

1 **REVISION 1**

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3 **Towards the wider application of ^{29}Si NMR spectroscopy to paramagnetic**

4 **transition metal silicate minerals: copper(II) silicates**

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14 **Running title:**

15 “NMR of paramagnetic copper silicates”

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ABSTRACT

20 ^{29}Si NMR has only rarely been applied to silicate minerals in which the predominant
21 cations have unpaired electron spins (e.g. most transition metals and REE), because of the
22 potential for serious line broadening and signal loss. However, as shown here, spectra for a series
23 of natural and synthetic copper(II) silicate minerals can be readily obtained, have paramagnetic
24 shifts far outside known chemical shift ranges, and potentially are very sensitive to structural
25 details involving interactions of paramagnetic cations and Si sites. Signals from different silicon
26 sites in the structures can be distinguished and quantified. Peak broadening due to magnetic
27 couplings and to disorder can be large, but not to the point of ‘non-observability’. NMR signal
28 loss can be related to specific, and in some cases improvable, technical issues such as excitation
29 bandwidth, sample spinning speed, and rapid nuclear spin relaxation. Two samples of the
30 ‘mineraloid’ chrysocolla from different copper ore deposits have very similar spectra with
31 significant paramagnetic shifts, suggesting strong Si-Cu interactions and a common
32 stoichiometry and short-range structure.

33 **Keywords:**

34 nuclear magnetic resonance, unpaired electron spins, diopside, shattuckite, cuprorivaite,
35 planchéite, chrysocolla

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INTRODUCTION

38 Since the earliest applications of high resolution, magic-angle spinning (MAS) ^{29}Si NMR
39 to minerals in the late 1970’s and early 1980’s, nearly all studies have been on materials with
40 low (< a few %) contents of paramagnetic cations, whose unpaired electron spins can interact
41 strongly with nuclear spins and can cause severe NMR line broadening, sometimes to the point

42 of making signal difficult or even impossible to observe with conventional methods (Engelhardt
43 and Michel 1987; Grimmer et al. 1983; MacKenzie and Smith 2002). This problem (actual or
44 expected) has generally limited application of this powerful experimental technique to groups of
45 minerals with low natural contents of transition metal and rare earth cations (e.g. feldspars,
46 feldspathoids, zeolites, silica polymorphs, clays, white micas, etc.) or to end members of
47 ferromagnesian solid solutions, often synthetic, with low or no Fe^{2+} or other transition metal
48 cations (e.g. olivines, pyroxenes, amphiboles, high pressure mantle phases, etc.).

49 However, paramagnetic interactions can also cause large, but observable, changes in
50 NMR peak positions far outside the ranges for diamagnetic materials, as has been particularly
51 well-studied for $^{6,7}\text{Li}$ and ^{31}P NMR in Li-rich transition metal oxides and phosphates of major
52 interest in advanced battery materials (Carlier et al. 2003; Grey and Dupré 2004; Tucker et al.
53 2002). In such materials, despite paramagnetic shifts of thousands of ppm and sometimes severe
54 line broadening, both simple MAS NMR and more advanced methods have revealed important
55 new details about ordering of cations, bonding, site symmetries, and other details of short range
56 structure (Strobridge et al. 2014; Yoon et al. 2004; Zeng et al. 2007). This extended work has led
57 to considerable advances in the theoretical understanding of the complex physics behind such
58 interactions, as well as a rapidly developing ability to calculate paramagnetic shifts from first
59 principles (Clément et al. 2012; Middlemiss et al. 2013). ^{31}P MAS NMR for many pure-phase
60 rare earth monazite-type or zircon-type phosphates are also readily observable, again with
61 paramagnetic shifts of up to 100's of ppm depending on the number of unpaired spins and details
62 of the electronic structure (Bose et al. 1979; Bregiroux et al. 2007). Pioneering studies of ^{119}Sn
63 and ^{89}Y MAS NMR in rare earth stannate pyrochlores introduced this approach to
64 mineralogically interesting oxides (Grey et al. 1989; Grey et al. 1990); very recent studies have

65 reported highly shifted ^{25}Mg (Lee et al. 2017) and ^{27}Al (Dogan et al. 2016) resonances in spectra
66 in doped transition metal oxide battery materials and in ^{89}Y and ^{137}Ba NMR data for transition
67 metal perovskite phases (Michaelis et al. 2012). Of particular interest in geochemistry are studies
68 of sorbed lithium and phosphate on iron oxyhydroxides, where ^7Li and ^{31}P NMR spectra are
69 strongly affected by the host Fe^{3+} cations (Kim et al. 2011; Kim et al. 2008).

70 In contrast, very few NMR studies of pure-phase paramagnetic transition metal silicates
71 have been reported with a rare exception of early work on Co_2SiO_4 (Saji et al. 1973). However,
72 recent reports on synthetic olivine and natural garnet solid solutions containing 0.1 to 20% FeO,
73 NiO, or CoO have shown that, at least in such fairly dilute systems, useful resolution can be
74 observed for Si sites with one or two paramagnetic first-neighbor cations, with shifts up to about
75 400 ppm above the known chemical shift range of -60 to -120 ppm (Begaudeau et al. 2012;
76 McCarty et al. 2015; Palke and Stebbins 2011a; Palke et al. 2015; Stebbins and Kelsey 2009;
77 Stebbins et al. 2017).

