

1 **Revision 1**

2 **Thermodynamics of Bastnaesite - a Major Rare Earth Ore Mineral**

3 **Radha Shivaramaiah¹, Andre Anderko², Richard E. Riman³ and Alexandra Navrotsky^{1,*}**

4 ¹Peter A. Rock Thermochemistry Laboratory and NEAT ORU, University of California Davis,

5 Davis, CA 95616, USA

6 ²OLI Systems, Inc. 240 Cedar Knolls Road, Cedar Knolls NJ 07927, USA

7 ³Department of Materials Science and Engineering, The State University of New Jersey, Piscataway, NJ

8 08854-8066, USA

9

10

11 **Abstract**

12 Bastnaesite, [RE-CO₃-OH/F] (RE = rare earth) is one of the major sources of rare earth
13 elements found in commercial deposits at Mountain Pass, California, Bayan Obo, China, and
14 elsewhere. Synthetic forms of bastnaesite have been explored for applications including optical
15 devices and phosphors. Determination of thermodynamic properties of these phases is critical for
16 understanding their origin, mining, and processing. We report the first experimental
17 determination of formation enthalpies of several OH and F bastnaesites based on high
18 temperature oxide melt solution calorimetry of well characterized synthetic samples. The
19 formation enthalpies from binary oxides and fluorides for all the bastnaesite samples are highly
20 exothermic, consistent with their stability in the garnet zone of the Earth's crust. Fluoride
21 bastnaesite, which is more abundant in nature than its hydroxide counterpart, is
22 thermodynamically more stable. For both OH and F bastnaesite, the enthalpy of formation
23 becomes more negative with increasing ionic radius of the RE³⁺ cation. This periodic trend is
24 also observed among rare earth phosphates and several other rare earth ternary oxides. For a
25 given RE, the formation enthalpies from binary oxides are more negative for orthophosphates
26 than for bastnaesites, supporting the argument that monazite could have formed by reaction of
27 bastnaesite and apatite at high temperature. The difference in formation enthalpy of monazite and
28 bastnaesite provides insight into energetics of such reactions along the rare earth series.

29

30

31

Introduction

32 Rare earth elements (RE) are technology drivers in today's world where they are
33 integrated into every part of modern military and civilian life, including rechargeable batteries
34 and other energy applications, phosphors for lighting and computer displays, catalytic converters,
35 and magnets (McLellan et al. 2013; Pathak et al. 2015; Gschneider 1981). Their wide application
36 has created a high demand, which makes their availability a crucial issue, or, in other words, a
37 critical material (Alonso et al. 2012; Humphries 2010). REs occur in nature both as major and
38 trace constituents of minerals (Keith 2011). Monazite, bastnaesite, xenotime, aeschynite and
39 allanite are minerals in which RE is an essential constituent. Bastnaesite, first discovered in 1818
40 at Bastnaes, Sweden (Glass and Smalley 1945), constitutes an important RE ore and occurs as a
41 high grade accessory mineral of igneous or hydrothermal origin, for example in Mountain Pass
42 (California), Bayan Obo (China) and Wingu Hill (Tanzania) (Head and Holley 1963). The
43 Molycorp operation at Mountain Pass CA (Pradip 1981a) was the largest producer of rare earths
44 until the late 1990s when low cost Chinese production hit the global market and now constitutes
45 97 % of the world's RE production. Today there is some interest in the U.S. and Canada in
46 renewed RE mining.

47 Bastnaesite represents a family of RE carbonate minerals $[\text{RE}(\text{OH}/\text{F})\text{CO}_3]$ which mostly
48 contain light rare earths (LRE) - La, Ce, Nd and Pr and trace amounts of heavy rare earths
49 (HRE) - Y, Dy, Lu, Eu and others (Fleischer 1978). Its structure consists of RE-F sheets
50 separated by carbonate ions and is related to other RE fluorocarbonate minerals such as parisite
51 $[\text{CaRE}_2(\text{CO}_3)_2\text{F}_2]$, roentgenite $[\text{Ca}_2\text{RE}_3(\text{CO}_3)_5\text{F}_3]$ and synchysite $[\text{CaRE}(\text{CO}_3)_2\text{F}]$, where
52 bastnaesite - like sheets are interlayered with sheets of CaCO_3 (Vlasov 1964). As the trivalent RE
53 ions have similar crystal chemistry to Ca^{2+} , they substitute for Ca in rock forming minerals, and

54 these RE bearing minerals are active participants in metamorphic reactions on the geologic time
55 scale (Lipin and Mckay 1989). These fluorocarbonates are known to have grown from fluids
56 containing both HF and CO₂ and are potential indicators of relative abundances of CO₂ and HF
57 (Jones and Wood 1992). Being carbonates, they also can play a role in controlling the global
58 carbon cycle, although they are not major sources or sinks of carbon. While both F and OH
59 bastnaesites are found in nature, the F form is predominant (Jones and Wood 1992). Several
60 reports document bastnaesites from various locations in terms of structure, composition, phase
61 relations and reactivity with other RE accessory minerals (Exley 1980; Smith et al. 1999).

62 The first efforts to synthesize bastnaesite phases date back to 1950s (George et al. 1959).
63 However it was not until the early 1970s that these phases were studied in terms of structural
64 aspects and compositional variability. In 1973 Haschke proposed phase equilibria in the
65 lanthanum - hydroxide - fluoride - carbonate system as a function of temperature and pressure by
66 synthesizing a series of solid solutions in the La(OH)_{1-x}F_xCO₃ system (Haschke 1975; Haschke
67 and Eyring 1971). These investigations provide possible reaction paths for the geochemical
68 formation of bastnaesite. There are recent reports on the crystallization mechanism of OH
69 bastnaesite under hydrothermal conditions (Vallina et al. 2014). In recent years synthetic
70 bastnaesites have gained increased attention as they are used as precursors for synthesis of rare
71 earth oxycarbonates (RE₂O₂CO₃) and oxyfluorides (REOF), which are excellent host lattices for
72 phosphors (Lee and Jung 2013; Janka and Schleid 2009). RE hydroxycarbonates are also of
73 interest for optical technologies (Chakhmouradian and Wall 2012). Despite this renewed interest,
74 there are no reliable thermodynamic data available in the literature for bastnaesites. Limited
75 reports on the thermal stability of bastnaesite phases indicate that fluoride phase always persists
76 to higher temperatures than its hydroxide counterpart (Hsu 1992).

77 This implies that the F – phase could be stable at magmatic conditions while the OH – phase
78 exists only in low temperature environments (Hsu 1992). Based on solubility product and
79 stability constant measurements from solubility and titration experiments, a speciation diagram
80 has been proposed for the CeFCO₃ system (Pradip et al. 2013).

81 In this paper we report the enthalpy of formation for some synthetic bastnaesite phases by
82 direct measurements using high temperature oxide melt solution calorimetry. For this purpose we
83 carried out synthesis and characterization of F – bastnaesites (La and Ce) and both polymorphs
84 of OH – bastnaesites of several RE (La, Ce, Y, Dy and Nd).

85 **Experimental Methods**

86 Synthesis

87 A summary of the synthesis conditions of both F and OH – bastnaesites is given in Table
88 1. The detailed procedure used for the syntheses is described below.

