

# 1 **Single-Crystal Equations of State of Magnesiowüstite at High Pressures**

## 2 **(Revision 1)**

3 Gregory J. Finkelstein<sup>1\*</sup>, Jennifer M. Jackson<sup>1</sup>, Wolfgang Sturhahn<sup>1</sup>, Dongzhou Zhang<sup>2</sup>, E. Ercan  
4 Alp<sup>3</sup>, Thomas S. Toellner<sup>3</sup>

5 <sup>1</sup>Seismological Laboratory, California Institute of Technology, 1200 E. California Blvd., MS  
6 252-21, Pasadena, California 91125, USA

7 <sup>2</sup>Partnership for Extreme Crystallography, University of Hawaii, School of Ocean and Earth  
8 Science and Technology, University of Hawaii, 1680 East West Road (POST Bldg 819E),  
9 Honolulu, Hawaii 96822, USA

10 <sup>3</sup>Advanced Photon Source, Argonne National Laboratory, 9700 South Cass Avenue, Argonne,  
11 Illinois 60439, USA

12 \*Email Address: [gjfinkel@gps.caltech.edu](mailto:gjfinkel@gps.caltech.edu)

13

14

15

16

17

18

19 Keywords: magnesiowüstite, single-crystal diffraction, elasticity, equations of state, Mössbauer  
20 spectroscopy, high pressure

21

22

23

## 24 **Abstract**

25 Solid solutions of (Mg,Fe)O with high iron enrichment may be an important component of  
26 ultralow-velocity zones at Earth's core-mantle boundary. However, to date there have been few  
27 high-precision studies on the elastic properties of these materials. In this study we present results  
28 on the compression of (Mg<sub>0.22</sub>Fe<sub>0.78</sub>)O magnesiowüstite in both neon and helium pressure media  
29 using single-crystal diffraction to ~55 GPa. In addition, our sample was characterized by time-  
30 domain synchrotron Mössbauer spectroscopy at ambient pressure using an extended time range  
31 that resulted in vastly improved energy resolution. The combination of these high-resolution  
32 techniques tightly constrains the presence of a defect-structure component at room pressure due  
33 to 4.7 mol% tetrahedrally-coordinated ferric iron, resulting in a renormalized composition of  
34 (Mg<sub>0.215</sub>Fe<sub>0.762 ± 0.023</sub>)O. Both high-pressure diffraction datasets are well described by a 3rd-order  
35 Birch-Murnaghan equation of state. The best fit-parameters for a crystal with cubic structure in  
36 helium are  $K_{0T} = 148(3)$  GPa,  $K'_{0T} = 4.09(12)$ , and  $V_0 = 78.87(6)$  Å<sup>3</sup>. Increasing differential  
37 stress in the neon-containing sample chamber was correlated with increasing apparent distortion  
38 of the initially cubic unit cell, requiring a lower-symmetry hexagonal cell to fit the data above  
39 ~20 GPa. For fit equations of state, we determine the pressure-dependent correlation ellipses for  
40 the equation of state parameters and compare with previously published single-crystal diffraction  
41 data from (Mg,Fe)O crystals in a helium medium. We make two main observations from the  
42 datasets using a helium pressure medium:  $K_{0T}$  decreases as a function of increasing iron content  
43 from periclase to wüstite and  $K'_{0T}$  is consistent with an approximately constant value of 4.0 that  
44 is independent of iron content, at least up to (Mg,Fe)O containing ~78 mol% FeO. In  
45 combination with previously reported thermal parameters, we compute the density of Mw78 at  
46 core-mantle boundary conditions and discuss the implications.

47

## 48 **Introduction**

49

50 Ferromagnesian oxides are a primary component of the earth's lower mantle, along with the  
51 dominant bridgmanite and minor  $\text{CaSiO}_3$  perovskite phases (Irifune 1994; Tschauner et al.  
52 2014). While the bulk of the lower mantle likely contains  $(\text{Mg,Fe})\text{O}$  ferropericlase solid  
53 solutions with ~10-40mol% of FeO (Fei et al. 1996; H. Mao 1997; Sinmyo et al. 2008), the  
54 presence of material significantly more enriched in Fe has been hypothesized as an explanation  
55 for some of the complex seismic structure observed near the core-mantle boundary (CMB). In  
56 particular, ultralow-velocity zones (ULVZs), 10s of km-thick patches at the CMB, show a ~10%  
57 reduction in seismic P-wave velocity and a ~30% reduction in seismic S-wave velocity from  
58 average lower-mantle values (Wen 1998; Thorne and Garnero 2004; Rost et al. 2006).

59

60 While a number of suggestions have been put forth in the literature to explain the origin of these  
61 features (Manga and Jeanloz 1996; Williams et al. 1998; W. L. Mao et al. 2006), one hypothesis  
62 that has gained traction in recent years proposes Fe-rich oxide components to explain the  
63 observed low sound velocities. Work by Wicks et al. (2010) used nuclear resonant inelastic  
64 scattering to show that the shear velocity of  $(\text{Mg}_{0.16}\text{Fe}_{0.84})\text{O}$  magnesiowüstite was ~55% slower  
65 than endmember MgO at CMB pressures, and that only ~10 vol. % of this material mixed with a  
66 silicate assemblage would be needed to match ULVZ seismic velocities. Subsequent simulations  
67 (Bower et al. 2011) that explored the geodynamic effects of enriching ULVZ material with  
68 magnesiowüstite produced morphologies consistent with observed ULVZ seismic topography  
69 (Sun et al. 2013).

70

71 While selected properties of (Mg,Fe)O with low iron concentrations have been determined at  
72 pressures above 1 Mbar and temperatures over 2000 K (Westrenen et al. 2005; Lin et al. 2006;  
73 2007; Z. Mao et al. 2011), until recently, the iron-rich members of this solid solution received  
74 less attention (Ohta et al. 2014; Wicks et al. 2015; 2017), and still little is known about their  
75 crystallographic and elastic properties. At ambient conditions, all (Mg,Fe)O solid solutions  
76 crystallize in the cubic B1 structure. Single-crystal x-ray diffraction (SXRD) has shown,  
77 however, that endmember wüstite, FeO, undergoes a slight rhombohedral distortion from cubic  
78 symmetry at ~18-23 GPa and 300 K (Hazen and Jeanloz 1984; Shu et al. 1998; Jacobsen et al.  
79 2005), whereas (Mg<sub>0.73</sub>Fe<sub>0.27</sub>)O remains cubic to at least 51 GPa (Jacobsen et al. 2005). A  
80 powder diffraction (PXRD) study on (Mg<sub>0.22</sub>Fe<sub>0.78</sub>)O reported a rhombohedral transition at 20-40  
81 GPa (Zhuravlev et al. 2010), whereas a PXRD study on (Mg<sub>0.06</sub>Fe<sub>0.94</sub>)O bracketed the cubic to  
82 rhombohedral transition between 13 and 24 GPa (Wicks et al. 2015). The rhombohedral  
83 distortion may also be associated with a magnetic transition at or near the structural transition  
84 pressure (Struzhkin et al. 2001; Jacobsen and Spetzler 2004; Kantor et al. 2004; Wicks et al.  
85 2010).

86

87 Diffraction methods are generally used in compression experiments to probe the volume  
88 reduction of the crystallographic unit cell with pressure, and thus bulk compressibility, as well as  
89 any structural transitions that may occur. Pressure-volume relationships are frequently modeled  
90 with an equation of state, which results in an estimate of a material's isothermal bulk modulus,  
91  $K_T$ , as well as its pressure derivative,  $K'_T$ . When constrained well by the measured data, these

92 parameters can be used in the calculation of sound velocities and density profiles and then  
93 compared to observed seismic data.

94

95 In a high-pressure apparatus such as a diamond anvil cell, where sample dimensions are  
96 frequently limited to tens of microns, diffraction from powdered samples often suffers from poor  
97 grain statistics, has difficult-to-fit overlapping peaks, and the scattered intensity is significantly  
98 weaker than that of similarly-sized single crystals. Combined, these effects tend to result in  
99 increased data scatter of compression curves measured from powdered samples compared with  
100 those from single crystals. Therefore, in order to more conclusively determine its compression  
101 behavior, here we expand on the previous study of ground, polycrystalline  $(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$  by  
102 compressing and measuring the SXR D of single crystals drawn from the same batch of sample  
103 used in the earlier work. Furthermore, we evaluate the influence of the pressure medium by  
104 running experiments in two different commonly used quasihydrostatic media: neon and helium.  
105 We complement the SXR D measurements with high-resolution time-domain synchrotron  
106 Mössbauer spectroscopy to tightly constrain the ferric iron concentration, which coupled with  
107 diffraction data, helps constrain the defect concentration.

108

## 109 **Experimental Procedures**

110

111 We prepared two diamond-anvil cells (DAC) for high-pressure experiments. One cell was a  
112 standard Princeton-design symmetric DAC with a cubic boron nitride seat on the cylinder side  
113 and a tungsten carbide seat on the piston side. Standard modified brilliant-cut diamond anvils  
114 with flat 300  $\mu\text{m}$  culets were used on both seats. The other DAC was a BX90 cell (Kantor et al.

115 2012) with symmetric Bohler-Almax-geometry seats (Boehler and De Hantsetters 2007) and  
116 diamonds with 250  $\mu\text{m}$  beveled culets. Each cell's  $\sim 40$   $\mu\text{m}$  thick rhenium gasket was loaded with  
117 a nearly identical sample configuration that consisted of two or three  $<10$   $\mu\text{m}$  thick ruby sphere  
118 pressure calibrants and platelets of synthetic  $(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$  magnesiowüstite, hereafter referred  
119 to as Mw78,  $<10$   $\mu\text{m}$  in thickness and  $\sim 20\text{-}30$   $\mu\text{m}$  in diameter. These small pieces were broken  
120 off from a larger single crystal. Additional details regarding the synthesis and characterization of  
121 this sample material may be found in previous publications (Jacobsen 2002; Mackwell et al.  
122 2005). The prepared DACs were subsequently gas-loaded with neon (symmetric DAC) or helium  
123 (BX90 DAC) pressure media at  $\sim 25,000$  PSI using either the Caltech or COMPRES/GSECARS  
124 gas-loading system (Rivers et al. 2008), respectively (see supplementary information for images  
125 of sample chambers at high pressure).

126

127 The ambient diffraction pattern of a portion of the magnesiowüstite sample was characterized at  
128 the 11.3.2 beamline of the Advanced Light Source at Lawrence Berkeley National Laboratory ( $\lambda$   
129  $= 0.6702$   $\text{\AA}$ ). Single crystal diffraction measurements at high pressures were collected on each  
130 DAC sample configuration at the Advanced Photon Source at Argonne National Laboratory. The  
131 He-medium experiment was conducted at the PX<sup>2</sup> BM-C beamline ( $\lambda = 0.4340$   $\text{\AA}$ ) and the Ne-  
132 medium experiment at the ID-D beamline ( $\lambda = 0.3100$   $\text{\AA}$ ) of the GSECARS sector. We achieved  
133 maximum pressures of 55.9(3) and 53.0(1.5) GPa, respectively. Downstream X-ray opening  
134 angles of the loaded cells were determined to be  $48^\circ$  (Ne) and  $60^\circ$  (He). The number of measured  
135 diffraction peaks at each pressure point is determined by lattice-parameter lengths, x-ray energy,  
136 and opening angle of the diamond anvil cell. After each pressure increase, the cell was allowed  
137 to rest for at least 20 minutes before collection of diffraction, minimizing pressure creep. Ruby

138 spectra were obtained immediately after the diffraction patterns were collected for a given  
139 compression step. Several experiments using similar methods for the collection of single crystal  
140 diffraction patterns at high pressure can be found in the recent literature, providing additional  
141 details on the procedure (Dera et al. 2013a; Finkelstein et al. 2014).