78 Such paramagnetic shifts in NMR peaks can be either positive or negative, and are
79 caused by a through-bond transfer of unpaired electron spin density (“Fermi contact shift”)
80 and/or a through-space dipolar coupling from a magnetic cation in an asymmetrical site
81 (“pseudocontact shift”) (Grey et al. 1989; Grey et al. 1990; Bertini et al. 2002; MacKenzie and
82 Smith 2002). These two mechanisms have been recently discussed in the mineralogical context
83 of rare-earth containing phosphate minerals and garnets (Palke and Stebbins 2011a,b; George et
84 al. 2013; McCarty and Stebbins 2016). The Fermi contact shift (also known as the “transferred
85 hyperfine coupling” when the unpaired electron spin is associated with an ion different from that
86 under observation by NMR) involves perturbation of the electron spin density by transfer
87 through one or more chemical bonds, and hence depends strongly on the degree of covalency in

88 the bonds (hence, their length) between the paramagnetic ion and the observed NMR nuclide.
89 This interaction can sometimes be detected several bonds away, but falls off rapidly with the
90 number of intervening bonds. It depends strongly as well on the nature and extent of orbital
91 overlap, which in turn are affected by the intervening bond angles and coordination
92 environments for both the paramagnetic center and the observed nuclide, and on the
93 gyromagnetic ratio γ of the NMR nuclide. In contrast, the pseudocontact shift mechanism
94 involves the direct, through-space magnetic dipolar coupling between the unpaired electron spin
95 and the nuclear spin, and hence falls off as the inverse cube of their separation. This effect on the
96 NMR resonance frequency is controlled by the anisotropy of the magnetic susceptibility tensor
97 for the paramagnetic center, which can vary widely in magnitude and symmetry depending on
98 the local structure, and is not averaged away by magic angle spinning. This tensor can sometimes
99 be measured by EPR spectroscopy; it may also be closely related to crystal field effects in optical
100 spectroscopy (e.g. George et al. 2013).

101 In some cases, the predominant shift mechanism can be deduced from geometric
102 considerations, areas and numbers of shifted peaks and approximate estimations (Grey et al.
103 1989; Grey et al. 1990; Palke and Stebbins 2011a,b); as noted above there has been considerable
104 recent progress towards more quantitative calculation for transition metal phosphates and oxides.
105 For a given site in a mineral, both may be important. However, in spite of such theoretical
106 advances, prediction of such shifts remains challenging, and little is known about what to expect
107 in silicates where paramagnetic components are predominant instead of dilute, beyond the
108 obvious guesses that shifts will be large and peak broadening may be severe. The much lower
109 resonant frequency (due to the lower gyromagnetic ratio γ) of ^{29}Si relative to better-studied ^7Li
110 and ^{31}P (Stebbins and Xue 2014), and even more importantly its much lower natural isotopic

111 abundance (4.7% vs. 93% and 100% respectively) will both lower sensitivity (i.e. obtainable
112 signal to noise ratio) significantly; however the lower γ of ^{29}Si will also reduce the strength of
113 electronic-nuclear dipolar couplings, which can be a major contributor to line broadening in
114 paramagnetic materials.

115 An obvious and important first step in the better understanding of the effects of
116 paramagnetic cations in ^{29}Si NMR spectra of transition-metal rich silicate minerals, and hence
117 potentially a more wide applicability of this method to large groups of natural minerals and
118 technological materials, will be to determine whether such spectra for pure-phase end member
119 compositions are in fact observable and whether they suggest that useful structural information
120 can be obtained. Cu^{2+} offers an intriguing, and as-yet unexplored, entry point, as the single
121 unpaired electron in this cation may reduce the interaction and line broadening relative to the
122 effects of other divalent transition metals, e.g. the four unpaired spins for high-spin Fe^{2+} . At the
123 same time, the coordination environment of Cu^{2+} in most silicates (square planar or square planar
124 with two more distant neighbors such as H_2O) is distinct from more common octahedral
125 environments more commonly found for Fe^{2+} , etc. Here it is demonstrated for the first time that
126 useful ^{29}Si NMR spectra can indeed be observed for a series of natural and synthetic Cu^{2+}
127 silicates, and that these are sensitive to local structure and composition, although all of the data
128 cannot yet be interpreted in detail. ^{29}Si NMR can also provide new insights into the structure of
129 the common, but problematical Cu^{2+} silicate chrysocolla, which is often X-ray “amorphous,” but
130 is considered by some researchers to be a “mineraloid” with considerable short-range order and
131 consistent stoichiometry (Farges et al. 2007; Frost et al. 2012).

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134 **SAMPLES AND EXPERIMENTAL METHODS**

135 Natural mineral samples were from the Stanford University Research Mineral Collection,
136 (Tab. 1), with identities confirmed by powder XRD. Electron microprobe analyses showed that
137 both diopside samples had nearly ideal composition, although a minor calcite impurity was
138 detected in some fragments of the Altyn Tyube sample. The shattuckite samples were also of
139 near-ideal composition, with only minor quartz impurities. The soft, fibrous nature of the
140 planchéite resulted in a relatively poor polish, but EPMA data gave a systematically low atomic
141 Cu/Si ratio of about 0.85 (vs. 1.0 ideal), probably because of the presence of about 1.5 wt% K₂O,
142 0.5% Na₂O and 3% MgO, suggesting possible K⁺, Na⁺ and Mg²⁺ substitution on the M sites of
143 this amphibole-like structure (Evans and Mrose 1977). EPMA and optical microscopy also
144 revealed that some fragments of this material had minor amounts of coarse particles of an iron
145 oxide phase, probably hematite. Most of this contaminant was removed from the NMR sample
146 by hand picking. Both chrysocolla samples were massive and uniformly blue-green in color, with
147 earthy lusters and without obvious color banding or other heterogeneities.

148 Some of the diopside was dehydrated by heating a powder at 600 °C for 1 h, which gave
149 the expected weight loss and yielded a black material as previously reported (Law et al. 2010)
150 with an XRD pattern as in previous studies (Breuer and Eysel 1989). Synthetic cuprorivaite
151 (“Egyptian blue” pigment) was synthesized in air from a stoichiometric mixture of CuO, CaCO₃,
152 and amorphous SiO₂ with 1 wt% Na₂CO₃ as a flux, heated for 23 h at 1000 °C (Bloise et al.
153 2016; Chakoumakos et al. 1993; Pradell et al. 2006). (The flux is helpful in allowing more
154 complete reaction at the relatively low synthesis temperatures needed to maintain Cu in the
155 divalent state.) The product was the expected deep blue color. EPMA indicated that about 95%
156 of the material formed coarse grains that had the correct stoichiometry, with the remainder of the

157 sample present as unreacted silica, Ca-rich silicates or Na-rich material remaining from the Na
158 fluxing process. Powder XRD on this material detected only cuprorivaite (Chakoumakos et al.
159 1993) and traces of cristobalite; optical microscopy indicated that much of the non-cuprorivaite
160 material was amorphous, presumably residual glass.