89 F – bastnaesites: REFCO₃ (RE = La and Ce) phases were synthesized by urea hydrolysis
90 in the presence of NH₄F as a source of F⁻. In a typical synthesis 0.02 mol of RE(NO₃)₃ was
91 dissolved in 20 mL of deionized (DI) water and was added dropwise to a 50 mL of solution
92 containing 0.02 mol of NH₄F. The mixture was stirred for 30 min and 0.04 mol of urea was
93 added. The resulting suspension was aged at 90 °C for 20 h and the precipitate obtained was
94 separated by centrifugation, washed with DI water and dried at 60 °C.

95 OH – bastnaesites: REOHCO₃ (RE = La, Ce, Y, Nd and Dy) were synthesized by three
96 different aqueous methods, (a) urea hydrolysis and (b) base hydrolysis of RECl₃ using sodium
97 carbonate and (c) water hydrolysis of RE(CO₃)₃. In a typical synthesis by urea hydrolysis 0.02
98 mol of RECl₃ and 0.01 mol of urea were dissolved in 50 mL of DI water and stirred at room

99 temperature for 30 min. The solution was then transferred to a Teflon lined autoclave and heated
100 at 180 °C for 3 h. The precipitate was obtained after cooling and subsequently separated by
101 centrifugation, washed with DI water 3 times, and dried at 60 °C. For base hydrolysis a 50 mM
102 solution of RECl_3 was added to a 50 mM solution of Na_2CO_3 at room temperature, stirred at
103 room temperature for 30 min and transferred to a Teflon-lined autoclave, followed by
104 hydrothermal treatment at 180 °C for 48 h. After cooling the autoclave to room temperature, the
105 precipitate was separated by centrifugation, washed with DI water several times and dried in an
106 oven at 60 °C. For hydrolysis of $\text{RE}(\text{CO}_3)_3$ in water, 5g of rare earth carbonate was taken in 100
107 mL water and heated first at 90 °C for 20 h and then at 100 °C for 1 h. The product was
108 separated by centrifugation and dried at 60 °C.

109 Characterization

110 Powder X-ray diffraction patterns (XRD) of the synthesized samples were recorded using
111 a Bruker AXS D8 Advance diffractometer with CuK α radiation, $K\alpha = 1.5418 \text{ \AA}$, to identify the
112 phases. The Rietveld technique was employed for structure refinement of all the samples using
113 X'pert HighScore Plus software. All the structure refinements were carried out using the
114 published structure models. When there was an impurity phase in the sample, phase fractions
115 were obtained by two phase refinements.

116 Thermogravimetric analysis (TGA) was performed to determine the water content in the
117 samples using a Netzsch 449 thermal analysis system in a dynamic Ar atmosphere (40mL/min,
118 25 - 800 °C, 5 °C/min, Pt crucible). The data were analyzed using Netzsch Proteus software. The
119 overall RE content in the samples was measured by inductively coupled plasma mass
120 spectrometry (ICP-MS) with an Agilent Technologies 7500a instrument. A known amount of
121 sample was dissolved in ~0.1 mL of ICP grade HCl and diluted with Millipore water. Fourier

122 transform infrared attenuated total reflectance spectra (ATR - FTIR) of all the samples were
123 recorded using a Bruker Model Alpha-P IR spectrometer (diamond ATR cell, 4 cm⁻¹ resolution,
124 400 - 4000 cm⁻¹).

125 High temperature oxide melt solution calorimetry

126 High temperature oxide melt solution calorimetry was performed using a custom built
127 isoperibol Tian-Calvet microcalorimeter as described by Navrotsky (1977; 1997). Molten
128 sodium molybdate was used as a solvent at 700 °C. The calorimeter was calibrated using the
129 heat content of high purity α -Al₂O₃. Oxygen was flushed through the glassware at 60 mL/min to
130 maintain constant atmosphere and bubbled through the solvent at 30 mL/min using a bubbling
131 tube to remove evolved gases, aid dissolution, and prevent local saturation of the solvent. In a
132 typical experiment ~5 mg of loosely pelletized sample was dropped into the solvent. The
133 measured heat effect (heat of drop solution, ΔH_{ds}), includes the heat content of the sample, the
134 heat effect associated with desorbing and vaporizing water, decarbonation and the heat of
135 dissolution of the oxide component of the sample in the high temperature solvent. Heats of drop
136 solution of the binary oxides were taken from previously published work (Navrotsky 2014).

137 **Results and Discussion**

138 Two F – bastnaesites (La and Ce) were successfully synthesized by urea hydrolysis in the
139 presence of NH₄F. Among the OH – bastnaesites, two phase pure samples with hexagonal
140 structure (La and Nd) were synthesized. Four OH – bastnaesite samples with orthorhombic
141 structure were prepared (La, Ce, Dy and Y), among which the one containing Dy was phase pure
142 and the others contained oxide impurities. A summary of samples used for calorimetric
143 measurements is given in Table 1

144 Rietveld fits of the PXRD patterns of the two F – bastnaesites are shown in Figure S1 in
145 the supplementary information. Both LaFCO₃ and CeFCO₃ crystallize in hexagonal crystal
146 structure. PXRD pattern of LaFCO₃ could be refined using published crystal structure (ICSD no:
147 26678) with $a = 7.180 (8) \text{ \AA}$ and $c = 4.912 (6) \text{ \AA}$. PXRD pattern of CeFCO₃ was refined using a
148 published crystal structure (ICSD no: 27591) with $a = 7.201 (1) \text{ \AA}$ and $c = 9.970 (2) \text{ \AA}$. The
149 results of refinements of both F – bastnaesites are listed in Table 2. In the crystal structure of
150 LaFCO₃, fluoride anions are present in two crystallographically distinct sites. La³⁺ cations form
151 hexagonal layers with both the fluoride ions. Oxygen atoms of the carbonate are present in two
152 different sites as well, resulting in two topologically different carbonate anions sandwiched
153 between the hexagonal layers formed by La³⁺ and F⁻. CeFCO₃ exhibits a similar structure, but
154 with an elongated c - parameter corresponding to two hexagonal layers formed by Ce and F ions,
155 unlike in LaFCO₃ where there is one layer per unit cell. A view of crystal structure along the c -
156 direction showing the hexagonal layers formed by Ce and F ions is shown in Figure 1a, which is
157 isostructural with LaFCO₃. The structure along the b -direction, showing carbonate in two
158 different orientations between these hexagonal layers, is shown in Figure 1b. The Ce³⁺ cation is
159 in irregular coordination with eight O and three in - plane F ions, resulting in an overall
160 coordination number of 11 for RE in a distorted tricapped trigonal prismatic arrangement.

161 OH – bastnaesite, on the other hand, is known to exist in two different polymorphic
162 modifications - hexagonal and orthorhombic (Sum 1986). Our attempts to synthesize both
163 polymorphs in the Y, La, Ce, Nd and Dy were partially successful. While most of our synthetic
164 efforts produced a mixture of the two polymorphs, we found one methodology to successfully
165 synthesize single phase precipitates for each of the REOHCO₃ systems, see Table 1. Both
166 hexagonal and orthorhombic polymorphs were obtained for LaOHCO₃ and the Rietveld fit of the

167 PXRD patterns are shown in Figure S2. The hexagonal polymorph was phase pure and the
168 orthorhombic polymorph had La_2O_3 impurities. Two phase refinement was performed for the
169 orthorhombic polymorph to quantify the amount of impurity, which was found to be 6.5 %. The
170 obtained refined lattice parameters and goodness of fit parameters are listed in Table 2. The two
171 polymorphs of OH – bastnaesites are not isostructural with F – bastnaesite, although the metal
172 sub-lattice looks similar for the hexagonal polymorph. The crystal structures of both polymorphs
173 are shown in Figure 2. In both cases each La is coordinated with 11 O atoms. Distortion in
174 carbonate polyhedra is observed, probably due to the hydrogen bonding with OH^- ions. Further
175 the OH anions are not coplanar with the La cations, unlike those in LaFCO_3 .

