure 3







Figure 5