161 NMR spectra were collected with a Varian 400 MHz spectrometer, at 79.5 MHz for ^{29}Si ,
162 with 3.2 mm MAS rotors, spinning at 10 to 20 kHz or non-spinning ('static'). Frequencies are
163 referenced to TMS at 0 ppm. To observe broad peaks a spin-echo pulse sequence ($90^\circ - \tau - 180^\circ$)
164 was used with the delay time τ of 50 to 100 μs to match an integral number of rotor periods and
165 to shift most of the NMR signal away from residual probe "ringing", an instrumental artifact that
166 is a source of what is essentially additional noise immediately after the last transmitted observe
167 pulse, in this system lasting a few 10's of μs . A 90° pulse length of 3 μs led to a relatively small
168 excitation bandwidth (ca. 150 kHz), and required moving the transmitter frequency by up to
169 several 1000's of ppm to initially locate the resonances and then to accurately observe spectra.
170 For the main NMR peaks of the Cu-silicates, relaxation was very rapid and allowed quantitative
171 spectra to be obtained with pulse delays of 0.1 s and total acquisition times of 3 to 20 h. Signals
172 for minor diamagnetic impurities relaxed more slowly and were observed with longer pulse
173 delays; their spectral components may be under-represented when not fully relaxed. Peak areas,
174 normalized to the nominal amount of SiO_2 in each sample, were ratioed to that of a pure natural
175 diopside (Wakefield, Quebec), the latter observed with a pulse delay of 1 h to ensure full
176 relaxation. Temperatures due to air frictional heating were measured from the ^{207}Pb NMR peak
177 shift in $\text{Pb}(\text{NO}_3)_2$ (Takahashi et al. 1999) as $31 \pm 2^\circ\text{C}$ at 10 kHz and $57 \pm 2^\circ\text{C}$ at 20 kHz. Even this
178 small difference produced large changes in positions of peaks with high paramagnetic shifts;
179 thermal gradients in rotors probably also increased the widths of the narrower peaks.

180

RESULTS AND DISCUSSION

181 **Diopside**

182 Table 1 gives NMR peak maxima, widths (full width at half maximum, FWHM), and
183 normalized peak areas. Diopside ($\text{Cu}_6\text{Si}_6\text{O}_{18} \cdot 6\text{H}_2\text{O}$) has a simple, well-known structure with
184 rings of six SiO_4 tetrahedra and a single Si site (with two bridging and two non-bridging
185 oxygens, or “Q²”), each with four closest Cu^{2+} neighbors in square planar coordination with two
186 more distant H_2O molecules (Ribbe et al. 1977). As in shattuckite and planchéite (below), each
187 Cu^{2+} shares at least one oxygen with another Cu^{2+} cation, forming continuous copper oxide
188 chains. The single ^{29}Si NMR peak of diopside (+980 ppm at 20 kHz spinning rate, ca. 57 °C) is
189 shifted far outside of the known range for diamagnetic silicates, i.e. -60 to -120 ppm (Fig. 1).
190 Effects of spinning rate on sample temperature, and hence on peak position, are well illustrated
191 by this mineral (Fig. 2) and are as expected for a large paramagnetic shift (Palke and Stebbins
192 2011). For these and other samples, the effects of temperature (T) on the shift scale with the
193 magnitude and sign of the shift as expected, and extrapolate roughly to the normal diamagnetic
194 chemical shift range when extrapolated to $1/T = 0$. Diopside samples from two different localities
195 had nearly identical spectra. Peak widths are among the narrowest observed for these copper
196 silicates (2 to 4 kHz or 25 to 40 ppm, in contrast to <1 ppm for ordered, diamagnetic silicates),
197 but are clearly affected by thermal gradients in the rotors, especially at the faster spinning
198 speeds. When data are collected with the spectrometer frequency centered on the resonance,
199 integrated peak areas (including spinning sidebands) are within 95% of values expected by
200 intensity calibration with the diopside standard, showing minimal loss of signal during the 50 to
201 100 μs echo delay. This in turn indicates that the spin-spin relaxation time T_2 is much larger than
202 the echo delay. The sidebands, especially when made more prominent by slower spinning (Fig.

203 2), map out the pattern observed without spinning, with an additional shift with temperature
204 decrease from about 31 °C at 10 kHz to about 24 °C without spinning. The static (non-spinning)
205 line shape resembles that produced by chemical shift anisotropy in a diamagnetic silicate,
206 indicating an orientation dependence to at least one part of the paramagnetic coupling (probably
207 the through-space dipolar interaction) between the unpaired electrons of Cu²⁺ and the ²⁹Si
208 nuclear spins.

209 The spectrum for dehydrated diopside (Cu₆Si₆O₁₈, not the metastable CuSiO₃ chain
210 silicate that can be produced at higher temperatures) is slightly broader with a paramagnetic shift
211 about 85 ppm lower, but both MAS and static spectra are overall quite similar to those of the
212 original, hydrated material: this change in position is relatively small given the >1600 ppm total
213 shift range observed for copper silicates (Fig. 2). Dehydrated diopside has been extensively
214 studied because of its unusual low temperature magnetic properties (Law et al. 2010) and, apart
215 from the loss of the H₂O molecules, has a crystal structure nearly identical to that of the hydrous
216 phase, with slightly shorter Cu-O and slightly longer Si-non-bridging oxygen bonds. Although
217 the color change from brilliant green to black on dehydration does suggest an important change
218 in the electronic structure around the Cu²⁺ ion, this change apparently does not strongly affect
219 either the crystal structure or the paramagnetic shift in the ²⁹Si NMR spectrum, probably because
220 in the hydrated material the H₂O molecules interact only weakly with the SiO₄ groups, but do
221 add two distant oxygens to the CuO₄ groups to form stretched “4+2” octahedra.