142

143 Experimental pressures were determined using the ruby scale of Dewaele et al. (2004). The  
144 uncertainty in pressure at each step was estimated as the standard deviation of the pressures  
145 measured for the different rubies in each sample chamber. Diffraction peaks were fit and lattice  
146 parameters refined in either the APEX 2 (ambient measurements) (Bruker 2012a) or the  
147 GSE\_ADA/RSV (high-pressure measurements) (Dera et al. 2013b) software package. Unit cells  
148 were indexed based on an analysis of reciprocal lattice difference vectors within the RSV  
149 interface using the CELL\_NOW (Bruker 2012b) program to (refer to discussion below and  
150 supplementary information for example images of indexed reciprocal lattices and unique  
151 diffraction peaks used). After converting lattice parameters to unit cell volumes, a 3<sup>rd</sup> order  
152 Birch-Murnaghan equation of state was fit to the pressure-volume data of each experimental run  
153 using the MINUTI software package (version 2.0) (Sturhahn 2015b). The 3<sup>rd</sup> order Birch-  
154 Murnaghan equation of state takes the form:

155

$$P(V) = \frac{3K_{0T}}{2} \left[ \left( \frac{V_0}{V} \right)^{\frac{7}{3}} - \left( \frac{V_0}{V} \right)^{\frac{5}{3}} \right] \left\{ 1 + \frac{3}{4} (K'_{0T} - 4) \left[ \left( \frac{V_0}{V} \right)^{\frac{2}{3}} - 1 \right] \right\}$$

156

157 The parameters of pressure, unit cell volume, zero-pressure unit cell volume, and isothermal  
158 zero-pressure bulk modulus and its pressure derivative are represented by the symbols  $P$ ,  $V$ ,  $V_0$ ,

159  $K_{0T}$ , and  $K'_{0T}$ , respectively. Using a new feature of MINUTI, we computed the covariance matrix  
160 of each of our equation of state fits as a function of pressure. The covariance matrices were used  
161 to calculate 68% and 95% confidence ellipses for our fit equation of state parameters, also as a  
162 function of pressure (Sturhahn 2015b); the reader is also referred to Angel (2000) for more  
163 information on covariance matrices and confidence ellipses as they relate to equations of state as  
164 a function of pressure.

165

166 To complement our diffraction experiments and aid in the determination of the ferric content of  
167 our sample, we applied synchrotron Mössbauer spectroscopy (SMS) at APS Sector 3-ID-B in  
168 hybrid top-up mode. The measured single-crystalline platelet of Mw78 was  $\sim 110 \mu\text{m}$  thick,  
169 resulting in an effective thickness of about 10. The x-rays were focused to  $\sim 10 \times 11 \mu\text{m}^2$  (D.  
170 Zhang et al. 2015), and a high-resolution monochromator provided an energy bandwidth of 1  
171 meV at the 14.4125 keV  $^{57}\text{Fe}$  nuclear transition energy (Toellner 2000). The SMS spectrum was  
172 fit using the CONUSS software package (Sturhahn 2015a). With our use of APS hybrid mode, a  
173 time range of 35 ns to 860 ns became accessible for the SMS spectrum. This extended time range  
174 translates into a much improved energy resolution of about 1.1  $\Gamma$  ( $\Gamma = 4.66 \text{ neV}$  is the natural line  
175 width of the 14.4125 keV  $^{57}\text{Fe}$  nuclear excited state) which compares to  $>2 \Gamma$  for conventional  
176 Mössbauer experiments using a radioactive source (Sturhahn 2001, Eq. 13). At present,  
177 experiments using the hybrid mode require long data collection times due to limited useful x-ray  
178 intensity. Therefore, our SMS spectrum collected with the extended time range took about 20  
179 hours of beamtime. However, there is a clear advantage in taking SMS spectra in hybrid mode.  
180 Out of the various kinds of Mössbauer spectroscopy, this configuration has the best energy

181 resolution to date, leading to markedly improved accuracy in the fit of potential minor sites and  
182 to finely resolved small differences in hyperfine parameters.

183

## 184 **Room-Pressure Results and Discussion**

185

186 At ambient conditions, the Mw78 crystal can be nominally indexed with a cubic unit cell (Table  
187 1). However, upon closer examination of the diffraction peaks at long exposure times, faint,  
188 diffuse satellites are observed around the primary peaks, with their intensity significantly  
189 decreasing at high two-theta angles (Figure 1). This additional scattering has been observed in  
190 several previous studies on the wüstite endmember (Hazen and Jeanloz 1984; Jacobsen et al.  
191 2005), but has not previously been reported for (Mg,Fe)O solid solutions. The complex intensity  
192 distribution around the satellite peaks of wüstite has been interpreted as being directly related to  
193 an inhomogeneous paracrystalline structure that consists of regions of high defect-cluster density  
194 and regions of low defect-cluster density (Welberry and Christy 1997; Saines et al. 2013). As the  
195 observed scattering due to the defect structure is extremely weak, in this paper we use the Bragg  
196 peak positions alone to determine lattice parameter values, even though we recognize that they  
197 do not fully describe the sample's structure.

198

199 By fitting an SMS spectrum from our sample, we were able to quantify its ferric and ferrous  
200 components. While the bulk (Mg,Fe)O structure is composed of octahedrally-coordinated ferrous  
201 Fe, the defect structure is thought to be made of clusters of tetrahedrally-coordinated ferric Fe.  
202 These two types of Fe can be readily distinguishable in an SMS spectrum, as the hyperfine fields

203 (e.g. isomer shift, quadrupole splitting) are sensitive to both oxidation state and coordination  
204 number.  
205  
206 Previously published SMS and conventional Mössbauer studies on (Mg,Fe)O and wüstite (e.g  
207 Johnson et al. (1969), Lin et al. (2006), McCammon et al. (1985), and Solomatova et al. (2016))  
208 have used several different models to fit the complex observed spectra, ranging from a single  
209 doublet (Lin et al. 2006) to multiple doublets and singlets (McCammon and Price 1985). After  
210 investigating these and similar models, we concluded that our spectrum required a model with  
211 four Fe sites to achieve an optimal and physically meaningful fit. These four sites consist of three  
212 doublets and one singlet. The SMS spectrum was collected without an external reference  
213 standard, such as a stainless steel foil, which would permit determination of all isomer shifts.  
214 However, by fixing three of the doublets to an isomer shift of 1.0 mm/s (a value that is  
215 reasonable for an octahedrally-coordinated ferrous Fe site), we determine the relative isomer  
216 shift of the fourth site. Combined, the three doublet sites correspond to ~94% of the Fe sites, and  
217 have quadrupole splitting values ranging from 0.6 to 1.3 mm/s (Table 1). In light of this  
218 information, we assign the three doublets to high-spin ferrous sites. The singlet, on the other  
219 hand, exhibited an isomer shift ~0.95 mm/s lower than the other sites, which is consistent only  
220 with tetrahedrally-coordinated high-spin ferric Fe (Dyar et al. 2006). This interpretation is  
221 distinct from select Mössbauer studies on wüstite, where the singlet site was interpreted as a  
222 component of the metallic iron sextet (McCammon and Price 1985). Renormalizing the  
223 measured Mw78 sample composition to take into account the vacancies necessitated by the ferric  
224 Fe content results in a composition of  $(\text{Mg}_{0.215}\text{Fe(II)}_{0.716}\text{Fe(III)}_{0.046-0.023})\text{O}$  if ferric and ferrous  
225 Fe are separated in the chemical formula, and  $(\text{Mg}_{0.215}\text{Fe}_{0.762-0.023})\text{O}$  if not distinguished.

226

## 227 **High-Pressure Results and Discussion**

### 228 **Phase Stability (subsection)**

229 At low pressures, the Mw78 crystal in neon has a diffraction pattern that can be indexed with the  
230 expected cubic unit cell. At high pressures, the diffraction peak positions are no longer  
231 compatible with this unit-cell geometry, and a lower-symmetry cell is required. We do not see  
232 evidence for twinning in the high-pressure phase. With the obtained coverage of reciprocal  
233 space, a hexagonal unit cell is sufficient for fitting the observed unit cell distortion. Deviations  
234 from cubic symmetry can be quantified by monitoring the  $c/a$  lattice parameter ratio. Figure 3a  
235 (bottom) shows the pressure dependence of the  $c/a$  ratio which has a value of  $\sqrt{6} \approx 2.45$  for a  
236 cubic material. A Mw78 crystal in neon remains cubic until  $\sim 20$  GPa, above which it becomes  
237 increasingly hexagonal (as observed by analysis of the reciprocal lattice difference vectors – see  
238 Figure 3b). In contrast, a Mw78 crystal in the helium pressure medium gives a  $c/a$  ratio near 2.45  
239 over the whole pressure range, indicating that the unit cell remains metrically cubic up to the  
240 maximum pressure of our study. Lattice parameter values refined in both the cubic and  
241 hexagonal settings are provided in Table 2.

242

243 Previous high-pressure studies on  $(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$  and  $(\text{Mg}_{0.06}\text{Fe}_{0.94})\text{O}$  in neon reported  
244 rhombohedral transitions around 20-40 GPa and 13-24 GPa, respectively (Zhuravlev et al. 2010).  
245 These results are compatible with our result on Mw78 in neon. High-pressure studies of helium-  
246 loaded wüstite single crystals (Shu et al. 1998; Jacobsen et al. 2005) reported fourfold non-  
247 merohedral twinning inherent to the mechanism of a rhombohedral transition in wüstite. Our  
248 single-crystal diffraction study on Mw78 in helium shows no evidence for a lower symmetry

249 distortion up to 55 GPa. The addition of MgO to wüstite appears to prevent such a transition and  
250 to extend the stability field of the cubic phase.

251

252 Since ruby spheres were placed in several locations inside the sample chamber of each DAC  
253 (refer to supplementary figures for pictures), we were able to use the standard deviation of the  
254 calibrated ruby pressures as a proxy for nonhydrostaticity experienced by the sample. As seen in  
255 Figure 3a (top), this quantity remains near zero for the helium-loaded sample chamber, but  
256 rapidly increases in the neon-loaded sample chamber above ~20 GPa. The nonhydrostaticity  
257 proxy shows the exact same pressure dependence as the hexagonal distortion of Mw78 described  
258 earlier. Previous work also investigated the quasihydrostatic pressure limits of helium and neon  
259 media using the ruby standard deviation method up to ~40 and ~50 GPa, respectively (Klotz et  
260 al. 2009). Their results suggest that while nonhydrostaticity in a neon medium increases rapidly  
261 above ~15 GPa, nonhydrostaticity in a helium medium increases gradually above ~23 GPa.