176 PXRD patterns of CeOHCO_3 (orthorhombic), YOHCO_3 (orthorhombic), NdOHCO_3
177 (hexagonal) and DyOHCO_3 (orthorhombic) are shown in Figure S3. CeOHCO_3 and YOHCO_3
178 contained RE_2O_3 impurities, quantified to be 5.9 and 6.3 wt % respectively. NdOHCO_3 and
179 DyOHCO_3 were phase pure. The refined lattice parameters of all the OH - bastnaesites are listed
180 in Table 2.

181 Thermogravimetric (TG) curves of the two F – bastnaesites and that of OH bastnaesites
182 are shown in Figure S4. A total mass loss of 20 % and 22 % was observed up to 550 °C for
183 LaFCO_3 and CeFCO_3 respectively, which corresponds to decarbonation, leading to residual
184 metal oxyfluorides. OH – bastnaesites showed no significant mass loss below 150 °C. A single
185 step mass loss was observed for CeOHCO_3 and DyOHCO_3 , indicating that decarbonation and
186 dehydroxylation occur simultaneously. In contrast, the other OH – bastnaesites showed two-step
187 mass losses, suggesting that decarbonation and dehydroxylation occur at different temperatures.

188 Attenuated total reflectance (ATR) spectra of F and OH – bastnaesites are shown in
189 Figure S5. Both the F – bastnaesites exhibit peaks corresponding to carbonate vibrations at 1440
190 and 865 cm^{-1} , whereas all the OH – bastnaesites exhibit multiple absorption bands in the same
191 frequency range due to lowering of local symmetry of carbonate. This could be due to distortion
192 of the carbonate ion by the polarizing influence of the OH group. Two different kinds of
193 distorted carbonate groups were also seen in these OH – bastnaesites. The observed peaks could
194 be attributed to ν_1 (1070 - 1100 cm^{-1}), ν_2 (850 - 880 cm^{-1}), ν_3 (1400 - 1510 cm^{-1}) and ν_4 (690 -
195 730 cm^{-1}) modes of carbonate. In addition, two weak but discernable sharp peaks were observed
196 between 3600 and 3400 cm^{-1} due to OH vibrations. The infrared spectrum of the orthorhombic
197 polymorphs is less complicated than that of the hexagonal polymorph.

198 High temperature oxide melt solution calorimetry

199 Enthalpies of drop solution of all the samples and the constituents used in thermodynamic
200 cycle are listed in Table 3. Enthalpies of formation of F – bastnaesites were calculated from a
201 mixture of constituent oxides and fluorides and the thermochemical cycle used for the
202 calculation is shown in Table 4. Since oxygen was bubbled through and over the solvent during
203 the experiments, the dissolved state of cerium is Ce^{4+} . The calculated enthalpy of formation of
204 LaFCO_3 and CeFCO_3 were -173.07 ± 2.40 and -141.24 ± 2.58 kJ/mol, respectively. The strongly
205 exothermic values indicate that these F bastnaesites are significantly stable compared to
206 mechanical mixture of their oxide and fluoride components and CO_2 gas.

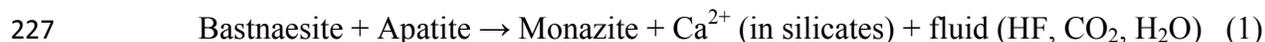
207 Thermodynamic cycles used for the calculation of formation enthalpies of OH –
208 bastnaesites from their constituent oxides are shown in Table 5. Enthalpy of formation of the two
209 hexagonal polymorphs LaOHCO_3 and NdOHCO_3 are -196.55 ± 1.64 and -165.01 ± 3.05 kJ/mol,
210 respectively.

211 Enthalpies of formation of the four orthorhombic polymorphs are -227.38 ± 2.87 , -166.03
212 ± 3.35 , -133.72 ± 1.90 and -105.60 ± 3.82 kJ/mol for La, Ce, Dy and Y OH – bastnaesites,
213 respectively. A plot of RE³⁺ cation radius (Shannon and Prewitt 1969) vs. enthalpy of formation
214 is shown in Figure 3.

215 **Discussion**

216 As seen from Figure 3, the formation enthalpies of the bastnaesite phases from RE₂O₃,
217 CO₂, and H₂O or HF become increasingly negative with increase in ionic radius of the RE³⁺
218 cation. A similar general trend has been observed among several other ternary compounds of
219 rare earths for a given structure, including phosphates, perovskites, aluminates and silicates
220 (Navrotsky 2001 and Ushakov et al. 2001).

221 Among RE mineral phases, bastnaesite is the lowest temperature RE mineral and is stable
222 within the garnet zone (Savko and Bazikov 2011). One of the hypotheses is that the bastnaesites
223 could have been formed during metamorphism (Savko and Bazikov 2011). Quite similar to
224 monazite, bastnaesites show light rare earth (LRE) selectivity and have a coordination number of
225 11 for RE cations. As temperature increases, bastnaesite is known to react with apatite to form
226 monazite (Savko and Bazikov 2011 and Smith et al 1999), as represented by reaction (1).

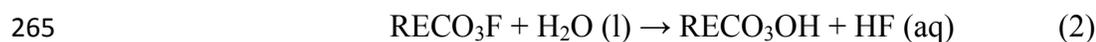


228 To consider this reaction for different bastnaesites, we compare the enthalpy of formation of
229 bastnaesites and RE phosphates. The trends in variation of formation enthalpy of monazite and
230 bastnaesite for different RE are similar, see Figure 4a, where with increasing ionic radius of
231 RE³⁺, the formation enthalpy becomes more negative (this work and Ushakov et al. 2001). While
232 the observed trends are similar, it is obvious that formation enthalpies from binary oxides are

233 more negative for phosphates than for bastnaesites. Furthermore the difference in formation
234 enthalpies of monazite and bastnaesite ($\Delta H_{f,diff} = \Delta H_{f,monazite} - \Delta H_{f,bastnaesite}$) for different RE gives
235 an indication of the variation of enthalpy and free energy with RE for reaction (1), assuming the
236 apatite, Ca-silicate, and fluid phases remain similar. A plot of difference in formation enthalpies
237 of RE monazite and RE bastnaesite is shown in Figure 4b. The difference is least negative for
238 both the F – bastnaesites, LaFCO₃ and CeFCO₃. The difference is comparatively more negative
239 for OH – bastnaesites, with La being least and Y being most negative. The difference thus seems
240 to become less negative with increasing radius of RE³⁺, consistent with the LRE preference of
241 bastnaesite. The least negative difference, observed for F – bastnaesite, both in case of La and
242 Ce, supports the observation that natural bastnaesites are strongly Ce and La - selective. Ce is the
243 element which is most abundant in natural bastnaesites and La is the second. In some monazite
244 deposits, the presence of F also indicates that monazite was formed at the expense of bastnaesite
245 and apatite (Savko and Bazikov 2011; Smith et al 1999) which is also supported by their
246 energetic stabilities.