222 Detailed studies of Li-rich transition metal phosphates (e.g. LiMPO₄, with M=Fe²⁺, Ni²⁺,
223 Co²⁺, most commonly with the olivine structure) have shown that paramagnetic shifts for their
224 ³¹P spectra are dominated by through-bond Fermi contact shifts (Middlemiss et al. 2013;
225 Strobridge et al. 2014), primarily (but not exclusively) from the first neighbor cations, and that

226 the observed total shift can be well approximated by the sum of observed and/or calculated shifts
227 for each individual M-O-P bond pathway. The electronic structure of the paramagnetic cation in
228 its site (or sites) in the structure is critical; the number of unpaired electron spins is important
229 but, surprisingly, not always predominant in the magnitude of the shift. The magnitude and even
230 the sign of the Fermi contact shift for a given bond pathway depend sensitively on the degree of
231 covalency and orbital overlap, which can be reflected in the bond distances and especially in the
232 M-O-P bond angles, in turn affected by cation coordination number and geometry. This complex
233 picture is likely to apply to transition metal-rich silicates as well, as the non-bridging oxygens on
234 the SiO₄ groups, each coordinated by one or more paramagnetic transition metal cations, are at
235 least roughly analogous to those in the PO₄ groups of the phosphates.

236 The observed paramagnetic shift for diopside (about +980 ppm at ca. 57 °C) might thus
237 be approximated as resulting predominantly from the sum of through-bond electron spin transfer
238 along four distinct bond pathways to the four first neighbor Cu²⁺ cations, each having one
239 unpaired electron spin. Two such pathways go through each of the two non-bridging oxygens in
240 the structure (O2 and O3). Cu-O and Si-O distances along each fall within a small range, as do
241 the four Cu-O angles, which vary from 119 to 132° (Ribbe et al. 1977). It is also important to
242 note that in contrast to the more regular octahedral coordination of other transition metal cations
243 in well-studied Fe-, Mn-, Co-containing battery materials (Li-rich oxides and phosphates), the
244 large geometric anisotropy of the CuO₄ groups (which may also have two additional distant
245 apical oxygens in some Cu silicates) is likely to generate a large magnetic anisotropy as well.
246 This could contribute to significant through-space pseudo-contact shifts in addition to the
247 surmised large, through-bond Fermi contact shift. At this time, however, it is not possible to
248 interpret these new results on copper silicates in more detail.

249

250 **Shattuckite**

251 Shattuckite ($\text{Cu}_5(\text{SiO}_3)_4(\text{OH})_2$), can, like diopside, be obtained in coarse, monomineralic
252 form and has a well-determined structure (Evans and Mrose 1977). Its single-chain structure has
253 two distinct Q^2 Si sites in equal populations (Fig. 3). Si1, which resembles the Si site in diopside,
254 has four CuO_4 closest neighbors, two for each non-bridging oxygen with Cu-O-Si angles of 109
255 to 121°. Si2, in contrast, has three close (Cu-O-Si of 126 to 128 °) and two distant CuO_4
256 neighbors (Cu-O-Si of 110 and 132°). Two relatively narrow NMR peaks are observed (Fig. 1),
257 with equal areas each comprising nearly 50% of that expected from the diopside standard. Again,
258 both shifts for this phase are far outside of the known range for chemical shifts, one above, one
259 below. The spectrum shown in Fig. 1 was collected with the transmitter frequency halfway
260 between the two resonances, which substantially reduced the intensities of both because of
261 limited excitation bandwidth. Data for area quantification were collected with the frequency
262 centered on each peak separately; the latter spectra also showed more symmetric spinning
263 sideband patterns, resembling those observed for diopside. The apparently asymmetric sideband
264 patterns for both peaks in the shattuckite data in Fig. 1 are not the consequence of an especially
265 anisotropic shift interaction, but instead illustrate the drop-off in intensity away from the center
266 of the spectrum, caused by the limited excitation bandwidth of the echo pulse sequence used.
267 The higher frequency peak in shattuckite is significantly broader than the low frequency peak,
268 and than the peak in diopside. Whether this is a consequence of disorder (e.g. in H positions) or
269 an effect of differential relaxation is not yet known.

270 The observed shift at +1580 ppm (peak A) can be tentatively assigned to Si1, whose large
271 positive shift is somewhat comparable to that in diopside, and that at -250 ppm (peak B) to Si2

272 because of the smaller number of short Cu-O-Si bond pathways in the latter, but further study
273 will be needed to confirm this hypothesis. Note that both positive and negative paramagnetic
274 shifts can be induced by a given paramagnetic cation, depending on the details of orbital
275 geometry, electron distributions in bonds (Middlemiss et al. 2013), and the relative roles of
276 Fermi contact shifts and, potentially, of pseudocontact interactions. OH⁻ groups in this phase
277 (and planchéite, below) may also significantly alter electronic distributions, if they change the
278 covalency of other Cu-O bonds linked to Si and thus extent of the electronic spin transfer.

279

280 **Planchéite**

281 The structure of planchéite (Cu₈(Si₄O₁₁)₂(OH)₄ · xH₂O, x ≈ 0.43) is considered to be less
282 precisely known than, but related to, that of shattuckite, with some difficulty in the structure
283 determination due to the fine grained, fibrous nature of the mineral (Evans and Mrose 1977). The
284 material studied here, a massive, uniformly colored greenish blue material from the copper
285 mining district in Guchab, Namibia, was comprised of a dense aggregate of fibrous material and
286 had a powder XRD pattern that matched that in the JCPDS data base, from data in Evans and
287 Mrose (1977). EPMA data indicated, however, that up to about 10% of the Cu²⁺ was replaced by
288 Mg²⁺, K⁺ and Na⁺, potentially increasing the disorder in the structure. The published structure
289 has a double silicate chain, and 4 inequivalent Si sites, with equal proportions of Q² (Si1, Si4)
290 and Q³ (Si2, Si3) tetrahedra. The former appear to each have four CuO₄ neighbors as in diopside,
291 the latter probably with two Cu neighbors each. Initially only the NMR peak labeled “B” was
292 observed (Fig. 1), paramagnetically shifted to +130 ppm but with a much broader line (19 kHz)
293 than that for diopside or shattuckite. Sample spinning rate was insufficient to narrow this
294 completely. After testing a range of transmitter frequencies, a second, even broader (≈70 kHz)

295 peak “A” was detected centered at about 1200 to 1400 ppm. It is likely that some type of
296 disorder or poor crystallinity in this fibrous, fine-grained mineral contributes to broadening of
297 both peaks. The spectrum shown in Figure 1 was collected with the transmitter frequency
298 midway between the two peak maxima. By analogy to the spectra for shattuckite, peak A can be
299 tentatively assigned to the Q² sites (more Cu neighbors, strongly total magnetic interaction) and
300 peak B to the Q³ sites, but again further study will be required to confirm this.