262

### 263 **Equations of State (subsection)**

264 Figure 4 shows compression curves for single-crystalline Mw78 in helium (indexed as cubic)  
265 and neon (indexed as cubic and hexagonal), and for powdered (Mg<sub>0.22</sub>Fe<sub>0.78</sub>)O in neon  
266 (Zhuravlev et al. 2010) (indexed as cubic). There is a discernible difference between the Mw78  
267 in neon and in helium, as the Mw78 sample in helium is more compressible than those in neon.  
268 The compression curves are smooth for each phase and are each well described by a 3<sup>rd</sup> order  
269 Birch-Murnaghan equation of state. Our optimization procedure using the MINUTI software  
270 (Sturhahn 2015b) incorporated the measured ambient unit cell volume of  $78.94 \pm 0.1 \text{ \AA}^3$  for our  
271 single-crystal cubic-indexed measurements as a prior on  $V_0$ . For the crystal in helium, the other

272 parameters,  $K_{0T}$ , and  $K'_{0T}$ , were allowed to vary freely. For the crystal in neon, due to the limited  
273 number of data points in respective phase regions, a strong prior of  $\pm 0.1$  was put on  $K'_{0T}$  for both  
274 the cubic (1.3-19.2 GPa) and hexagonal (24.1-53.3 GPa) phases. The best fits to the data were  
275 achieved with  $V_0$ ,  $K_{0T}$ , and  $K'_{0T}$  values of 78.742(14)  $\text{\AA}^3$ , 163.0(1.0) GPa, and 4.02(10) for the  
276 cubic-indexed single crystal Mw78 in neon and 78.87(6)  $\text{\AA}^3$ , 148(3) GPa, and 4.09(12) for the  
277 cubic-indexed single crystal Mw78 in helium,  $V_0$ ,  $K_{0T}$ , and  $K'_{0T}$  values of 78.5(2)  $\text{\AA}^3$ , 166(5)  
278 GPa, and 4.01(10) for the cubic-indexed powdered Mw78 in neon, and  $V_0$ ,  $K_{0T}$ , and  $K'_{0T}$  values  
279 of 58.7(4)  $\text{\AA}^3$ , 176.8(1.1) GPa, and 4.00(10) for the hexagonal-indexed single crystal Mw78 in  
280 neon. These fits are plotted in Figure 4. While both single-crystal cubic compression curves have  
281 similar  $K_{0T}$  values ( $\sim 4$ ), the cubic phase of the sample in neon exhibits a significantly higher  
282 value for  $K_{0T}$  compared with the sample in helium. This corresponds to higher values of  $V/V_0$  for  
283 the sample in neon at pressures above  $\sim 5$  GPa (Figure 4), and is consistent with Mw78 being  
284 very sensitive to nonhydrostaticity, as neon crystallizes at  $\sim 4.8$  GPa (Klotz et al. 2009). While in  
285 our diffraction experiments we cannot confidently resolve a hexagonal distortion in Mw78 until  
286 above  $\sim 20$  GPa, the unit cell volume appears to be affected by neon crystallization below  $\sim 20$   
287 GPa.

288

289 In order to provide context for our results, we re-evaluated equations of state for data on several  
290 (Mg,Fe)O compositions using MINUTI or direct comparison with our study (Table 3). Pressure-  
291 volume data were sourced from single-crystal diffraction studies by Jacobsen et al. (2002; 2005;  
292 2008) and a powder diffraction study by Zhuravlev et al. (2010). Since Zhuravlev et al. (2010)  
293 did not collect data near zero pressure, we applied a prior of  $4.0 \pm 0.1$  on  $K'_{0T}$  in the equation of  
294 state fit. The Jacobsen et al. (2005) data on  $\text{Fe}_{0.93}\text{O}$  covered a limited pressure range, and

295 therefore we also applied a prior of  $4.0 \pm 0.1$  on  $K'_{0T}$ . This value is consistent with the given  
296  $Fe_{0.93}O$  data, as well as with the Mw78 single-crystal diffraction measurements in helium. All  
297 other parameters remained unconstrained during fitting.

298

299 All examined (Mg,Fe)O datasets are fit well by a 3<sup>rd</sup> order Birch-Murnaghan equation of state.  
300 We show the 68% and 95% confidence ellipses (joint probability) for each pair of equation of  
301 state fit parameters as a function of pressure in Figure 5. Confidence ellipses, which are derived  
302 from the fit-parameter covariance matrix, visualize magnitude and sign of parameter correlations  
303 and provide insight into the acceptable parameter space. The primary correlation in a 3<sup>rd</sup> order  
304 Birch-Murnaghan equation of state fit is between  $K_T$  and  $K'_T$ . As discussed in Angel (2000), this  
305 parameter correlation is initially negative but continuously changes to a positive correlation with  
306 pressure. Consequently, for every equation of state fit, there is some pressure where  $K_T$  and  $K'_T$   
307 are uncorrelated and uncertainties of the fit parameters are lowest. This pressure primarily  
308 depends on the pressure range of a dataset but is also sensitive to the data distribution and  
309 uncertainties. In our fits, the pressure where  $K_T$  and  $K'_T$  are uncorrelated is between 5 and 20  
310 GPa, depending on the dataset. Above this pressure,  $K_T$  and  $K'_T$  are positively correlated, and  
311 their uncertainties increase rapidly for pressures outside of the measured range.

312

313 In Figure 6, we present F-f plots for all our (Mg,Fe)O equation of state fits, with separate  
314 subplots for samples compressed in helium, neon, and 4:1 methanol-ethanol pressure media. F-f  
315 plots show normalized pressure (F) as a function of Eulerian strain (f), and can be useful for  
316 estimating  $K_{0T}$  and  $K'_{0T}$  parameter values of a 3<sup>rd</sup> order Birch-Murnaghan equation of state. In an  
317 F-f plot, if a straight line can be fit to a given dataset, its F-axis intercept corresponds to the

318 approximate value of  $K_{0T}$ . A positive or negative slope of the line indicates a value of  $K'_{0T}$  that is  
319 either greater than or less than 4, respectively. A slope of zero means that  $K'_{0T}$  equals 4 which  
320 effectively reduces the order and represents a 2<sup>nd</sup> order Birch-Murnaghan equation of state. If a  
321 straight line does not properly model the F-f plot a higher-order equation of state is required to fit  
322 the pressure-volume data. Specific results for each pressure medium will be discussed below.

323

324 The F-f plots of the samples in both helium and neon media show near-zero slopes, which  
325 indicates a compositionally independent  $K'_{0T}$  value of about 4 that is consistent with observations  
326 of low-Fe compositions with  $\text{FeO} \leq 17$  mol% from ultrasonic and Brillouin measurements. There  
327 have been no high-pressure reports of both the compressional and shear velocities to compute the  
328 bulk modulus, required for constraining a pressure derivative, for samples with greater iron  
329 content. The positive F-f slopes in the samples compressed in 4:1 methanol-ethanol require a  
330 different explanation, as this pressure medium is believed to remain hydrostatic to  $\sim 10$  GPa  
331 (Klotz et al. 2009). Jacobsen et al. (2005) speculated that they may be related to the limited  
332 maximum compression of the samples or to differences between the quartz pressure scale and  
333 others that are used more commonly (e.g. ruby fluorescence).

334

335 The F-intercepts on the F-f plots suggest a decrease in  $K_{0T}$  with increasing iron content for  
336 samples compressed in helium and 4:1 methanol-ethanol, the pressure media in which the  
337 compression of multiple (Mg,Fe)O compositions has been measured. These trends, with slopes  
338 of about  $-0.15$  GPa/mol%, are shown in Figure 7. Even though the trends have similar slopes, the  
339 methanol-ethanol trend is offset in  $K_{0T}$  by about 3 GPa. It should be noted that the methanol-  
340 ethanol  $K_{0T}$  values are not necessarily expected to be similar to those in other pressure media, as

341 these datasets cannot be fit with a  $K'_{0T}$  near 4, and require a significantly higher value of  $\sim 5.5$   
342 (see Table 3). The best-fit  $K_{0T}$  values of 163.0(1.0) GPa and 166(5) GPa for the cubic-indexed  
343 Mw78 single crystal and powder in neon, respectively, are consistent with each other. They show  
344 essentially the same result of  $K_{0T} = 160(2)$  GPa and  $K'_{0T} = 4.12(14)$  that was reported for  
345 (Mg,Fe)O powder with 48 mol% FeO in neon using a spin-crossover equation of state  
346 (Solomatova et al. 2016).

347

### 348 **Compositional Dependence of the Bulk Modulus (subsection)**

349 Values for  $K_{0T}$  obtained from equation of state fits can also be compared to results using other  
350 methods such as Brillouin spectroscopy and ultrasonic interferometry, which determine sound  
351 velocities rather than density. Sound velocities and density are then combined to calculate the  
352 adiabatic bulk modulus  $K_S$ , which is related to  $K_T$  via  $K_S = K_T (1 + \alpha\gamma T)$ , where  $\alpha$  is the  
353 volumetric thermal expansion coefficient,  $\gamma$  is the Grüneisen parameter, and  $T$  is temperature (in  
354 our case, 298 K). Lack of sufficient data for the compositional dependence of  $\alpha\gamma$  in the  
355 (Mg,Fe)O system makes it unclear what values should be used, and therefore for Brillouin and  
356 ultrasonic studies we are plotting  $K_{0S}$ .

357

358 As reported previously by Jacobsen et al. (2002),  $K_{0S}$  trends differently for solid solutions  
359 containing low and high Fe content. Crystals with FeO mol% up to  $\sim 15\%$  show an increasing  
360  $K_{0S}$  with increasing iron content, albeit with small-number statistics, while those with higher  
361 FeO mol% show the opposite effect. This is in subtle contrast to equation of state fits to pressure-  
362 volume data from diffraction, which support a linear decrease in  $K_{0T}$  from MgO to  $Fe_{1-x}O$ . It  
363 should be noted that this trend is observed for single crystals that likely have an increasing

364 amount of ferric iron towards the  $\text{Fe}_{1-x}\text{O}$  end-member (this study, Jacobsen et al. 2002).  
365 Additional work is needed to reconcile this discrepancy, with a potential focus on the differential  
366 sensitivity of diffraction, ultrasonic, and Brillouin measurements to small proportions of a defect  
367 structure.

368

### 369 **Implications**

370

371 In this paper, we studied the room-temperature elastic behavior of single-crystalline Mw78 in  
372 helium and neon pressure media. We refined the compositional formula to  $(\text{Mg}_{0.215}\text{Fe}_{0.762-0.023})\text{O}$   
373 by constraining the ferric Fe content with high-resolution synchrotron Mössbauer spectroscopy.  
374 As discussed earlier, iron-rich magnesiowüstite may be present in the Earth's core-mantle  
375 boundary (CMB) region. An extrapolation of Mw78 elastic properties to pressure and  
376 temperature conditions near the CMB requires additional information. The density of Mw78  
377 resulting from the present study in helium has to be supplemented with thermal parameters such  
378 as a pressure-dependent Debye temperature and Grüneisen parameter. Here we use  $\theta_{\text{D0}} = 426$  K,  
379  $\gamma_0 = 1.72(8)$ ,  $q = 0.5$  as reported by Wicks et al. (2015) for  $(\text{Mg}_{0.06}\text{Fe}_{0.94})\text{O}$ , a composition similar  
380 to that explored in the current study. We also assume the absence of phase transitions, as iron-  
381 rich magnesiowüstite is expected to be cubic near the CMB (Fischer et al. 2011; Wicks et al.  
382 2015). For the extrapolation of density, we use a Birch-Murnaghan Mie-Grüneisen equation of  
383 state as implemented into the MINUTI software.