247 In nature F – bastnaesite is dominant and OH – bastnaesite is rare. OH – bastnaesites are
248 known to occur only in low temperature zones which are essentially devoid of F. The stability
249 relations at $P_F = 1$ kbar and $T = 400 - 900$ °C in a part of (Ce, La)-F-H-C-O system indicate that
250 the F – bastnaesites of both La and Ce have stability fields wider than their OH – counterparts
251 by 50 °C and 120 °C respectively (Hsu 1992). This finding is also supported by Haschke (1975).
252 From the only two synthetic analogues available for comparison, it can be seen from Table 3 that
253 the enthalpy of formation calculated from binary oxides of the OH – bastnaesites of both La and
254 Ce are more negative than those of their F – counterparts calculated from a mixture of binary
255 oxide and fluoride end members. These values do not help for direct comparison, since they were

256 calculated from different end members. We thus compare the enthalpy of formation values
257 calculated from elements for these phases as shown in Table 4 and 5. The enthalpies of
258 formation calculated from elements indicate that the F – bastnaesites are more stable than their
259 OH – counterparts (last column of Table 3). Also, the enthalpy of reaction calculated for the
260 conversion of RECO₃F to RECO₃OH, which corresponds to substitution of F⁻ by OH⁻, is positive
261 (Eq. 2) for both La and Ce (100.59 and 69.76 kJ/mol for La containing hexagonal and
262 orthorhombic OH - polymorphs respectively and 106.49 kJ/mol for Ce). These positive values
263 strongly suggest that the reaction is not spontaneous, since entropy effects are expected to be
264 relatively small.



266 Thus the F – bastnaesites are thermodynamically more stable than their OH – analogues.
267 This observation explains why F – bastnaesite is more common in the Earth’s crust than OH –
268 bastnaesite and also explains the ease of formation of the former in laboratory conditions even
269 under low F content in the aqueous phase.

270 **Implications**

271 While in reality bastnaesite is a complex solid solution with different RE and in some
272 cases with both OH⁻ and F⁻, in this paper we have reported the thermodynamic properties of F⁻
273 and OH⁻ end members. The new data should be useful for modelling the formation of the mineral
274 and leaching and processing of the ore. The calorimetric data confirm the increasing stability of
275 bastnaesites with light rare earths and the greater stability and persistence to higher temperature
276 of the fluoride end members compared to the hydroxides, thus providing a thermodynamic

277 justification for the pattern of bastnaesite occurrences in nature. The data provide a starting point
278 for the modelling of complex bastnaesite solid solutions.

279 **Acknowledgments**

280 The first steps of this study, including synthesis and characterization, were supported by
281 the National Science Foundation, Division of Earth Sciences with contributions from the NSF -
282 wide SUSChem initiative (grant EAR - 1321410). The latter part of the study, including
283 calorimetry, data analysis, and manuscript preparation, was supported by the Critical Materials
284 Institute, an Energy Innovation Hub funded by the U.S. Department of Energy, Office of Energy
285 Efficiency and Renewable Energy, Advanced Manufacturing Office.

286

287

References

- 288 Alonso, E., Sherman, A.M., Wallington T.J., Everson M.P., Field F.,R, Roth R., Kirchain, R E.
289 (2012) Evaluating rare earth element availability: a case with revolutionary demand from
290 clean technologies. *Environmental Science and Technology*, 46, 3406–341
- 291 Chakhmouradian, A.R., Wall, F. (2012) Rare earth elements: minerals, mines, magnets.
292 *Elements*, 8, 333-340
- 293 Donnay, G., and Donnay, J.D.H. (1953) The crystallography of bastnaesite, parisite, roentgenite,
294 and synchisite. *American Mineralogist*, 38, 932-938
- 295 Exley, R. A. (1980) Microprobe studies of REE-rich accessory minerals: Implications for Skye
296 granite petrogenesis and REE mobility in hydrothermal systems. *Earth and Planetary
297 Science Letters*, **1980**, 48, 97-110
- 298 Fleischer, M. (1978) Relative proportions of the lanthanides in minerals of the bastnaesite group.
299 *Canadian Mineralogist*, 16, 361-363
- 300 Glass, J. J., Smalley, R. G. (1945) Bastnaesite. *American Mineralogist*, 30, 601-615
- 301 Gschneidner, K. A. (1981) Industrial Applications of Rare Earth Elements, ACS Symposium
302 Series, 164, 1-297
- 303 Haschke, J. M. (1975) The lanthanum hydroxide fluoride carbonate system: The preparation of
304 synthetic bastnaesite. *Journal of Solid State Chemistry*, 12, 115-121
- 305 Haschke, J. M., Eyring, L. (1971) Hydrothermal equilibriums and crystal growth of rare earth
306 oxides, hydroxides, hydroxynitrates, and hydroxycarbonates. *Inorganic Chemistry*, 10,
307 2267-2274
- 308 Head, E.L., Holley, C.E. (1963) Rare-earth carbonate, *Rare Earth Research Proceedings of Third
309 Conference*, Gordon and Breach, 51-63

- 310 Hsu, L. C. (1992) Synthesis and stability of bastnaesites in a part of the system (Ce,La)-F-H-C-
311 O. Mineralogy and Petrology, 47, 87-101
- 312 Humphries, M. (2010) Rare Earth Elements: The Global Supply Chain: Congressional Research
313 Service; The Library of Congress: Washington, DC, USA.
- 314 Janka, O., Schleid, T. (2009) Facile Synthesis of Bastnaesite-Type LaF[CO₃] and Its Thermal
315 Decomposition to LaOF for Bulk and Eu³⁺-Doped Samples. European Journal of
316 Inorganic Chemistry, 3, 357-362.
- 317 Jansen, G. J., Magin, G. B., Levin, B. Synthesis of bastnaesite (1959) American
318 Mineralogist, 44, 180-190
- 319 Vlasov, K. A. (1964) Geochemistry, mineralogy and genetic types of deposits of rare elements.
320 Mineralogy of Rare Elements, 2, 264
- 321 Keith, R. L. (2011) Principal Rare Earth Elements Deposits of the U. S.: A Summary of
322 Domestic Deposits and a Global Perspective, 1-96
- 323 Kirillov, A.S. (1964) Hydroxyl-bastnasite, a new mineral variety. Zapiski Vsesoyuznogo
324 Mineralogicheskogo Obshchestva, 159, 93-95
- 325 Lee, M.H., Jung, W.S. (2013) Hydrothermal Synthesis of LaCO₃OH and Ln³⁺-doped LaCO₃OH
326 Powders under Ambient Pressure and Their Transformation to La₂O₂CO₃ and La₂O₃.
327 Bulletin of the Korean Chemical Society, 34, 3609-3614
- 328 Lipin, B.R., McKay, G.A. (1989) Geochemistry and mineralogy of rare earth elements.
329 Mineralogical Society of America, 21, 259-302
- 330 Maksimovic, Z., Panto, G. (1985) Hydroxyl-bastnaesite-(Nd), a new mineral from Montenegro,
331 Yugoslavia. Mineralogical Magazine, 49, 717-720
- 332 McLellan, B.C., Corder, G.D., Ali, S.H. (2013) Sustainability of Rare Earths – An Overview of