301 With the standard echo delay of 100 μs, the overall integrated area (both peaks) is about
302 74% of that expected from the diopside standard. This value was only slightly higher (76%) at
303 the minimum practical echo delay of 50 μs (Fig. 4), indicating that the “missing” observed
304 intensity is not solely the result of signal decay (spin-spin relaxation) during the pulse sequence,
305 but is more likely dominated by insufficient excitation bandwidth. Given peak overlap and
306 unconstrained peak shapes, fitting of the spectra to obtain relative areas has low accuracy, but the
307 broader “A” peak could plausibly account for half of the total as expected from our tentative
308 assignment to the Q² sites in this structure.

309 In the planchéite spectra, a small, much narrower third peak (“D”) was also detected,
310 whose position may be consistent with a small or negligible paramagnetic shift, suggesting the
311 presence of a minor unknown impurity phase undetected by XRD or EPMA.

312

313 **Cuprorivaite (“Egyptian blue”)**

314 Among the known anhydrous copper silicates, cuprorivaite (CaCuSi₄O₁₀) appears to be
315 relatively straightforward to synthesize. It is the main constituent of the “Egyptian blue” pigment
316 invented millennia ago, apparently as a simulant for the prized lapis lazuli gemstone (Pabst,
317 1959), and also occurs rarely in nature.

318 The structure of cuprorivaite has corrugated silicate sheets with a single Q³ Si site
319 (Chakoumakos et al. 1993), CuO₄ square planar coordination and only a single Cu²⁺ first
320 neighbor to each Si (Fig. 3). The relatively large paramagnetic shift of +720 ppm (Fig. 1) is thus
321 somewhat surprising, but this phase does have underbonded non-bridging oxygens and a
322 resulting unusually short Si-NBO distance of 1.59 Å. This should contribute to a greater through-
323 bond transfer of unpaired electron spin density from the adjacent Cu²⁺ cation.

324 The spectra for cuprorivaite also have a small peak centered near to -92 ppm, which
325 required pulse delays of up to 10 s for full relaxation. Under these conditions, its calibrated peak
326 area indicates that it represents about 2% of the total Si in the sample, and could be due to a
327 small amount of a Ca-silicate impurity. The minor cristobalite impurity detected by EPMA and
328 XRD could relax even more slowly and be nearly undetectable by NMR in low concentrations;
329 in contrast, any Cu-rich glass phase that is present (again, suspected at the few % level from
330 EPMA data) could be unobservable in these ²⁹Si NMR experiments because of extreme
331 broadening caused by structural disorder and the accompanying widely varying Cu-Si
332 interactions.

333 For the main (ca. +720 ppm) peak for the cuprorivaite, the low integrated peak intensity
334 (in the initial spectrum with a 100 μs echo delay) of only about 6 % (Tab. 1) of the expected
335 value cannot be accounted for by incomplete reaction or the presence of impurities, which are
336 estimated from optical microscopy and EPMA imaging to be less than about 5% of the total
337 sample. Experiments with longer pulse delays demonstrate that the main peak is fully relaxed at
338 the 0.1 s pulse delay used; spectra collected with the transmitter frequency centered up to at least
339 5000 ppm above and below this resonance showed no additional signals. The relatively modest
340 peak width (ca. 7 kHz) did not suggest extreme broadening or signal loss during acquisition.

341 However, data collected at three different echo delays (1, 2, and 3 MAS rotor periods or 50, 100,
342 and 150 μ s) did demonstrate a surprisingly rapid signal decay (Fig. 4). As a first approximation,
343 this decay (measured here as the ratio of the integrated peak area (M) to that expected from the
344 calibration with the diopside standard (M_0)) is expected to be exponential with the echo delay
345 time τ , such that $\ln(M/M_0) \approx -\tau/T_2$. The time constant is the spin-spin relaxation time T_2
346 (approximated as a constant) and is estimated from the slope to be about 35 μ s. When
347 extrapolated back to 0 in the echo delay, the predicted intensity is close to 100% of that
348 expected.

349 Unlike reduced observable intensities for planchéite, the rapid decay of the signal in
350 cuprorivaite is more likely to be connected to especially fast nuclear spin relaxation,
351 characterized by a short T_2 and probably by a short nuclear spin-lattice relaxation time $T_{1,n}$. In
352 materials with significant contents of paramagnetic centers, relaxation can often be dominated by
353 electron-nuclear interactions, where the *electronic* spin-relaxation time, $T_{1,e}$ plays a key role
354 (Bakmutov et al. 2009; Grey and Dupré 2004). If $T_{1,e}$ is much shorter than the inverse of the
355 nuclear Larmor frequency ω_L (i.e. the NMR observation frequency for a given nuclide),
356 coupling of the electronic and nuclear spins is inefficient, and relaxation rates are far above
357 minimum values and accompanying peak broadening is not severe. This is often the case for rare
358 earth cations in oxide materials (Grey et al. 1990). $T_{1,e}$ for paramagnetic transition metal cations
359 may be slower, given strong coupling to nuclear spins, faster relaxation, and potentially more
360 severe line broadening. At first glance for the copper silicates studied here, this problem might
361 be expected to be minimized for cuprorivaite, as each Si has only 1 Cu first cation neighbor, in
362 contrast to 3 or 4 for diopside, shattuckite, and planchéite, thus reducing the total number of
363 interactions between Cu and Si. However, cuprorivaite is distinctive in this group of minerals by