384

385 Pressure at the CMB is well established to be 135.8 GPa from the preliminary reference Earth  
386 model (PREM) (Dziewonski and Anderson 1981), however, temperature at the CMB and the

387 temperature profile near the CMB are less well agreed on. Candidate geotherms from van der  
388 Hilst et al. (2007) and Zhang et al. (2016) suggest that the upper bound for the core-mantle  
389 boundary temperature is  $\sim 4000$  K, but the addition of minor amounts of light elements will likely  
390 depress this value. To compute geophysically-relevant parameters of crystalline Mw78 at the  
391 CMB, we assume a CMB temperature of 3800 K (L. Zhang and Fei 2008; Fischer and Campbell  
392 2010; Wicks et al. 2015; Kato et al. 2016). Extrapolating to the CMB pressure of 135.8 GPa at  
393 300 K using the  $V_0$ ,  $K_{0T}$ , and  $K'_{0T}$  parameters obtained from our single-crystal experiments in  
394 neon and helium (Table 3), we obtain a density of  $8.03(3)$  g/cm<sup>3</sup> for the sample in neon (using  
395 the equation of state parameters from the cubic-indexed data set) and  $8.17(3)$  g/cm<sup>3</sup> for the  
396 sample in helium. The  $\sim 2\%$  percent density difference between these two calculations is about a  
397 factor of five higher than the uncertainty in density for the individual single-crystal equation of  
398 state studies presented here. This illustrates the benefits of using an equation of state from high  
399 quality measurements. Applying thermal considerations to compute the density at 3800 K at the  
400 same pressure results in a density of  $7.51$  g/cm<sup>3</sup> using the equation of state of the sample in neon  
401 and  $7.63$  g/cm<sup>3</sup> for the sample in helium.

402

403 To place the value for density computed from our study using a helium pressure medium in the  
404 context of the 1D Earth model PREM, Mw78 is about 37% more dense than the average mantle  
405 at the CMB. On the core side of the CMB, Mw78 is about 23% less dense than the outer core. If  
406 magnesiowüstite with a similar composition to our sample existed above the core-mantle  
407 boundary, it would form morphologies explored in work by Bower et al. (2011) that depend on a  
408 small volume fraction mixed with silicates. It has recently been suggested that giant impacts over  
409 Earth's history could have resulted in the temperature-dependent dissolution of small amounts of

410 Mg in the outer core, which then could have precipitated and floated to the core-mantle boundary  
411 as the core cools, helping to drive the geodynamo over billions of years (O'Rourke and  
412 Stevenson 2016). The precipitated Mg could react either with lower mantle material or any  
413 oxygen present as a light element in the outer core, to form MgO (Badro et al. 2016) or  
414 (Mg,Fe)O. As indicated by the density deficit of  $(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$  compared with PREM, this  
415 material could be emplaced on the core-side of the core-mantle boundary, potentially percolate  
416 or diffuse into the lower mantle, and then stabilize at the base of the mantle.

417

418 To date, only a select number of (Mg,Fe)O compositions have been investigated using single-  
419 crystal diffraction at high pressure, and few Brillouin and ultrasonic studies exist with which to  
420 compare the results. In this paper, we provided constraints on the elastic behavior of Mw78 in  
421 helium and neon pressure media. From the significant differences of the corresponding  
422 compression behavior, we conclude that nonhydrostaticity has a strong influence on the  
423 appearance of distortions away from cubic symmetry in iron-rich (Mg,Fe)O compositions. In  
424 addition, the derived equations of state are clearly different, in particular at the high pressures  
425 relevant to lower mantle conditions. The study of single crystals in a helium pressure medium  
426 presents a valuable baseline upon which future studies can build, for example, by extending the  
427 pressure range and elevating the temperature.

428

#### 429 **Acknowledgements**

430 This work was supported by National Science Foundation (EAR) CSEDI-1161046, CAREER-  
431 0956166, and the Caltech Seismological Laboratory Director's Postdoctoral Fellowship. The  
432 GSECARS gas-loading system, APS Sector 3, and GSECARS PX<sup>2</sup> are partially supported by

433 COMPRES. We are thankful for the single crystal sample provided to us by Stephen J.  
434 Mackwell. We would like to thank Natalia Solomatova, Rachel Morrison, Przemek Dera, Jiyong  
435 Zhao, Wenli Bi, Christine Beavers, Simon Teat, Vitali Prakapenka, Clemens Prescher, and  
436 Sergey Tkachev for their assistance with our experiments.

437

## 438 **References**

439

440 Angel, R.J. (2000) Equations of state. *Reviews in Mineralogy and Geochemistry*, 41, 35–59.

441 Badro, J., Siebert, J., and Nimmo, F. (2016) An early geodynamo driven by exsolution of mantle  
442 components from Earth's core. *Nature*, 536, 326–328.

443 Boehler, R., and De Hantsetters, K. (2007) New anvil designs in diamond-cells. *High Pressure*  
444 *Research*, 24, 391–396.

445 Bower, D.J., Wicks, J.K., Gurnis, M., and Jackson, J.M. (2011) A geodynamic and mineral  
446 physics model of a solid-state ultralow-velocity zone. *Earth and Planetary Science Letters*,  
447 303, 193–202.

448 Bruker (2012a) APEX2. Bruker AXS Inc., Madison, Wisconsin, USA.

449 Bruker (2012b) CELL\_NOW. Bruker AXS Inc., Madison, Wisconsin, USA.

450 Dera, P., Finkelstein, G.J., Duffy, T.S., Downs, R.T., Meng, Y., Prakapenka, V., and Tkachev, S.  
451 (2013a) Metastable high-pressure transformations of orthoferrosilite  $\text{Fs}_{82}$ . *Physics of the*  
452 *Earth and Planetary Interiors*, 221, 15–21.

453 Dera, P., Zhuravlev, K., Prakapenka, V., Rivers, M.L., Finkelstein, G.J., Grubor-Urosevic, O.,  
454 Tschauer, O., Clark, S.M., and Downs, R.T. (2013b) High pressure single-crystal micro X-  
455 ray diffraction analysis with GSE\_ADA/RSV software. *High Pressure Research*, 1–19.

456 Dewaele, A., Loubeyre, P., and Mezouar, M. (2004) Equations of state of six metals above 94  
457 GPa. *Physical Review B*, 70, 094112.

458 Dyar, M.D., Agresti, D.G., Schaefer, M.W., Grant, C.A., and Sklute, E.C. (2006) Mössbauer  
459 spectroscopy of earth and planetary materials. [dx.doi.org](http://dx.doi.org), 34, 83–125.

460 Dziewonski, A.M., and Anderson, D.L. (1981) Preliminary reference Earth model. *Physics of the*  
461 *Earth and Planetary Interiors*, 25, 297–356.

462 Fei, Y., Wang, Y., and Finger, L.W. (1996) Maximum solubility of FeO in  $(\text{Mg}, \text{Fe})\text{SiO}_3$ -  
463 perovskite as a function of temperature at 26 GPa: Implication for FeO content in the lower  
464 mantle. *Journal of Geophysical Research*, 101, 11525.

- 465 Finkelstein, G.J., Dera, P.K., Jahn, S., Oganov, A.R., Holl, C.M., Meng, Y., and Duffy, T.S.  
466 (2014) Phase transitions and equation of state of forsterite to 90 GPa from single-crystal X-  
467 ray diffraction and molecular modeling. *American Mineralogist*, 99, 35–43.
- 468 Fischer, R.A., and Campbell, A.J. (2010) High-pressure melting of wüstite. *American*  
469 *Mineralogist*, 95, 1473–1477.
- 470 Fischer, R.A., Campbell, A.J., Shofner, G.A., Lord, O.T., Dera, P., and Prakapenka, V.B. (2011)  
471 Equation of state and phase diagram of FeO. *Earth and Planetary Science Letters*, 304, 496–  
472 502.
- 473 Hazen, R.M., and Jeanloz, R. (1984) Wüstite (Fe<sub>1-x</sub>O): A review of its defect structure and  
474 physical properties, 22, 37–46.
- 475 Irifune, T. (1994) Absence of an aluminous phase in the upper part of the Earth's lower mantle.
- 476 Jacobsen, S.D. (2002) Structure and elasticity of single-crystal (Mg,Fe)O and a new method of  
477 generating shear waves for gigahertz ultrasonic interferometry. *Journal of Geophysical*  
478 *Research*, 107, 2037.
- 479 Jacobsen, S.D., and Spetzler, H. (2004) Shear waves in the diamond-anvil cell reveal pressure-  
480 induced instability in (Mg,Fe)O. *Proceedings of the National Academy of Sciences of the*  
481 *United States of America*, 101, 5867–5871.
- 482 Jacobsen, S.D., Holl, C.M., Adams, K.A., Fischer, R.A., Martin, E.S., Bina, C.R., Lin, J.F.,  
483 Prakapenka, V.B., Kubo, A., and Dera, P. (2008) Compression of single-crystal magnesium  
484 oxide to 118 GPa and a ruby pressure gauge for helium pressure media. *American*  
485 *Mineralogist*, 93, 1823–1828.
- 486 Jacobsen, S.D., Lin, J.-F., Angel, R.J., Shen, G., Prakapenka, V.B., Dera, P., Mao, H.-K., and  
487 Hemley, R.J. (2005) Single-crystal synchrotron X-ray diffraction study of wüstite and  
488 magnesiowüstite at lower-mantle pressures. *Journal of Synchrotron Radiation*, 12, 577–583.
- 489 Johnson, D.P. (1969) Mössbauer study of the local environments of <sup>57</sup>Fe in FeO. *Solid State*  
490 *Communications*.
- 491 Kantor, I., Prakapenka, V., Kantor, A., Dera, P., Kurnosov, A., Sinogeikin, S., Dubrovinskaia,  
492 N., and Dubrovinsky, L. (2012) BX90: A new diamond anvil cell design for X-ray  
493 diffraction and optical measurements. *Review of Scientific Instruments*, 83.
- 494 Kantor, I.Y., McCammon, C.A., and Dubrovinsky, L.S. (2004) Mössbauer spectroscopic study  
495 of pressure-induced magnetisation in wüstite (FeO). *Journal of Alloys and Compounds*, 376,  
496 5–8.
- 497 Kato, C., Hirose, K., Nomura, R., Ballmer, M.D., Miyake, A., and Ohishi, Y. (2016) Melting in  
498 the FeO-SiO<sub>2</sub> system to deep lower-mantle pressures: Implications for subducted banded  
499 iron formations. *Earth and Planetary Science Letters*, 440, 56–61.