- 333 the State of Knowledge. *Minerals*, 3, 304–317
- 334 Navrotsky, A. (1977) Progress and new directions in high temperature calorimetry. *Physics and*
335 *Chemistry of Minerals*, 2, 89–104
- 336 Navrotsky, A. (2001) Systematic trends and prediction of enthalpies of formation of refractory
337 lanthanide and actinide ternary oxide phases. *Ceramic Transactions*, 137-145
- 338 Navrotsky, A. (1997) Progress and new directions in high temperature calorimetry: revisited.
339 *Physics and Chemistry of Minerals*, 24, 222–241
- 340 Navrotsky, A. (2014) Progress and new directions in calorimetry: a 2014 perspective. *Journal of*
341 *American Ceramic Society*. 97, 3349-3359
- 342 Pathak, A. K., Khan, M. Gschneider, K. A., McCallum, R.W., Zhou, L., Sun, K., Dennis, K.W.,
343 Zhou, C., Pinkerton, F. E., Kramer, M. J., Pecharsky, V. K. (2015) Cerium: an unlikely
344 replacement of dysprosium in high performance Nd–Fe–B permanent magnets. *Advanced*
345 *Materials*, 27, 2663–2667
- 346 Pradip (1981) The surface properties and flotation of rare earth minerals, Ph.D. Thesis,
347 University of California, Berkeley.
- 348 Pradip., Li, H. C.C., Fuestenau, D. W. (2013) The Synthesis and Characterization of Rare
349 Earth Fluocarbonates. *KONA Powder and Particle Journal*, 30, 193-200
- 350 Savko, K. A., Bazikov, N. S. (2011) Phase equilibria of bastnaesite, allanite, and monazite:
351 bastnaesite-out isograd in metapelites of the Vorontsovskaya group, Voronezh
352 crystalline massif. *Petrology*, 19, 445-469
- 353 Shannon, R. D., Prewitt, C. T. (1969) Effective ionic radii in oxides and fluorides. *Acta*
354 *Crystallographica*. B25, 927-946

- 355 Smith, M.P., Henderson, P., and Peishan, Z., (1999) Reaction Relationships in the Bayan Obo
356 Fe-REE-Nb Deposit, Inner Mongolia, China: Implications for the Relative Stability of
357 Rare Earth Element Phosphates and Fluorocarbonates, Contributions to Mineralogy and
358 Petrology, 134, 294–310.
- 359 Sun, J., Kyotani, T., Tomita, A., (1986) Preparation and characterization of lanthanum carbonate
360 hydroxide, Journal of Solid State Chemistry, 64, 94-99
- 361 Ushakov, S.V., Helean, K.B., Navrotsky, A., and Boatner, L.A., (2001) Thermochemistry of
362 Rare-Earth Orthophosphates, Journal of Materials Research, 16, 2623–2633
- 363 Vallina, B., Blanco, J. D.R, Blanco, J. A., Benning, L. G.,(2014) The effect of heating on the
364 morphology of crystalline neodymium hydroxycarbonate, NdCO₃OH, Mineralogical
365 Magazine, 78, 1391-1397
- 366 Williams-Jones, A. E. Wood, S. A., (1992) A preliminary petrogenetic grid for REE
367 fluorocarbonates and associated minerals, Geochimica et Cosmochimica Acta, 56, 725-
368 738
- 369

370 **Figure Captions**

371 Figure 1. Crystal structure of CeFCO₃ shown along (a) c-axis and along (b) b- axis

372 Figure 2. Structures of orthorhombic (left panel) and hexagonal (right panel) polymorphs of
373 LaOHCO₃.

374 Figure 3. Enthalpy of formation of F/OH-bastnaesite as a function of ionic radii of RE³⁺ cation

375 Figure 4. Comparison of enthalpy of formation of bastnaesite with REPO₄. In the inset is a plot
376 of difference in enthalpy of formation of bastnaesite and monazite for different REs.

377

378

379

380

381 **Table 1.** Bastnaesite samples used for calorimetric studies, their synthesis conditions and crystal
382 systems

Sample	Method of synthesis/Temperature	Crystal system	% oxide impurities
LaFCO ₃	Urea hydrolysis /90 °C	Hexagonal	-
CeFCO ₃	Urea hydrolysis/ 90 °C	Hexagonal	-
LaOHCO ₃ -Hex	Hydrolysis of La ₂ (CO ₃) ₃ / 90 °C	Hexagonal	-
LaOHCO ₃ -ortho	Urea hydrolysis/ 180 °C	Orthorhombic	6.1
CeOHCO ₃	Urea hydrolysis/ 180 °C	Orthorhombic	5.9
DyOHCO ₃	Hydrolysis of REE nitrate with Na ₂ CO ₃ /220 °C	Orthorhombic	-
YOHCO ₃	Urea hydrolysis/ 180 °C	Orthorhombic	6.3
NdOHCO ₃	Hydrolysis of REE nitrate with Na ₂ CO ₃ /220 °C	Hexagonal	-

383

384

385

386 **Table 2.** Summary of results of Rietveld refinements of F and OH - bastnaesites

387

388

389

390

391

392

393

394

395

396

397

398

399

400

401

402

Bastnaesite	Space group	Cell parameters/Å
LaFCO ₃	P-6 2m	$a = 7.180$ (8) $c = 4.912$ (6)
CeFCO ₃	P-6 2c	$a = 7.201$ (1) $c = 9.970$ (2)
LaOHCO ₃	P-6	$a = 12.647$ (2) $c = 10.025$ (4)
LaOHCO ₃	Pnma	$a = 7.270$ (1), $b = 5.022$ (8) $c = 8.590$ (1)
CeOHCO ₃	P212121	$a = 5.005$ (4), $b = 8.535$ (7), $c = 7.305$ (6)
YOHCO ₃	P212121	$a = 4.809$ (1), $b = 6.957$ (3), $c = 8.466$ (5)
DyOHCO ₃	P212121	$a = 4.834$ (3), $b = 6.980$ (1), $c = 8.444$ (2)
NdOHCO ₃	P-6	$a = 12.327$ (1) $c = 9.880$ (4)

403 **Table 3.** Thermochemical data for F and OH - bastnaesites and for the constituents used for
 404 enthalpy of formation calculation

Sample	Enthalpy of drop solution (kJ/mol)	Enthalpy of formation from oxides at 298 K (kJ/mol)	Enthalpy of formation from elements (kJ/mol)
LaF ₃	95.87 ± 2.9		-1732.1 ± 1.2
La ₂ O ₃	-225.1 ± 3.16		-1791.6 ± 1.0
CeF ₃	89.58 ± 2.3		-1726.1 ± 1.6
Ce ₂ O ₃		371.6 ± 5.0	-1799.8 ± 1.8
CeO ₂	74.37 ± 0.75		-1090.4 ± 0.8
Y ₂ O ₃	-120.74 ± 0.94		-1932.8 ± 5.2
Nd ₂ O ₃	-163.36 ± 3.44		-1806.8 ± 3.0
Dy ₂ O ₃	-114.88 ± 2.2		-1863.4 ± 5.0
LaFCO ₃	162.50 ± 1.95	-173.07 ± 2.4	-1729.37 ± 2.72
CeFCO ₃	126.27 ± 1.77	-141.24 ± 2.58	-1708.85 ± 4.09
LaOHCO ₃ -hex	150.57 ± 0.45	-196.55 ± 1.64	-1628.78 ± 2.53
LaOHCO ₃ -ortho	181.40 ± 2.4	-227.38 ± 2.87	-1659.61 ± 4.48
CeOHCO ₃	115.74 ± 2.10	-166.03 ± 3.35	-1602.36 ± 3.65
DyOHCO ₃	142.85 ± 1.55	-133.72 ± 1.90	-1601.85 ± 5.15
NdOHCO ₃	149.90 ± 2.52	-165.01 ± 3.05	-1604.89 ± 5.74
YOHCO ₃	111.80 ± 4.8	-105.6 ± 4.82	-1608.49 ± 6.87