364 having only isolated CuO_4 groups, instead of continuous chains of edge or corner-shared Cu
365 polyhedra. Interactions among unpaired Cu^{2+} electronic spins are thus likely to be much weaker
366 in cuprorivaite than in the other three minerals. Given that electronic spin-lattice relaxation is
367 likely to be affected or even controlled by these electronic interactions, it is reasonable to
368 hypothesize that the $T_{1,e}$ for cuprorivaite might be relatively slow, promoting more rapid ^{29}Si
369 nuclear spin relaxation and signal loss during acquisition of the NMR signal. Further
370 measurements of T_2 , $T_{1,n}$ and $T_{1,e}$ in these types of materials by NMR and EPR methods could
371 prove fruitful for better understanding of such spin dynamics. Details of relaxation rates and
372 mechanisms in paramagnetic systems can also be structurally informative, for example in MnO-
373 containing microporous silica-based materials (Bakmutov et al. 2009).

374

375 **Chrysocolla**

376 Chrysocolla is a hydrated, hydroxyl-bearing copper silicate whose formula (neglecting its
377 minor Al content) is often given as approximately $\text{Cu}_2\text{H}_2\text{Si}_2\text{O}_5(\text{OH})_4 \cdot x\text{H}_2\text{O}$ (Hariu et al. 2013).
378 Its weak X-ray diffraction has indicated poor or even non-crystallinity, and thus a poorly defined
379 structure. Much of what we know about this potentially highly variable material thus has come
380 from spectroscopy (e.g. XPS, XAS), including reports on material from the localities of the
381 samples described here (Farges et al. 2007; Frost et al. 2012; Hariu et al. 2013). The goal here is
382 not to resolve all questions about this mineral or “mineraloid”, but to test the applicability of
383 paramagnetic ^{29}Si NMR to gain new types of information about its short-range structure, perhaps
384 as a guide to future studies. Our two samples are both from well-studied copper mining centers
385 (Tab. 1) and appear to be mineralogically uniform on microscopic examination, without obvious
386 heterogeneities, color banding, etc.

387 The ^{29}Si NMR spectra for both samples are nearly identical (Fig. 5), with peak maxima at
388 +380 ppm and linewidths of about 37 kHz (460 ppm); MAS at 20 kHz did not produce
389 significant line narrowing. About 80% of the expected signal is detected. Spectra from both
390 samples also have small, much narrower peaks centered at about -85 ppm with narrow and
391 broader components (line widths about 35 to 80 ppm). These relax more slowly than the main
392 peaks, and are within the range of normal diamagnetic chemical shifts. These may come from
393 small amounts of amorphous silica, probably with minor Cu contents that contribute to widths
394 greater than is typical in the absence of paramagnetic components. The proportions of these
395 impurities are not precisely determined, but peak areas suggest that they comprise less than 10 to
396 20% of the silica in the samples.

397 The relatively large paramagnetic shifts for both chrysocolla samples suggest that most of
398 their Si is closely associated with Cu^{2+} cations, with one or more in the first coordination shells.
399 The extreme peak widths are likely the result of disorder as expected in a poorly crystalline
400 material, combined with paramagnetic interactions (^{29}Si peak widths for diamagnetic silicate
401 glasses are typically about 10 to 25 ppm). However, the very similar overall peak positions and
402 shapes for the two samples are unlikely to be the result of a random physical mixture of some
403 type of copper compound (e.g. $\text{Cu}(\text{OH})_2$) that is dispersed in amorphous silica or other colloid
404 (Farges et al. 2007), but suggest instead a common stoichiometry and a short- to medium-range
405 structure that is similar in both specimens not only on average but in the overall distribution of Si
406 environments. Such a conclusion has also been reached in experimental studies of synthetic sol-
407 gel products, which although remaining X-ray amorphous, do have reproducible medium-range
408 order (Hariu et al. 2013). Future NMR studies of a wider range of chrysocolla and other
409 “mineraloid” or nano-crystalline silicates is likely to prove interesting.

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IMPLICATIONS

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New data presented here show that useful ^{29}Si NMR spectra can be observed for several Cu^{2+} silicates, suggesting that there may be a much wider range of applications of this method to a variety of paramagnetic minerals containing cations with unpaired electron spins. Major technical and theoretical challenges remain, however. In some cases, line widths are small enough to be narrowed significantly by magic angle spinning; in others higher resolution spectra will be obtainable with NMR probes with faster spinning rates, which can now reach at least 70 kHz in advanced designs. Rapid spin-lattice relaxation times, which allow for rapid pulsing in the NMR experiment, often allow spectra to be collected with good signal-to-noise ratios even though peaks may be orders of magnitude broader than in diamagnetic minerals. When carefully calibrated with a diamagnetic intensity standard, observed signals may be lower than expected because of very broad peaks and limited excitation bandwidth, or because of rapid signal decay during acquisition in the case of unusually rapid nuclear spin relaxation. Signals can be shifted hundreds to thousands of ppm above or below of the known range for diamagnetic silicates, requiring careful experimental design to even initially determine spectrometer transmitter frequency; temperature variations of 10's of $^{\circ}\text{C}$, induced by air frictional heating during sample spinning, can shift resonances by 100's of ppm. Although as yet we know little about the specific details of the structure that control the large paramagnetic shifts and line widths, these will likely involve the short to medium range relationships between the magnetic cations and Si sites. Potentially important new constraints on bonding, electronic structure, and order/disorder in both pure phases and solid solutions may be obtainable in future studies. Advanced theoretical work, as well as more complex NMR methods that facilitate the accurate detection of very broad NMR

433 resonances (Pell and Pintacuda 2015), will be especially important in taking full advantage of
434 this newly realized experimental tool.

435

436

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439 electron microprobe analyses. We thank two helpful reviewers and the associate editor for useful
440 comments on the manuscript.