- 500 Klotz, S., Chervin, J.-C., Munsch, P., and Le Marchand, G. (2009) Hydrostatic limits of 11  
501 pressure transmitting media. *Journal of Physics D: Applied Physics*, 42, 075413.
- 502 Lin, J.-F., Gavriluk, A.G., Struzhkin, V.V., Jacobsen, S.D., Sturhahn, W., Hu, M.Y., Chow, P.,  
503 and Yoo, C.-S. (2006) Pressure-induced electronic spin transition of iron in  
504 magnesiowüstite-(Mg,Fe)O. *Physical Review B*, 73, 113107.
- 505 Lin, J.F., Vanko, G., Jacobsen, S.D., Iota, V., Struzhkin, V.V., Prakapenka, V.B., Kuznetsov, A.,  
506 and Yoo, C.S. (2007) Spin transition zone in earth's lower mantle. *Science*, 317, 1740–1743.
- 507 Mackwell, S., Bystricky, M., and Sproni, C. (2005) Fe–Mg interdiffusion in (Mg,Fe)O. *Physics  
508 and Chemistry of Minerals*, 32, 418–425.
- 509 Manga, M., and Jeanloz, R. (1996) Implications of a metal-bearing chemical boundary layer in  
510 D" for mantle dynamics. *Geophysical Research Letters*.
- 511 Mao, H. (1997) Multivariable dependence of Fe-Mg partitioning in the lower mantle. *Science*,  
512 278, 2098–2100.
- 513 Mao, W.L., Mao, H.-K., Sturhahn, W., Zhao, J., Prakapenka, V.B., Meng, Y., Shu, J., Fei, Y.,  
514 and Hemley, R.J. (2006) Iron-rich post-perovskite and the origin of ultralow-velocity zones.  
515 *Science*, 312, 564–565.
- 516 Mao, Z., Lin, J.-F., Liu, J., and Prakapenka, V.B. (2011) Thermal equation of state of lower-  
517 mantle ferropericlase across the spin crossover. *Geophysical Research Letters*, 38.
- 518 McCammon, C.A., and Price, D.C. (1985) Mössbauer spectra of  $\text{Fe}_x\text{O}$  ( $x > 0.95$ ). *Physics and  
519 Chemistry of Minerals*, 11, 250–254.
- 520 Ohta, K., Fujino, K., Kuwayama, Y., Kondo, T., Shimizu, K., and Ohishi, Y. (2014) Highly  
521 conductive iron-rich (Mg,Fe)O magnesiowüstite and its stability in the Earth's lower mantle.  
522 *Journal of Geophysical Research (Solid Earth)*, 119, 4656–4665.
- 523 O'Rourke, J.G., and Stevenson, D.J. (2016) Powering Earth's dynamo with magnesium  
524 precipitation from the core. *Nature*, 529, 387–389.
- 525 Rivers, M., Prakapenka, V.B., Kubo, A., Pullins, C., Holl, C.M., and Jacobsen, S.D. (2008) The  
526 COMPRES/GSECARS gas-loading system for diamond anvil cells at the Advanced Photon  
527 Source. *High Pressure Research*, 28, 273–292.
- 528 Rost, S., Garnero, E.J., and Williams, Q. (2006) Fine-scale ultralow-velocity zone structure from  
529 high-frequency seismic array data. *Journal of Geophysical Research*, 111, B09310.
- 530 Saines, P.J., Tucker, M.G., Keen, D.A., Cheetham, A.K., and Goodwin, A.L. (2013) Coupling of  
531 the local defect and magnetic structure of wüstite  $\text{Fe}_{1-x}\text{O}$ . *Physical Review B*, 88, 134418–  
532 8.
- 533 Shu, J., Mao, H., Hu, J., and Fei, Y. (1998) Single-crystal X-ray diffraction of wüstite to 30 GPa

- 534 hydrostatic pressure. *Neues Jahrbuch für Mineralogie Abhandlungen*.
- 535 Sinmyo, R., Hirose, K., Nishio Hamane, D., Seto, Y., Fujino, K., Sata, N., and Ohishi, Y. (2008)  
536 Partitioning of iron between perovskite/postperovskite and ferropericlae in the lower  
537 mantle. *Journal of Geophysical Research (Solid Earth)*, 113, B11204.
- 538 Solomatova, N.V., Jackson, J.M., Sturhahn, W., Wicks, J.K., Zhao, J., Toellner, T.S., Kalkan, B.,  
539 and Steinhardt, W.M. (2016) Equation of state and spin crossover of (Mg,Fe)O at high  
540 pressure, with implications for explaining topographic relief at the core-mantle boundary.  
541 *American Mineralogist*, 101, 1084–1093.
- 542 Struzhkin, V.V., Mao, H.-K., Hu, J., Schwoerer-Böhning, M., Shu, J., Hemley, R.J., Sturhahn,  
543 W., Hu, M.Y., Alp, E.E., Eng, P., and others (2001) Nuclear inelastic x-ray scattering of FeO  
544 to 48 GPa. *Physical Review Letters*, 87, 255501–4.
- 545 Sturhahn, W. (2015a) CONUSS (COherent NUclear resonant Scattering by Single crystals) open  
546 source software. [www.nrixs.net](http://www.nrixs.net).
- 547 Sturhahn, W. (2015b) MINUTI (MINeral physics UTILities) open source software.  
548 [www.nrixs.net](http://www.nrixs.net).
- 549 Sun, D., Helmberger, D.V., Jackson, J.M., Clayton, R.W., and Bower, D.J. (2013) Rolling hills  
550 on the core–mantle boundary. *Earth and Planetary Science Letters*, 361, 333–342.
- 551 Thorne, M.S., and Garnero, E.J. (2004) Inferences on ultralow-velocity zone structure from a  
552 global analysis of SPdKS waves. *Journal of Geophysical Research*.
- 553 Toellner, T.S. (2000) Monochromatization of synchrotron radiation for nuclear resonant  
554 scattering experiments. *Hyperfine Interactions*, 125, 3–28.
- 555 Tschauner, O., Ma, C., Beckett, J.R., Prescher, C., Prakapenka, V.B., and Rossman, G.R. (2014)  
556 Discovery of bridgmanite, the most abundant mineral in Earth, in a shocked meteorite.  
557 *Science*, 346, 1100–1102.
- 558 Van der Hilst, R.D., De Hoop, M.V., Wang, P., Shim, S.H., Ma, P., and Tenorio, L. (2007)  
559 Seismostratigraphy and thermal structure of Earth's core-mantle boundary region. *Science*,  
560 315, 1813–1817.
- 561 Welberry, T.R., and Christy, A.G. (1997) Defect distribution and the diffuse X-ray diffraction  
562 pattern of wustite, Fe<sub>1-x</sub>O. *Physics and Chemistry of Minerals*, 24, 24–38.
- 563 Wen, L. (1998) Ultra-Low Velocity Zones Near the Core-Mantle Boundary from Broadband  
564 PKP Precursors. *Science*, 279, 1701–1703.
- 565 Westrenen, W.V., Li, J., Fei, Y., Frank, M.R., Hellwig, H., Komabayashi, T., Mibe, K., Minarik,  
566 W.G., Orman, J.A.V., Watson, H.C., and others (2005) Thermoelastic properties of  
567 (Mg<sub>0.64</sub>Fe<sub>0.36</sub>)O ferropericlae based on in situ X-ray diffraction to 26.7GPa and 2173K.  
568 *Physics of the Earth and Planetary Interiors*, 151, 163–176.

- 569 Wicks, J.K., Jackson, J.M., and Sturhahn, W. (2010) Very low sound velocities in iron-rich  
570 (Mg,Fe)O: Implications for the core-mantle boundary region. *Geophysical Research Letters*,  
571 37.
- 572 Wicks, J.K., Jackson, J.M., Sturhahn, W., and Zhang, D. (2017) Sound velocity and density of  
573 magnesiowüstites: Implications for ultralow-velocity zone topography. *Geophysical*  
574 *Research Letters*.
- 575 Wicks, J.K., Jackson, J.M., Sturhahn, W., Zhuravlev, K.K., Tkachev, S.N., and Prakapenka, V.B.  
576 (2015) Thermal equation of state and stability of (Mg<sub>0.06</sub>Fe<sub>0.94</sub>)O. *Physics of the Earth and*  
577 *Planetary Interiors*, 249, 28–42.
- 578 Williams, Q., Revenaugh, J., and Garnero, E. (1998) A correlation between ultra-low basal  
579 velocities in the mantle and hot spots. *Science*, 281, 546–549.
- 580 Zhang, D., Jackson, J.M., Zhao, J., Sturhahn, W., Alp, E.E., Hu, M.Y., Toellner, T.S., Murphy,  
581 C.A., and Prakapenka, V.B. (2016) Temperature of Earth's core constrained from melting of  
582 Fe and Fe<sub>0.9</sub>Ni<sub>0.1</sub> at high pressures. *Earth and Planetary Science Letters*, 447, 72–83.
- 583 Zhang, D., Jackson, J.M., Zhao, J., Sturhahn, W., Alp, E.E., Toellner, T.S., and Hu, M.Y. (2015)  
584 Fast temperature spectrometer for samples under extreme conditions. *Review of Scientific*  
585 *Instruments*, 86, 013105.
- 586 Zhang, L., and Fei, Y. (2008) Melting behavior of (Mg,Fe)O solid solutions at high pressure.  
587 *Geophysical Research Letters*, 35.
- 588 Zhuravlev, K.K., Jackson, J.M., Wolf, A.S., Wicks, J.K., Yan, J., and Clark, S.M. (2010)  
589 Isothermal compression behavior of (Mg,Fe)O using neon as a pressure medium. *Physics*  
590 *and Chemistry of Minerals*, 37, 465–474.

591

## 592 **Figure Captions**

593

594 Figure 1. Pseudo-precession *hk0* image of Mw78 diffraction at ambient conditions collected in  
595 air at ALS. Weak satellite peaks and diffuse scattering can be seen around the main Bragg  
596 reflections.

597

598 Figure 2. Synchrotron Mössbauer spectrum of Mw78 single crystal at room pressure.

599 The raw data and associated  $1\sigma$  error bars are shown as light-gray markers, and the best-fit  
600 model is shown as a purple line. The inset shows the calculated energy spectrum using the best-  
601 fit parameters from the measured time spectrum and the contributions to the fit from each  
602 individual site (pink, red, blue, and green correspond to the  $\text{Fe}_A^{2+}$ ,  $\text{Fe}_B^{2+}$ ,  $\text{Fe}_C^{2+}$ , and  $\text{Fe}_D^{3+}$  sites  
603 described in Table 1). In this calculation the absolute isomer shift is arbitrary, and only the  
604 relative isomer shift is meaningful.

605

606 Figure 3. a) Top: Ruby standard deviation as a function of mean ruby pressure for both helium-  
607 (orange circles) and neon-loaded (cyan circles) diamond anvil cells. Bottom:  $c/a$  ratio as a  
608 function of mean ruby pressure for both helium- (orange circles) and neon-loaded (cyan circles)  
609 diamond anvil cells. Uncertainty in pressure is represented as one ruby standard deviation. Error  
610 bars are not given for the ruby standard deviation, and, when not visible, other error bars are  
611 smaller than the symbol size. Uncertainty in pressure is represented as one ruby standard  
612 deviation, and error in  $c/a$  ratio represents a  $1\sigma$  uncertainty that has been propagated from lattice  
613 parameter fits. When not visible, error bars are smaller than the symbol size. b) The reciprocal  
614 lattice of the single-crystal in neon at 53.3 GPa. Red vectors outline the hexagonal unit cell and  
615 its relationship to a pseudo-cubic unit cell.