405 **Table 4.** Thermochemical cycle used for the calculation of enthalpy of formation of F-
 406 bastnaesites

<u>LaFCO₃</u>	
(1) $\text{LaCO}_3\text{F (s, 25 °C)} \rightarrow 0.33 \text{La}_2\text{O}_3 \text{ (sol, 700 °C)} + 0.33\text{LaF}_3 \text{ (sol, 700 °C)} + \text{CO}_2 \text{ (g, 700 °C)}$	ΔH_1
(2) $\text{La}_2\text{O}_3 \text{ (s, 25 °C)} \rightarrow \text{La}_2\text{O}_3 \text{ (sol, 700 °C)}$	ΔH_2
(3) $\text{LaF}_3 \text{ (s, 25 °C)} \rightarrow \text{LaF}_3 \text{ (sol, 700 °C)}$	ΔH_3
(4) $\text{CO}_2 \text{ (g, 25 °C)} \rightarrow \text{CO}_2 \text{ (g, 700 °C)*}$	ΔH_4
(5) $2\text{La (s, 25 °C)} + 3/2\text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{La}_2\text{O}_3 \text{ (s, 25 °C)}$	ΔH_5
(6) $\text{La (s, 25 °C)} + 3/2\text{F}_2 \text{ (g, 25 °C)} \rightarrow \text{LaF}_3 \text{ (s, 25 °C)}$	ΔH_6
(7) $\text{C (s, 25 °C)} + \text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{CO}_2 \text{ (g, 25 °C)}$	ΔH_7
(8) $0.33 \text{La}_2\text{O}_3 \text{ (s, 25 °C)} + 0.33\text{LaF}_3 \text{ (s, 25 °C)} + \text{CO}_2 \text{ (g, 25 °C)} \rightarrow \text{LaCO}_3\text{F (s, 25 °C)}$	$\Delta\text{H}_{\text{f,ox}}$
(9) $\text{La (s, 25 °C)} + 1/2\text{F}_2 \text{ (g, 25 °C)} + \text{C (s, 25 °C)} + 3/2 \text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{LaCO}_3\text{F (s, 25 °C)}$	$\Delta\text{H}_{\text{f,el}}$
$\Delta\text{H}_{\text{f,ox}} = -\Delta\text{H}_1 + 0.33\Delta\text{H}_2 + 0.33\Delta\text{H}_3 + \Delta\text{H}_4$	
$\Delta\text{H}_{\text{f,el}} = \Delta\text{H}_{\text{f,ox}} + 0.33\Delta\text{H}_5 + 0.33\Delta\text{H}_6 + \Delta\text{H}_7$	
<u>CeFCO₃</u>	
(1) $\text{CeCO}_3\text{F (s, 25 °C)} + 0.165\text{O}_2 \text{ (g, 700 °C)} \rightarrow 0.667 \text{CeO}_2 \text{ (sol, 700 °C)} + 0.33\text{CeF}_3 \text{ (sol, 700 °C)} + \text{CO}_2 \text{ (g, 700 °C)}$	ΔH_1
(2) $\text{CeO}_2 \text{ (s, 25 °C)} \rightarrow \text{CeO}_2 \text{ (sol, 700 °C)}$	ΔH_2
(3) $\text{CeF}_3 \text{ (s, 25 °C)} \rightarrow \text{CeF}_3 \text{ (sol, 700 °C)}$	ΔH_3
(4) $\text{CO}_2 \text{ (g, 25 °C)} \rightarrow \text{CO}_2 \text{ (g, 700 °C)*}$	ΔH_4
(5) $\text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{O}_2 \text{ (g, 700 °C)}$	ΔH_5
(6) $\text{Ce}_2\text{O}_3 \text{ (s, 25 °C)} + 0.5\text{O}_2 \text{ (g, 25 °C)} \rightarrow 2\text{CeO}_2 \text{ (s, 25 °C)}$	ΔH_6
(7) $2\text{Ce (s, 25 °C)} + 3/2 \text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{Ce}_2\text{O}_3 \text{ (s, 25 °C)}$	ΔH_7
(8) $\text{Ce (s, 25 °C)} + 3/2 \text{F}_2 \text{ (g, 25 °C)} \rightarrow \text{CeF}_3 \text{ (s, 25 °C)}$	ΔH_8
(9) $\text{C (s, 25 °C)} + \text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{CO}_2 \text{ (g, 25 °C)}$	ΔH_7
(10) $0.33 \text{Ce}_2\text{O}_3 \text{ (s, 25 °C)} + 0.33\text{CeF}_3 \text{ (s, 25 °C)} + \text{CO}_2 \text{ (g, 25 °C)} \rightarrow \text{CeCO}_3\text{F (s, 25 °C)}$	$\Delta\text{H}_{\text{f,ox}}$
(11) $\text{Ce (s, 25 °C)} + 1/2\text{F}_2 \text{ (g, 25 °C)} + \text{C (s, 25 °C)} + 3/2 \text{O}_2 \text{ (g, 25 °C)} \rightarrow \text{CeCO}_3\text{F (s, 25 °C)}$	
$\Delta\text{H}_{\text{f,ox}} = -\Delta\text{H}_1 + 0.667\Delta\text{H}_2 + 0.33\Delta\text{H}_3 + \Delta\text{H}_4 - 0.165\Delta\text{H}_5 + 0.33\Delta\text{H}_6$	
$\Delta\text{H}_{\text{f,el}} = \Delta\text{H}_{\text{f,ox}} + 0.33\Delta\text{H}_7 + 0.33\Delta\text{H}_8 + \Delta\text{H}_9$	

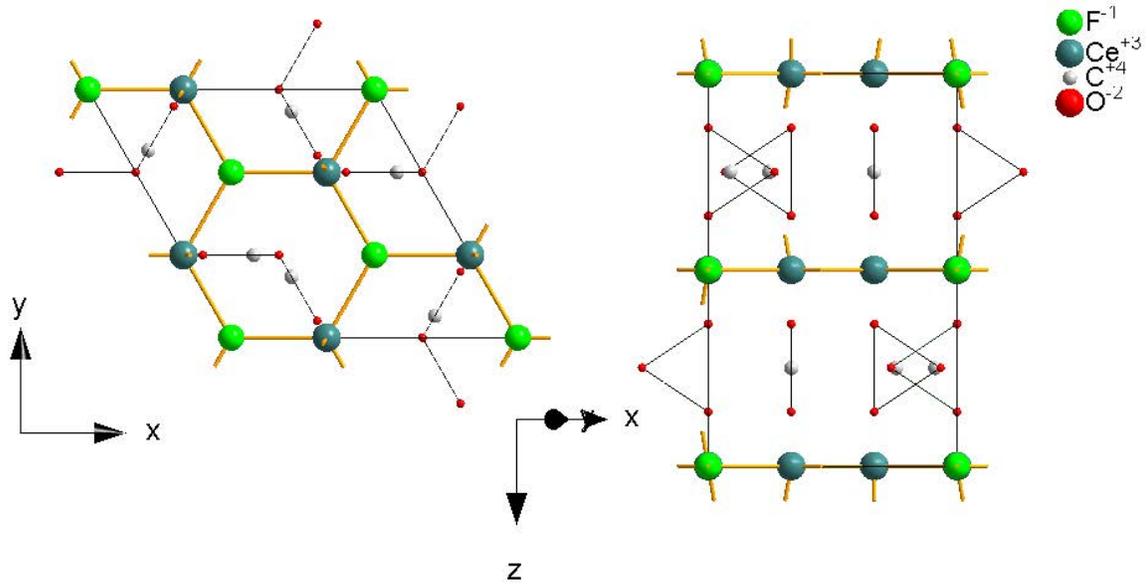
407 * $\text{CO}_2 = 32.07 \text{ kJ/mol}$ and $\text{O}_2 = 21.74 \text{ kJ/mol}$ (Calculated from heat capacity measurements).