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580
581
582

583 **Table 1.** Samples, ²⁹Si NMR data, and numbers of first neighbor transition metal cations. Site
 584 assignments are preliminary estimates. All data are from MAS spectra with 20 kHz spinning
 585 speed (≈57 °C), 0.1 s pulse delay, and 100 μs echo delay, unless otherwise noted.
 586

sample, locality, formula	peak	site/ phase	no. Cu ²⁺	peak max., ppm	width (FWHM) kHz	relative area ^a
diopase , SU#1431 Altyn Tyube, Kaz. Cu ₆ (Si ₆ O ₁₈)6H ₂ O	A	Si	4	980±20	3.4	0.97±0.05
“ 10 kHz spin rate	A	Si	4	1065±20	2.0	
“ dehydrated	A	Si	4	895±20	8.6	0.94±0.05
diopase , SU#7516 Guchab, Namibia	A	Si	4	980±20	2.8	0.96±0.05
shattuckite , SU#8247 Ajo, AZ Cu ₅ (SiO ₃) ₄ (OH) ₂	A	Si1?	4	1580±10	9.6	0.46±0.05
	B	Si2?	3	-250±10	4.0	0.49±0.05
planchéite , SU#7797 Guchab, Namibia Cu ₈ (Si ₄ O ₁₁) ₂ (OH) ₄ xH ₂ O	A	Si1+Si4?	4	1200±200	70 ^b	A+B+D:
	B	Si2+Si3?	2?	130±20	19 ^b	0.74±0.05
	D	? ^d	?	-120±5	2-4 ^c	
cuprorivaite , synthetic CaCuSi ₄ O ₁₀	A	Si	1	720±20	7.0	0.06±0.003 ^f 0.24±0.007 ^f
	C	? ^e	(0)	-92±5	1.5	0.02±0.01
chrysocolla , SU#8461 Katanga, Zaire ≈Cu ₂ H ₂ Si ₂ O ₅ (OH) ₄ xH ₂ O	A	Si	1-4?	380±20	37 ^b	0.7±0.1
	O	a-SiO ₂ ^g	(0)	-87±5	2-7 ^c	≈ 0.03
chrysocolla , SU#28475 Inspiration mine, Gila Ct, AZ	A	Si	2-4?	380±20	37 ^b	0.8±0.1
	O	a-SiO ₂ ^g	(0)	-87±5	2.8-7 ^c	≈ 0.03
“, 4 s pulse delay	A	Si	2-4?	380±20	37 ^b	0.8±0.1
	O	a-SiO ₂ ^g	(0)	-87±5	2.8-7 ^c	0.1±0.05

587 ^a Peak areas include any observed sidebands and are normalized to reported wt% silica, then
 588 ratioed to area for Wakefield diopside (SU#31011) standard. Values <1 indicate signal loss.

589 ^b Spinning speed insufficient for full narrowing.

590 ^c Composite peak with narrow and broader components.

591 ^d Possibly an unknown diamagnetic impurity.

592 ^e Probably Cu-bearing Ca-silicate impurity.

593 ^f Upper value is for echo delay of 100 μs, lower value is for echo delay of 50 μs.

594 ^g Probably amorphous silica.

595

596

597

Figure captions

598 **Figure 1.** ^{29}Si MAS NMR spectra for copper silicates (Tab. 1). Peaks labeled “A” and “B” are
599 for Si sites in the main phase; “C” may be from a Ca-silicate impurity; “D” may be due to an
600 unknown diamagnetic impurity. “*” marks spinning sidebands. The entire known range of
601 chemical shifts for SiO_4 groups in diamagnetic silicates lies between the two dashed lines. Data
602 shown for cuprorivaite were collected with an echo delay of 50 μs ; for others with 100 μs . For
603 planchéite and shattuckite, spectra shown were collected with transmitter frequency centered
604 between two main peaks; for cuprorivaite and diopside the transmitter frequency was centered on
605 the single main peak.

606

607 **Figure 2.** ^{29}Si NMR spectra for diopside and dehydrated diopside, with spinning speeds labeled.
608 The offset of the peak from 20 to 10 kHz (arrow) is caused by a reduction in temperature of
609 about 26 °C at the lower spinning speed. The complex MAS peak shapes for the hydrated sample
610 may be due in part to non-uniform temperature in the rotor. Note that the frequency scale is
611 expanded relative to that in Fig. 1.

612

613 **Figure 3.** Local structures around the Si sites in representative copper silicates. In shattuckite
614 ($\text{Cu}_5(\text{SiO}_3)_4(\text{OH})_2$), Si1 (a) and Si2 sites (b) are linked to form corner-shared chains; nearly
615 planar CuO_4 groups are edge and corner shared, also to form chains. Si1 has four close Cu
616 neighbors; Si2 has three. H positions are not shown. In cuprorivaite ($\text{CaCuSi}_4\text{O}_{10}$) (c), each Si
617 has three Si and only one CuO_4 neighbor, which is isolated from other Cu. Ca site is not shown.
618 (Color online.)

619

620

621 **Figure 4.** Natural logarithm of the observed, integrated ^{29}Si NMR signal intensity (M) for
622 cuprorivaite and planchéite, normalized to that expected from the pure diopside standard (M_0),
623 vs. the echo delay time in μs . For cuprorivaite, extrapolation back to zero delay gives a ratio
624 close to 1 (i.e. $\ln(M/M_0) = 0$), i.e. the full expected intensity. For planchéite, signal loss is less
625 severe but is not a strong function of delay time and does not extrapolate to the full expected
626 value.

627

628 **Figure 5.** ^{29}Si MAS NMR spectra for two chrysocolla samples (Tab. 1), with the same ppm scale
629 as Fig. 1. Peaks marked “O” are probably due to amorphous “opaline” silica. Pulse delays are
630 noted. The entire known range of chemical shifts for SiO_4 groups in diamagnetic silicates lies
631 between the two dashed lines.