616

617 Figure 4. Normalized unit cell volumes of Mw78 and equations of state fit using MINUTI  
618 (Sturhahn 2015b). Orange, dark cyan/green, and light cyan symbols correspond to diffraction  
619 from a single crystal of Mw78 in helium (this study, indexed using a cubic unit cell), a single  
620 crystal of Mw78 in neon (this study, indexed using a cubic unit cell below and a hexagonal unit  
621 cell above 20 GPa), and powdered  $(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$  in neon (Zhuravlev et al. 2010), respectively.

622 Circles represent measured unit cell volumes normalized by the zero-pressure unit cell volume  
623 resulting from a fit by a 3<sup>rd</sup> order Birch-Murnaghan equation of state. Uncertainty in pressure is  
624 represented as one ruby standard deviation. Shaded regions represent the fit 3<sup>rd</sup> order Birch-  
625 Murnaghan equations of state with 1 $\sigma$  fitting errors. Where not visible, vertical error bars on  
626 individual data points are smaller than the symbol size.

627

628 Figure 5. a) Confidence ellipses of  $K_T$  and  $K'_T$  at 68% and 95% levels for several (Mg,Fe)O  
629 compositions at 0, 20, and 130 GPa. Orange, cyan, and brown ellipses represent diffraction  
630 measurements in helium, neon, and 4:1 methanol-ethanol pressure media (Jacobsen 2002;  
631 Jacobsen et al. 2005; 2008; Zhuravlev et al. 2010), respectively. Darker ellipses correspond to  
632 samples from this study (the ellipses corresponding to the sample in Ne are derived from the  
633 covariance matrix of the equation of state fit to the low-pressure cubic-indexed data points). b)  
634 Confidence ellipses of  $V$  and  $K_T$ . c) Confidence ellipses of  $V$  and  $K'_T$ .

635

636 Figure 6. Plots of normalized pressure ( $F$ ) as a function of Eulerian strain ( $f$ ) corresponding to  
637 compression data (filled circles with error bars) and equation of state fits (solid lines) for various  
638 (Mg,Fe)O compositions (Jacobsen 2002; Jacobsen et al. 2005; 2008; Zhuravlev et al. 2010). Top:  
639 Single-crystal studies in a helium pressure medium. Middle. Single-crystal and powder studies in  
640 a neon pressure medium. Dark green and dark cyan data points are from the same sample, and  
641 represent fits to the portions of the data that can be indexed using cubic and hexagonal unit cells,  
642 respectively. For the data from the samples in a neon pressure medium, a strong prior was placed  
643 on  $K'_{OT}$ , causing it to stay near a value of  $4.0 \pm 0.1$  during fitting. Bottom: Single-crystal studies  
644 in a 4:1 methanol-ethanol pressure medium. Dark symbols/lines correspond to data collected for

645 this study and light symbols/lines correspond to previous studies. The error bars in F and f  
646 correspond to  $1\sigma$  uncertainties, and when not visible, are smaller than the symbol size (see Angel  
647 (2000) for details on F-f plot error bar calculations). Uncertainties are not given for the lines  
648 corresponding to the equation of state fits.

649

650 Figure 7.  $K_{0T}$  (diffraction) and  $K_{0S}$  (ultrasonic, Brillouin) compositional trends for (Mg,Fe)O  
651 solid solutions from single-crystal studies. Ultrasonic and Brillouin studies are represented by  
652 red and purple symbols, respectively. Single-crystal diffraction studies in helium, neon, and 4:1  
653 methanol-ethanol pressure media are represented by orange, cyan, and brown symbols,  
654 respectively. The dashed lines show weighted linear least-square fits to the data. In the equations  
655 that correspond to the best-fit lines,  $x$  is the iron mol% and  $y$  is either  $K_{0T}$  (diffraction) or  $K_{0S}$   
656 (ultrasonic, Brillouin) in units of GPa. Compositional error bars are smaller than the symbol size  
657 and  $K_0$  error bars are either  $1\sigma$  uncertainties from equation of state fitting (diffraction) or those  
658 reported in the literature (Brillouin, ultrasonic).

659

**Table 1. Mw78 room-pressure best-fit hyperfine parameters**

Site	Coordination Environment	Weight Fraction (%)	IS (mm/s)	QS (mm/s)	FWHM (mm/s)
Fe <sub>A</sub> <sup>2+</sup>	Octahedral	69.7(1.0)	1.0 <sup>b</sup>	0.592(3)	0.1319(14) <sup>c</sup>
Fe <sub>B</sub> <sup>2+</sup>	Octahedral	17.3(1.0)	1.0 <sup>b</sup>	0.913(5)	0.1319(14) <sup>c</sup>
Fe <sub>C</sub> <sup>2+</sup>	Octahedral	7 <sup>a</sup>	1.0 <sup>b</sup>	1.307(7)	0.1319(14) <sup>c</sup>
Fe <sub>D</sub> <sup>3+</sup>	Tetrahedral	6 <sup>a</sup>	0.044(10)	0	0.427(12)

Thickness = 137.6(1.3)  $\mu\text{m}$ , Lamb Mössbauer Factor = 0.75

IS = Isomer Shift, QS = Quadrupole Splitting, FWHM = Full-Width at Half Maximum

<sup>a</sup>These site fractions were manually varied to optimize the fit.

<sup>b</sup>The isomer shift for these sites was fixed so that the relative isomer shift for the ferric Fe site could be reported.

<sup>c</sup>The widths of the doublet sites were constrained to be identical to each other.

**Table 2. Lattice parameters and unit cell volumes of Mw78**

Ambient – Cubic Unit Cell				
P (GPa)	<i>a</i> (Å)	V (Å <sup>3</sup> )		
0.0001	4.2898(4)*	78.94(2)		
Helium Medium – Cubic Unit Cell				
Ruby 1 λ (nm)	Ruby 2 λ (nm)	Avg. P (GPa)	<i>a</i> (Å)	V (Å <sup>3</sup> )
694.95	694.95	1.823(16)	4.2703(12)	77.87(7)
696.13	696.14	5.128(32)	4.2433(12)	76.40(7)
697.41	697.45	8.786(63)	4.2132(12)	74.79(7)
698.64	698.67	12.294(64)	4.1860(13)	73.35(7)
699.98	700.01	16.199(65)	4.1596(13)	71.97(7)
701.02	701.06	19.301(72)	4.1420(12)	70.97(7)
702.29	702.33	23.119(75)	4.1152(15)	69.69(8)
703.38	703.44	26.467(115)	4.1003(11)	68.94(6)
704.22	704.28	29.044(136)	4.0838(14)	68.11(7)
705.13**	705.15	31.820(52)	4.0707(13)	67.46(7)
706.40	706.44	35.856(86)	4.0491(13)	66.38(7)
707.70	707.74	40.004(94)	4.0314(15)	65.52(7)
708.40	708.46	42.319(127)	4.0235(13)	65.12(7)
709.73	709.77	46.645(71)	4.0034(11)	64.16(5)
710.78	710.84	50.177(145)	3.9891(15)	63.48(8)
711.48	711.53	52.517(129)	3.9781(14)	62.95(7)
712.35	712.44	55.542(222)	3.9665(16)	62.40(8)
Helium Medium – Hexagonal Unit Cell				
Avg. P (GPa)	<i>a</i> (Å)	<i>c</i> (Å)	<i>c/a</i>	V (Å <sup>3</sup> )
1.823(16)	3.0184(17)	7.404(10)	2.453(4)	58.42(10)
5.128(32)	3.0011(16)	7.345(9)	2.448(4)	57.29(9)
8.786(63)	2.9778(16)	7.307(10)	2.454(4)	56.11(10)
12.294(64)	2.9592(17)	7.255(10)	2.452(4)	55.02(10)
16.199(65)	2.9388(14)	7.221(9)	2.457(4)	54.01(8)
19.301(72)	2.9265(16)	7.179(10)	2.453(4)	53.24(9)
23.119(75)	2.9063(18)	7.151(10)	2.461(4)	52.31(10)
26.467(115)	2.8994(14)	7.102(8)	2.449(3)	51.70(8)
29.044(136)	2.8858(19)	7.086(11)	2.455(5)	51.10(11)
31.820(52)	2.8761(17)	7.067(10)	2.457(4)	50.63(10)
35.856(86)	2.8620(17)	7.021(10)	2.453(4)	49.80(10)
40.004(94)	2.8490(19)	6.993(11)	2.455(5)	49.16(11)
42.319(127)	2.8418(16)	6.991(10)	2.460(4)	48.89(9)
46.645(71)	2.8295(13)	6.953(15)	2.457(6)	48.21(11)
50.177(145)	2.820(2)	6.914(12)	2.452(5)	47.62(11)
52.517(129)	2.8110(19)	6.903(11)	2.456(5)	47.24(10)
55.542(222)	2.8008(19)	6.898(12)	2.463(5)	46.86(10)
Neon Medium – Cubic Unit Cell				

Ruby 1 $\lambda$ (nm)	Ruby 2 $\lambda$ (nm)	Avg. P (GPa)	$a$ (Å)	$V$ (Å <sup>3</sup> )
694.76	694.77	1.314(20)	4.2749(5)	78.12(3)
696.19	696.20	5.294(20)	4.2423(5)	76.35(3)
697.51	697.52	9.029(21)	4.2162(5)	74.95(3)
699.12	699.14	13.682(42)	4.1827(5)	73.18(3)
701.02	701.02	19.244	4.1505(6)	71.5(4)
702.73	702.53	24.084(43)	4.1246(8)	70.17(5)
704.45	704.07	29.081(84)	4.1023(13)	69.04(7)
706.12	705.66	34.18(1.03)	4.0806(15)	67.95(8)
708.20	707.68	40.73(1.19)	4.0546(19)	66.66(10)
709.14	708.66	43.86(1.12)	4.039(3)	65.90(15)
710.26	709.73	47.46(1.25)	4.028(3)	65.36(15)
712.03	711.43	53.28(1.44)	4.003(4)	65.14(20)

Neon Medium – Hexagonal Unit Cell

Avg. P (GPa)	$a$ (Å)	$c$ (Å)	$c/a$	$V$ (Å <sup>3</sup> )
1.314(20)	3.0240(7)	7.409(3)	2.4500(12)	58.68(4)
5.294(20)	3.0000(8)	7.347(3)	2.4490(12)	57.26(4)
9.029(21)	2.9813(8)	7.304(3)	2.4499(13)	56.22(4)
13.682(42)	2.9581(7)	7.243(3)	2.4485(12)	54.89(4)
19.244	2.9318(8)	7.201(3)	2.4562(13)	53.60(4)
24.084(43)	2.9102(9)	7.170(4)	2.4637(16)	52.59(5)
29.081(84)	2.8891(12)	7.152(5)	2.476(3)	51.79(6)
34.18(1.03)	2.8720(13)	7.122(5)	2.480(3)	50.88(6)
40.73(1.19)	2.8508(16)	7.086(6)	2.486(3)	49.87(8)
43.86(1.12)	2.833(2)	7.092(9)	2.503(4)	49.29(10)
47.46(1.25)	2.8214(18)	7.084(7)	2.511(3)	48.83(8)
53.28(1.44)	2.7953(18)	7.070(7)	2.529(3)	47.84(8)

\* Using the relationships  $a_{hex} = a_{cubic}/\sqrt{2}$  and  $c_{hex} = \sqrt{3} a_{cubic}$  to transform cubic to hexagonal resulting in  $a_{hex} = 3.0333$  Å and  $c_{hex} = 7.4302$  Å and  $V_{hex} = 59.21$  Å<sup>3</sup>.

