

Supplementary Material: Appendix D

At equilibrium there is equality of chemical potentials (μ) between H₂O dissolved in the melt (m) and exsolved as a vapor (v)

$$\mu_{H_2O}^m = \mu_{H_2O}^v \quad (D1)$$

The corresponding free energy balance for this equilibrium is

$$0 = \Delta G_{(T,P)}^o + RT \ln \frac{f_{H_2O}^v}{a_{H_2O}^m} \quad (D2)$$

where $\Delta G_{(T,P)}^o$ is the standard state free energy change for the equilibrium represented by Eq. D1, $f_{H_2O}^v$ is the fugacity of H₂O vapour, and $a_{H_2O}^m$ is the activity of H₂O in the melt. Expansion of Eq. D2 in terms of standard state values of enthalpy ($\Delta H_{(T,P)}^o$) and entropy ($\Delta S_{(T,P)}^o$) yields:

$$0 = \Delta H_{(T,P)}^o - T\Delta S_{(T,P)}^o + RT \ln \frac{f_{H_2O}^v}{a_{H_2O}^m} \quad (D3)$$

The activity of H₂O in the melt can be approximated as the square of H₂O content measured, assuming all H₂O is speciated as OH at low pressure, low concentrations and high temperatures. This assumption is corroborated by Silver and Stolper (1988), Ihinger et al. (1999) and Zhang et al. (2007) and coincides with previous H₂O solubility models (Nicholls, 1980; Burnham 1994; Sahagian and Proussevitch 1996). Thus,

$$a_{H_2O}^m = [x_{H_2O}]^2 \quad (D4)$$

The fugacity of H₂O vapour is the fugacity for H₂O in the vapour at the experimental conditions ratioed to the standard state conditions: a pure ideal gas at atmospheric pressure (i.e. 101325 Pa). Assuming that the bubbles in our experiments contain pure H₂O fluid at the model internal bubble pressure (cf. Table 3):

$$f_{H_2O}^v = \frac{P_{H_2O}}{101325} = \frac{101828}{101325} \cong 1 \quad (D5)$$

Substituting these values in to equation D3 yields

$$0 = \Delta H_{(T,P)}^o - T\Delta S_{(T,P)}^o + RT \ln \frac{1}{[x_{H_2O}]^2} \quad (D6)$$

The corresponding equilibrium constant (K_1) is

$$K_1 = \frac{1}{[x_{H_2O}]^2} = [x_{H_2O}]^{-2} = \exp \frac{-\Delta H^o}{RT} \exp \frac{\Delta S^o}{R} \quad (D7)$$

This can be rearranged to form the Arrhenian equation:

$$\ln x_{H_2O} = \frac{\Delta H^o}{2RT} + \frac{-\Delta S^o}{2R} \quad (D8)$$

where slope (m) and intercept (b) are:

$$m = \frac{\Delta H^o}{2R} \quad (D9)$$

$$b = \frac{-\Delta S^o}{2R} \quad (D10)$$

We obtained estimates for ΔH^o and ΔS^o from fitting Eq. D8 to our experimental data over the temperature range 900 to 1050°C (cf. Fig. 6a).

References

- Burnham, C.W. (1994) Development of the Burnham model for prediction of H₂O solubility in magmas. *Reviews in Mineralogy and Geochemistry*, 30, 123-129.
- Ihinger, P. D., Zhang, Y., and Stolper, E. M. (1999) The speciation of dissolved water in rhyolitic melt. *Geochimica et Cosmochimica Acta*, 63, 3567-3578.
- Nicholls, J. (1980) A Simple Thermodynamic Model for Estimating the Solubility of H₂O in Magmas. *Contributions to Mineralogy and Petrology*, 74, 211–220.
- Sahagian, D.L., and Proussevitch, A.A. (1996) Thermal effects of magma degassing. *Journal of Volcanology and Geothermal Research*, 74, 19–38.
- Silver, L., and Stolper, E. (1985) A Thermodynamic Model for Hydrous Silicate Melts. *Journal of Geology*, 93, 161–177.
- Zhang, Y., Xu, Z., Zhu, M., and Wang, H. (2007) Silicate melt properties and volcanic eruptions. *Reviews of Geophysics*, 45, RG4004.