632

FIGURE 1

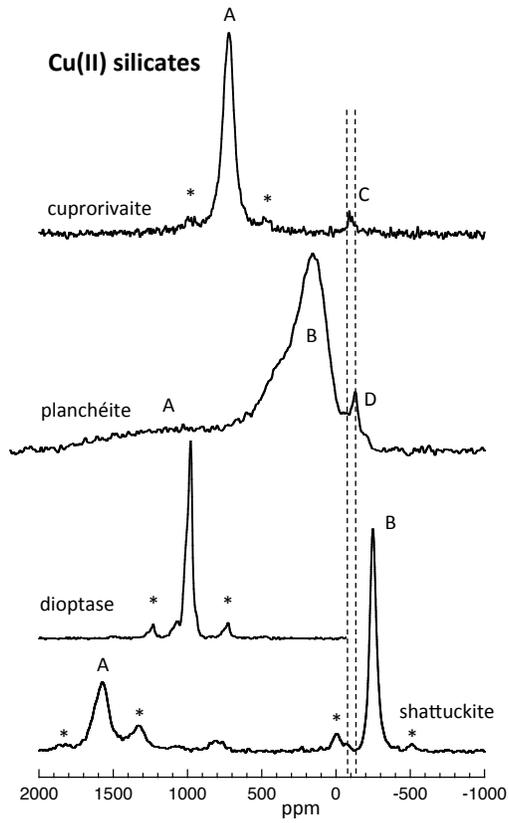


FIGURE 2

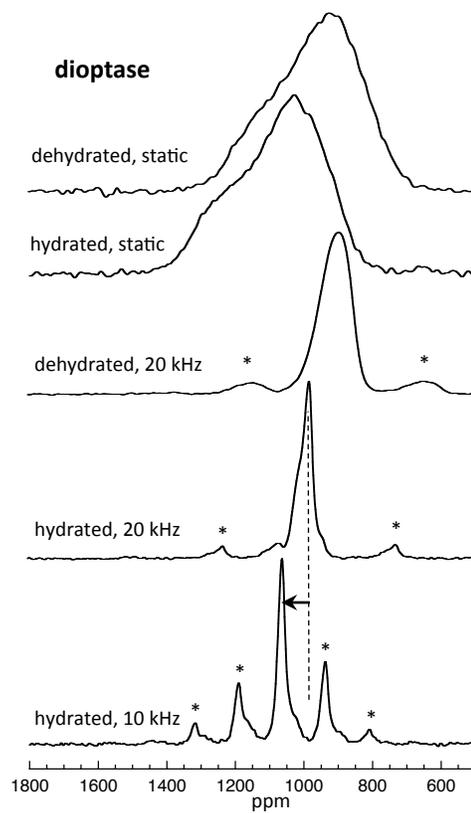


FIGURE 3

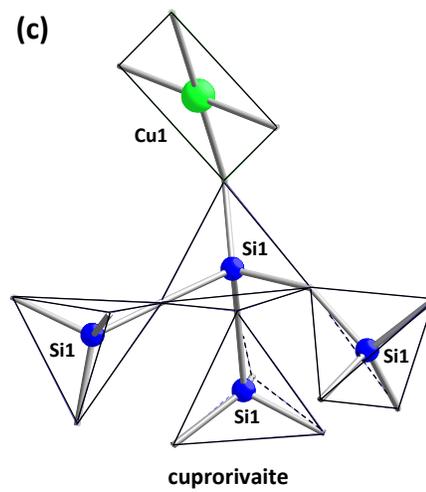
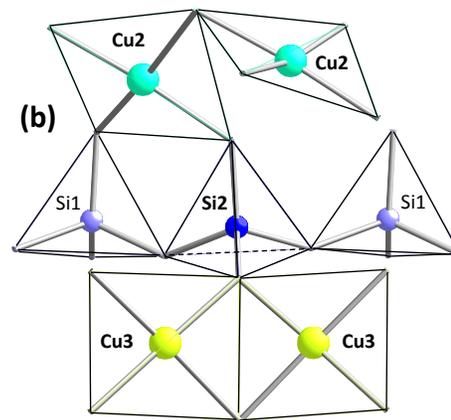
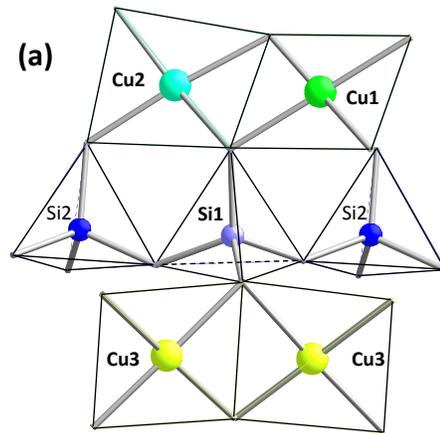


FIGURE 4

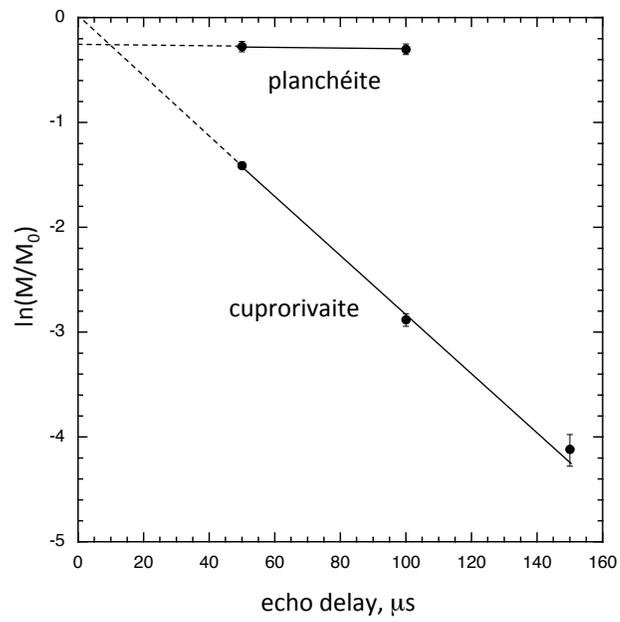


FIGURE 5