\*\*Ruby 1 was not measured at this pressure point. To estimate a pressure comparable with the other data points, we used the other data points to fit the ruby 1 wavelength as a linear function of ruby 2 wavelength. The resulting equation, which we used to estimate the ruby 1 pressure ( $P_{ruby1}$ ), was:  $P_{ruby1} = 0.9973 * P_{ruby2} + 1.8804$  (GPa).  $P_{ruby1}$  and  $P_{ruby2}$  are in units of GPa.

**Table 3. Selected 300 K (Mg,Fe)O Equation of State Fit Parameters**

Work**	Composition	Medium	Technique	P Range (GPa)	$V_0$ ( $\text{\AA}^3$ )	$K_{0T}$ (GPa)	$K'_{0T}$
Jacobsen et al. (2005)	$\text{Fe}_{0.93}\text{O}$	Helium	SXRD	0-22.8	79.41(4)	145.7(1.2)	3.99(6) <sup>b</sup>
				Published Fit:	79.41(4)	146(2)	4.0 (fixed)
This Study (cubic)	$(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$	Helium	SXRD	1.8-55.5	78.87(6) <sup>a</sup>	148(3)	4.09(12)
This Study (cubic)	$(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$	Neon	SXRD	1.3-19.2	78.742(14) <sup>a</sup>	163.0(1.0)	4.02(10) <sup>b</sup>
This Study (hex.)	$(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$	Neon	SXRD	24.1-53.3	58.7(4)	176.8(1.1)	4.00(10) <sup>b</sup>
Zhuravlev et al. (2010)	$(\text{Mg}_{0.22}\text{Fe}_{0.78})\text{O}$	Neon	PXRD	7.9-39.8	78.5(2)	166(5)	4.01(10) <sup>b</sup>
				Published Fit:	78.49(21)	166(5)	4.0 (fixed)
Jacobsen et al. (2002)	$(\text{Mg}_{0.25}\text{Fe}_{0.75})\text{O}$	4:1 Meth-Eth	SXRD	0-7.2	78.082(3)	151.3(7)	5.6(2)
				Published Fit:	78.082(3)	151.3(6)	5.55(19)
Jacobsen et al. (2002)	$(\text{Mg}_{0.44}\text{Fe}_{0.56})\text{O}$	4:1 Meth-Eth	SXRD	0-8.9	77.457(4)	155.2(8)	5.7(3)
				Published Fit:	77.453(4)	155.8(9)	5.5(2)
Jacobsen et al. (2002)	$(\text{Mg}_{0.73}\text{Fe}_{0.27})\text{O}$	4:1 Meth-Eth	SXRD	0-9.3	76.336(3)	158.4(5)	5.48(11)
				Published Fit:	76.336(3)	158.4(4)	5.49(11)
Jacobsen et al. (2005)	$(\text{Mg}_{0.73}\text{Fe}_{0.27})\text{O}$	Helium	SXRD	0-51.1	77.30(4)	154.8(1.9)	3.97(10)
				Published Fit:	76.30(9)	154(3)	4.0(1)
Jacobsen et al. (2008)	MgO	Helium	SXRD	0-118.1	74.707(5)	160.08(14)	4.030(4)
				Published Fit:	74.697(6)	160.2*	4.03*

\* From (2000) MgO primary pressure scale

\*\* If symmetry is not noted, the datasets represent cubic forms of (Mg,Fe)O

<sup>a</sup> Prior of  $V_0 = 78.94 \pm 0.1 \text{ \AA}^3$  used during equation of state fit

<sup>b</sup> Prior of  $K'_{0T} = 4.0 \pm 0.1$  used during equation of state fit