408 **Table 5.** Thermochemical cycle used to calculate enthalpies of formation of OH-bastnaesites.

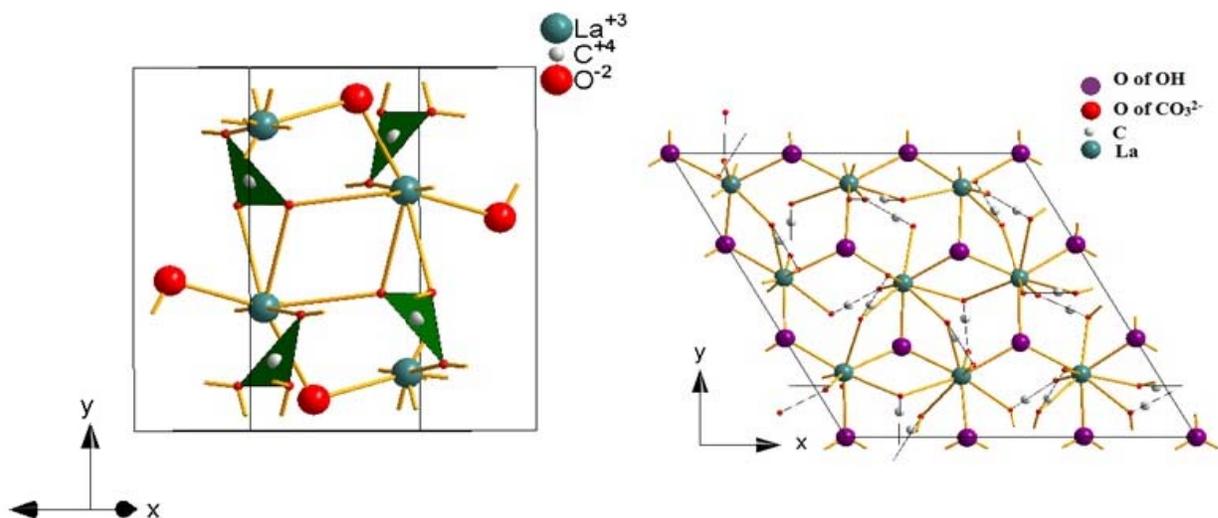
409

RECO₃OH (RE – La, Y, Nd and Dy)	
410	(1) RECO ₃ OH _(s, 25 °C) → 0.5 RE ₂ O _{3(sol, 700 °C)} + CO _{2(g, 700 °C)} + 0.5 H ₂ O _(g, 700 °C) ΔH ₁
	(2) RE ₂ O _{3(s, 25 °C)} → RE ₂ O _{3(sol, 700 °C)} ΔH ₂
411	(3) H ₂ O _(l, 25 °C) → H ₂ O _(g, 700 °C) ΔH ₃
	(4) CO _{2(g, 25 °C)} → CO _{2(g, 700 °C)} ΔH ₄
412	(5) 2RE _(s, 25 °C) + 3/2 O _{2(g, 25 °C)} → Ce ₂ O _{3(s, 25 °C)} ΔH ₅
413	(6) C _(s, 25 °C) + O _{2(g, 25 °C)} → CO _{2(g, 25 °C)} ΔH ₆
414	(7) H ₂ + 1/2 O _{2(g, 25 °C)} → H ₂ O _(l, 25 °C) ΔH ₇
	(8) 0.5 RE ₂ O _{3(s, 25 °C)} + CO _{2(g, 25 °C)} + 0.5 H ₂ O _(l, 25 °C) → RECO ₃ OH _(s, 25 °C) ΔH _{f,ox}
415	(9) RE _(s, 25 °C) + C _(s, 25 °C) + 1/2 H _{2(g, 25 °C)} + 2O _{2(g, 25 °C)} → RECO ₃ OH _(s, 25 °C) ΔH _{f,el}
416	ΔH _{f,ox} = - ΔH ₁ + 0.5ΔH ₂ + 0.5 ΔH ₃ + ΔH ₄
	ΔH _{f,el} = ΔH _{f,ox} + 0.5ΔH ₅ + 0.5 ΔH ₇ + ΔH ₆
417	
418	CeCO₃OH
419	
420	(1) CeCO ₃ OH _(s, 25 °C) + 0.25O _{2(g, 700 °C)} → CeO _{2(sol, 700 °C)} + CO _{2(g, 700 °C)} + 0.5 H ₂ O _(g, 700 °C) ΔH ₁
421	
422	(2) CeO _{2(s, 25 °C)} → CeO _{2(sol, 700 °C)} ΔH ₂
423	
424	(3) CO _{2(g, 25 °C)} → CO _{2(g, 700 °C)} ΔH ₃
425	(4) O _{2(g, 25 °C)} → O _{2(g, 700 °C)} ΔH ₄
426	
427	(5) Ce ₂ O _{3(s, 25 °C)} + 0.5O _{2(g, 25 °C)} → 2CeO _{2(s, 25 °C)} ΔH ₅
428	(6) H ₂ O _(l, 25 °C) → H ₂ O _(g, 700 °C) ΔH ₆
429	
430	(7) 2Ce _(s, 25 °C) + 3/2 O _{2(g, 25 °C)} → Ce ₂ O _{3(s, 25 °C)} ΔH ₇
431	(8) C _(s, 25 °C) + O _{2(g, 25 °C)} → CO _{2(g, 25 °C)} ΔH ₈
432	
433	(9) H ₂ + 1/2 O _{2(g, 25 °C)} → H ₂ O _(l, 25 °C) ΔH ₉
434	(10) 0.5Ce ₂ O _{3(s, 25 °C)} + CO _{2(g, 25 °C)} + 0.5 H ₂ O _(l, 25 °C) → CeCO ₃ OH _(s, 25 °C) ΔH _{f,ox}
435	ΔH _{f,ox} = - ΔH ₁ + ΔH ₂ + ΔH ₃ - 0.25ΔH ₄ + 0.5ΔH ₅ + 0.5 ΔH ₆
436	ΔH _{f,el} = ΔH _{f,ox} + 0.5ΔH ₇ + 0.5 ΔH ₉ + ΔH ₈
437	
438	
439	

440 # ΔH_{f,el} for CO₂ and H₂O is -393.51±1.3 and -285.83±0.42 kJ/mol respectively (From Robie and
 441 Hemingway 1995)



459



460

461

462 **Figure 2.** Structures of orthorhombic (left panel) and hexagonal (right panel) polymorphs of