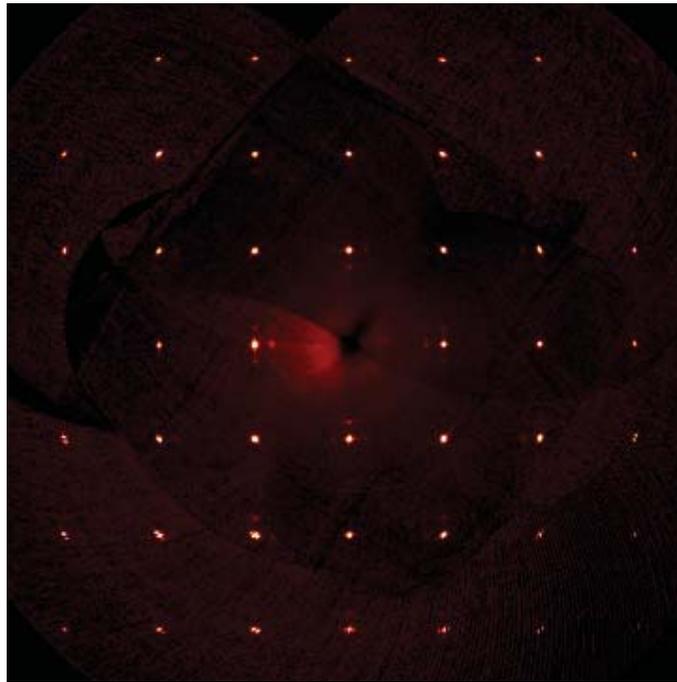


Figure 1

Figure 2

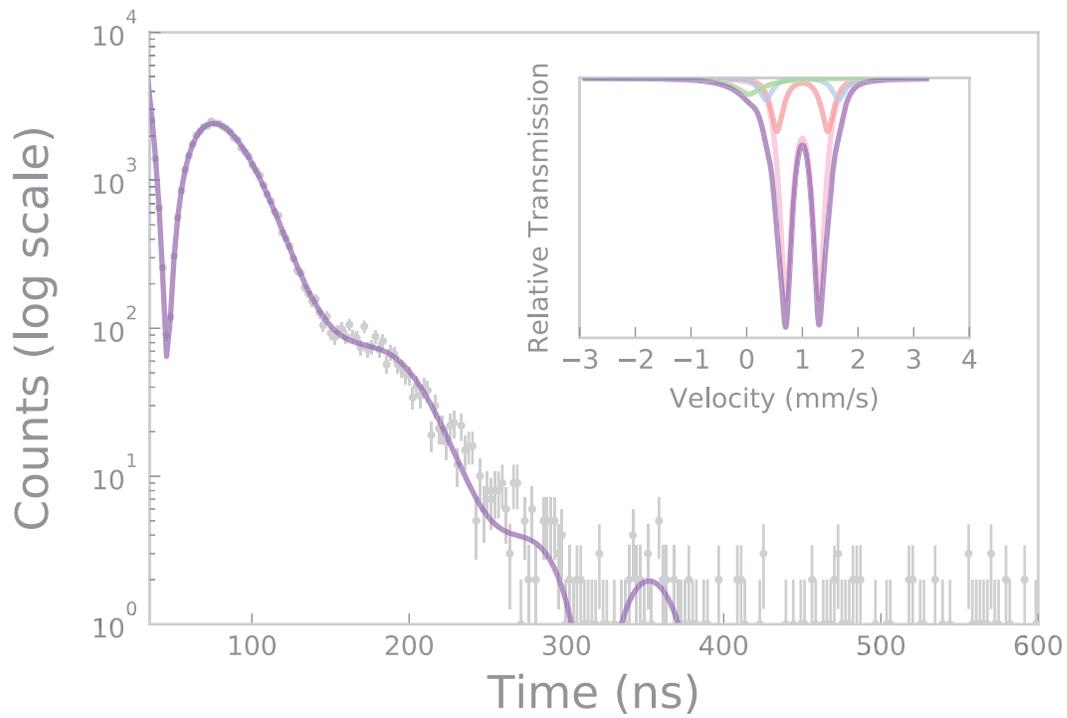


Figure 3a

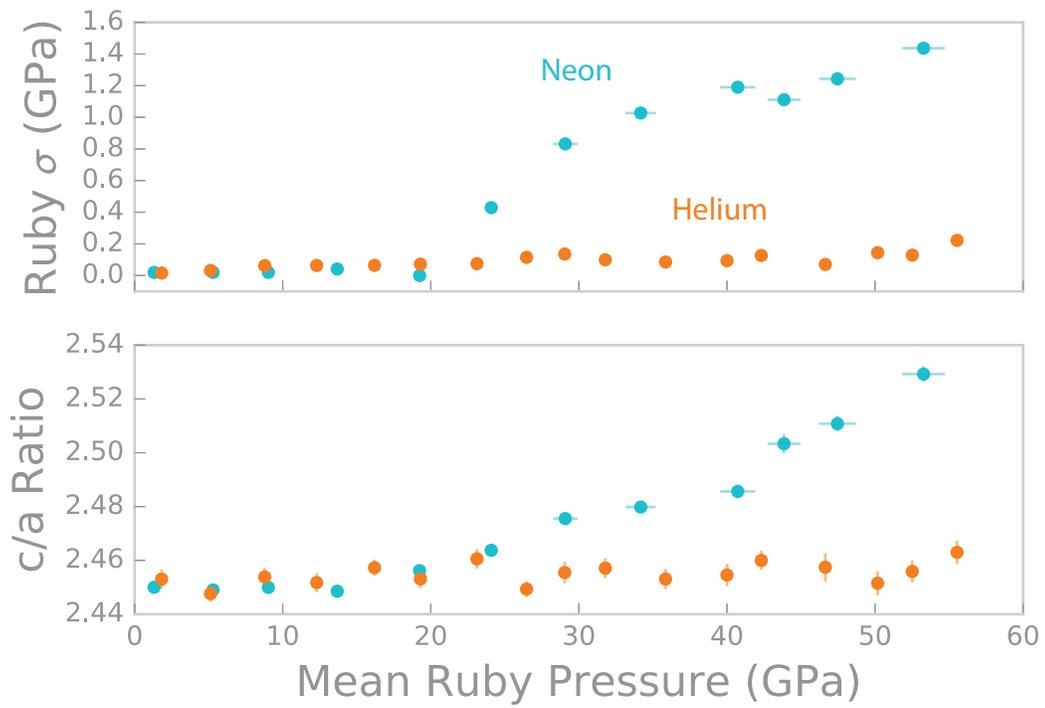


Figure 3b

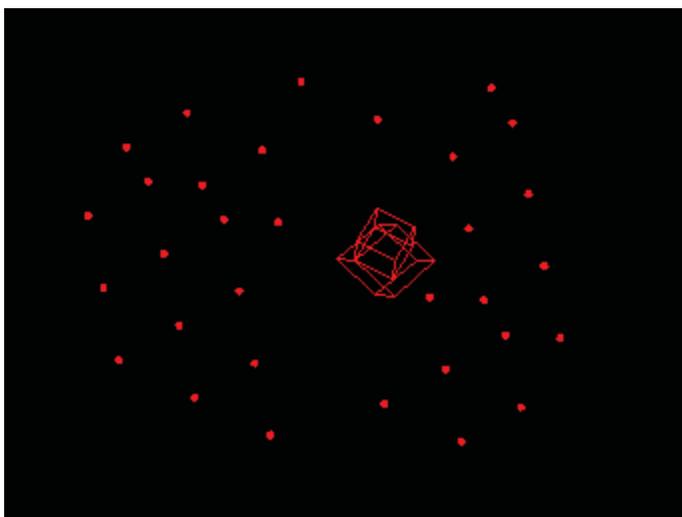
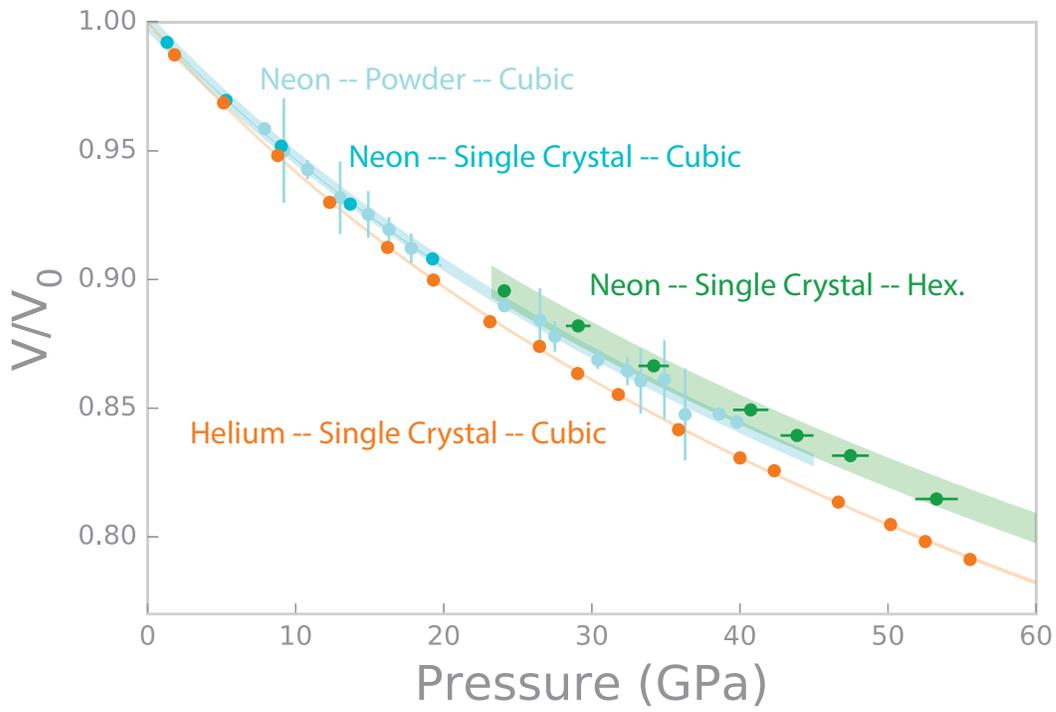


Figure 4



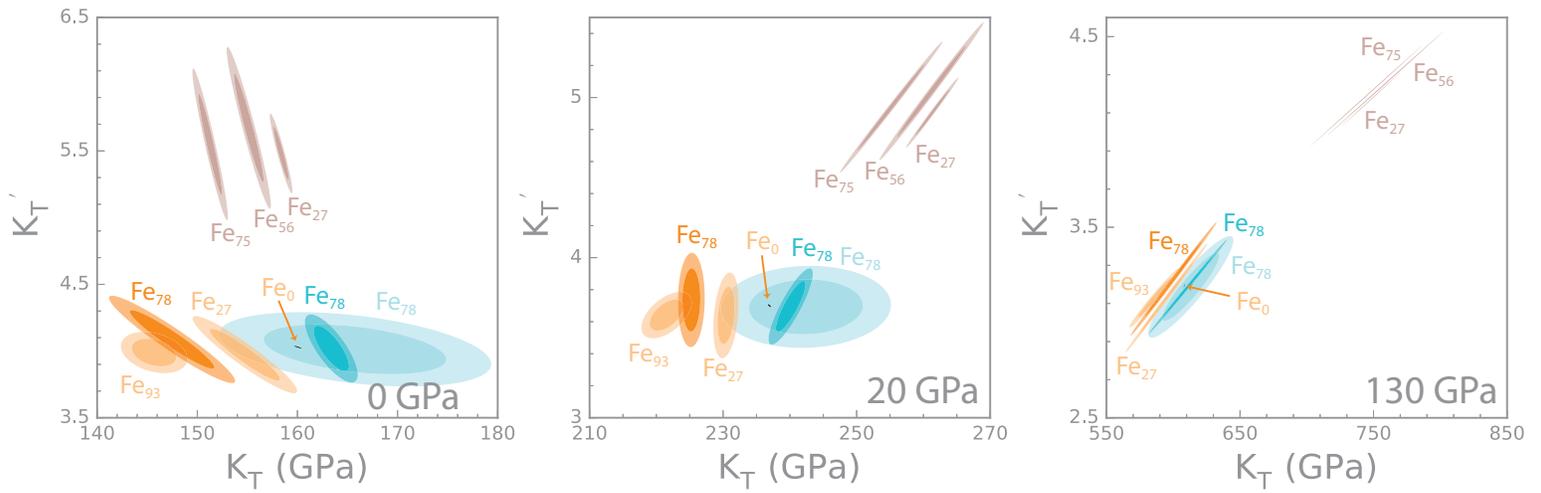


Figure 5a

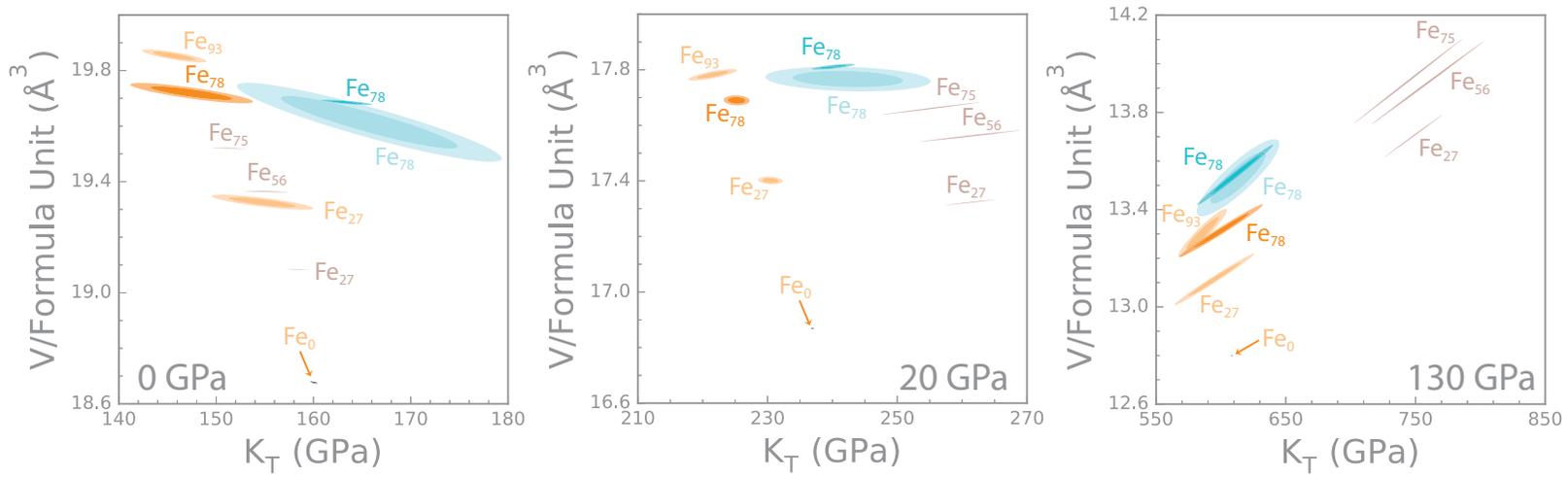


Figure 5b

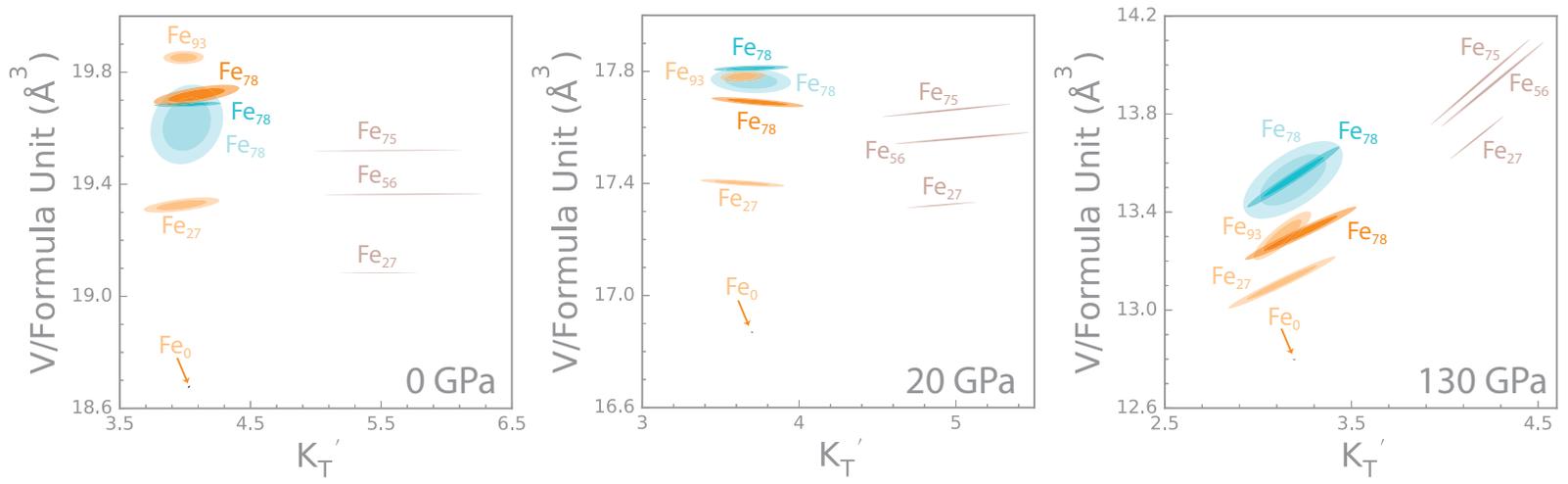


Figure 5c

Figure 6

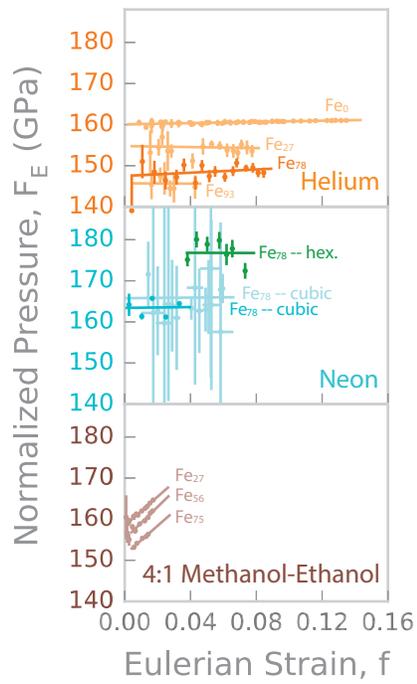


Figure 7