463 LaOHCO₃.

464

465

466

467

468

469

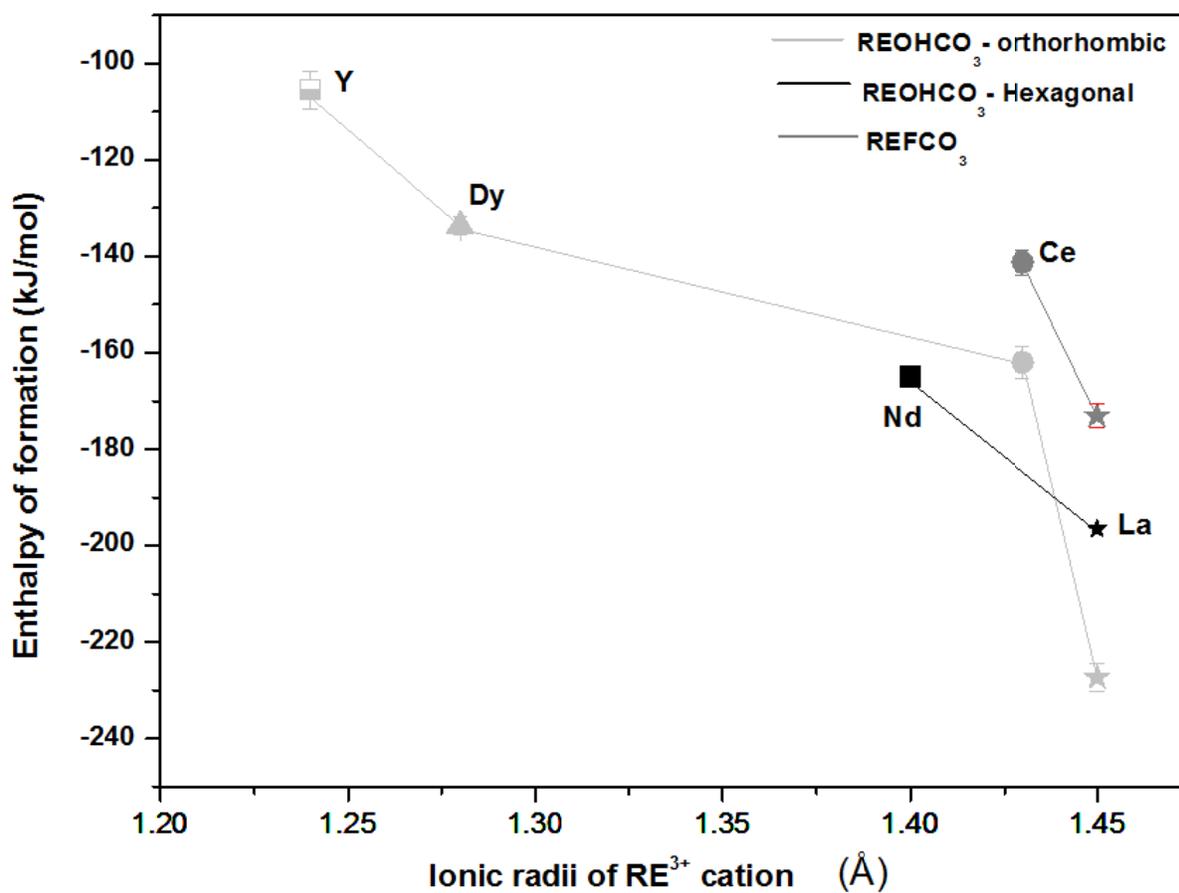
470

471

472

473

474



475

476

477 **Figure 3.** Enthalpy of formation of F/OH-bastnaesite as a function of ionic radii of RE³⁺ cation

478

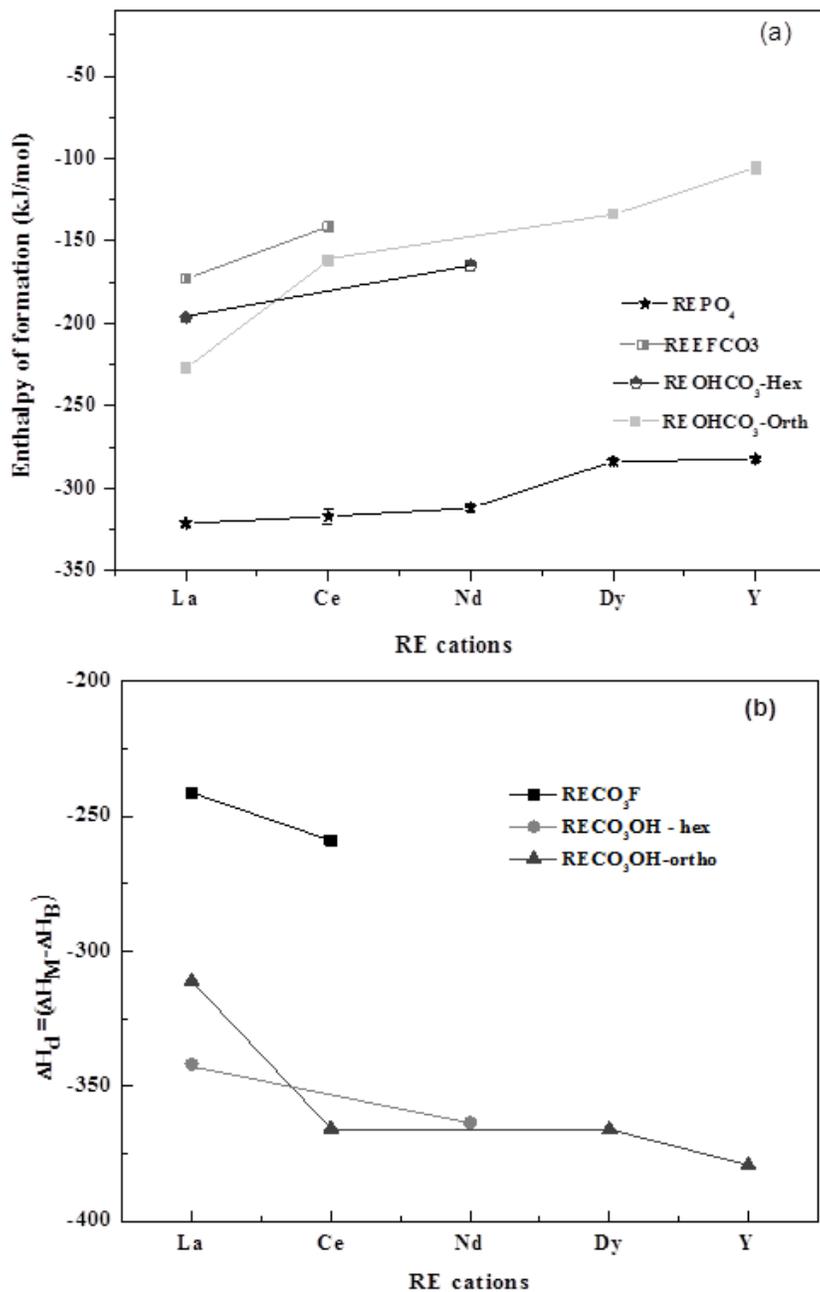
479

480

481

482

483



484

485

486 **Figure 4.** (a) A plot of enthalpy of formation values obtained for bastnaesite and monazite as a

487 function of different RE. (b) a plot of difference in enthalpy of formation of bastnaesite and

488 monazite for different REs